

## 14. CORRELATION OF CALCAREOUS NANNOPLANKTON AND PLANKTONIC FORAMINIFER BIOSTRATIGRAPHY OFF THE COAST OF SOUTHERN CHILE<sup>1</sup>

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### ABSTRACT

The biostratigraphy established for ODP Leg 141 Sites 859 through 863, drilled off the coast of southern Chile, is based mainly on calcareous nannoplankton and planktonic foraminifers. Because of the scarcity of the fossils and the low diversity of the assemblages, the biostratigraphic resolution is quite low. The uppermost sediments contain well-preserved fossil assemblages that indicate warm surface water influences at the southern mid-latitudes. The oldest sediments may belong to the uppermost lower Pliocene. The Pliocene/Pleistocene boundary is determined by the first occurrence of the *Gephyrocapsa oceanica* group at 1.6 Ma.

### INTRODUCTION

Five sites were drilled during the Ocean Drilling Program (ODP) Leg 141 on the Chile margin in the vicinity of the Nazca/Antarctica/South America Plate Triple Junction (Fig. 1). Sites 859, 860, and 861 are located on a west-east transect across the accretionary complex at different water depths. Site 862 was drilled near the crest of the Taitao Ridge of oceanic origin, and Site 863 was drilled at the base of the trench slope of the Chile Trench at the point where the Chile Ridge is being subducted.

The biostratigraphy for these sites is based mainly on results from calcareous nannoplankton and planktonic foraminifer investigations obtained from the same samples.

In general, the foraminifers, the nannoplankton assemblages as well as the siliceous microfossils are of low species diversity as a result of low water temperatures, dissolution, and dilution by detrital material. Thus, the biostratigraphic resolution is poor. Observed fluctuations in the abundance and diversity of the fossil assemblages may be related to climatic changes.

### METHODS

Samples were taken from the core catchers and about one sample per section of core. The preparation methods used to obtain foraminifers differed depending on the degree of sediment induration. Soft sediments were washed over a 63 µm screen. Slightly indurated sediments were soaked in diluted H<sub>2</sub>O<sub>2</sub> solution and then washed. Indurated samples were freeze-dried, soaked in hot paraffin, and boiled in a soda lye until disintegration occurred. This method was repeated until a reasonable amount of loose particles was obtained. After the washed samples were dried, the fossils were separated under the binocular microscope. The planktonic/benthic ratios were calculated. The state of preservation of planktonic foraminifers is described as follows: G = good (little or no fragmentation, overgrowth, and/or dissolution); M = moderate (some signs of fragmentation, overgrowth, and/or dissolution); P = poor (severe fragmentation, heavy overgrowth, and/or dissolution). Samples with reworked foraminifers sometimes in different colors, translucent-white and yellow-orange, are marked in the tables as *c*. Deformed specimens of foraminifers are

indicated as *d* and partly silicified ones as *s*. The abundances of foraminifers and nannofossils given in the range charts are based on semiquantitative estimate (B = barren, T = traces, R = rare, F = few, C = common, A = abundant). Smear slides were prepared for the investigation of calcareous nannoplankton. The samples were studied using a polarizing microscope.

### BIOSTRATIGRAPHY

#### Calcareous Nannoplankton

Stratigraphic results of Leg 141 were obtained by the investigation of calcareous nannofossils from about 550 samples. Most of the samples are barren or nannoplankton poor. Age determinations are based on the Standard Nannoplankton Zonation (Martini, 1971; Martini and Müller, 1986).

The nannoplankton assemblages are of low diversity owing to cold-water temperatures, dissolution of the calcareous microfossils, and dilution by the high amount of detrital material. Only a few mainly long-ranging species are present, but not enough to establish a detailed biostratigraphy.

The Pleistocene sequence can be subdivided broadly into upper and lower Pleistocene by using the Zone NN19/NN20 boundary at 0.47 Ma. The Pliocene/Pleistocene boundary, normally determined by the extinction of the discoasters, is difficult to recognize because this group, typical of a warm water environment, is absent from the Leg 141 material. The last occurrence of *Cyclococcolithus macintyreii*, proposed by Bizon and Müller (1977) for the determination of the Pliocene/Pleistocene boundary at 1.6 Ma, is also not a useful biostratigraphic event within this region. *Cyclococcolithus macintyreii* found also in areas dominated by lower surface-water temperatures is generally a good substitute for the discoasters. However, only rare specimens of this species were observed in the studied samples. Because of these difficulties, we decided to use the first occurrence (FO) of the group *Gephyrocapsa oceanica* (Raffi and Rio, 1979), a *Gephyrocapsa* with a large central opening.

A subdivision of the upper Pliocene in Leg 141 sediments (nannoplankton Zones NN16–NN18) is impossible. At some sites, small specimens of the genus *Reticulofenestra* were found. But by using only the light microscope, it is not possible to determine if this form represents a small variety of *Reticulofenestra pseudoumbilica*, which would give an early Pliocene age (Zone NN15), or if it belongs to part of Zone NN16, because it has been observed from this stratigraphic level in other regions.

It is most likely that the entire recovered sequences belong to the Pleistocene and upper Pliocene; in any case, they are not older than 3.2–3.4 Ma. This is shown by the presence of a very small species of the genus *Gephyrocapsa*, which has its first appearance at about this

<sup>1</sup> Lewis, S.D., Behrmann, J.H., Musgrave, R.J., and Cande, S.C. (Eds.), 1995. *Proc. ODP Sci. Results*, 141: College Station, TX (Ocean Drilling Program).

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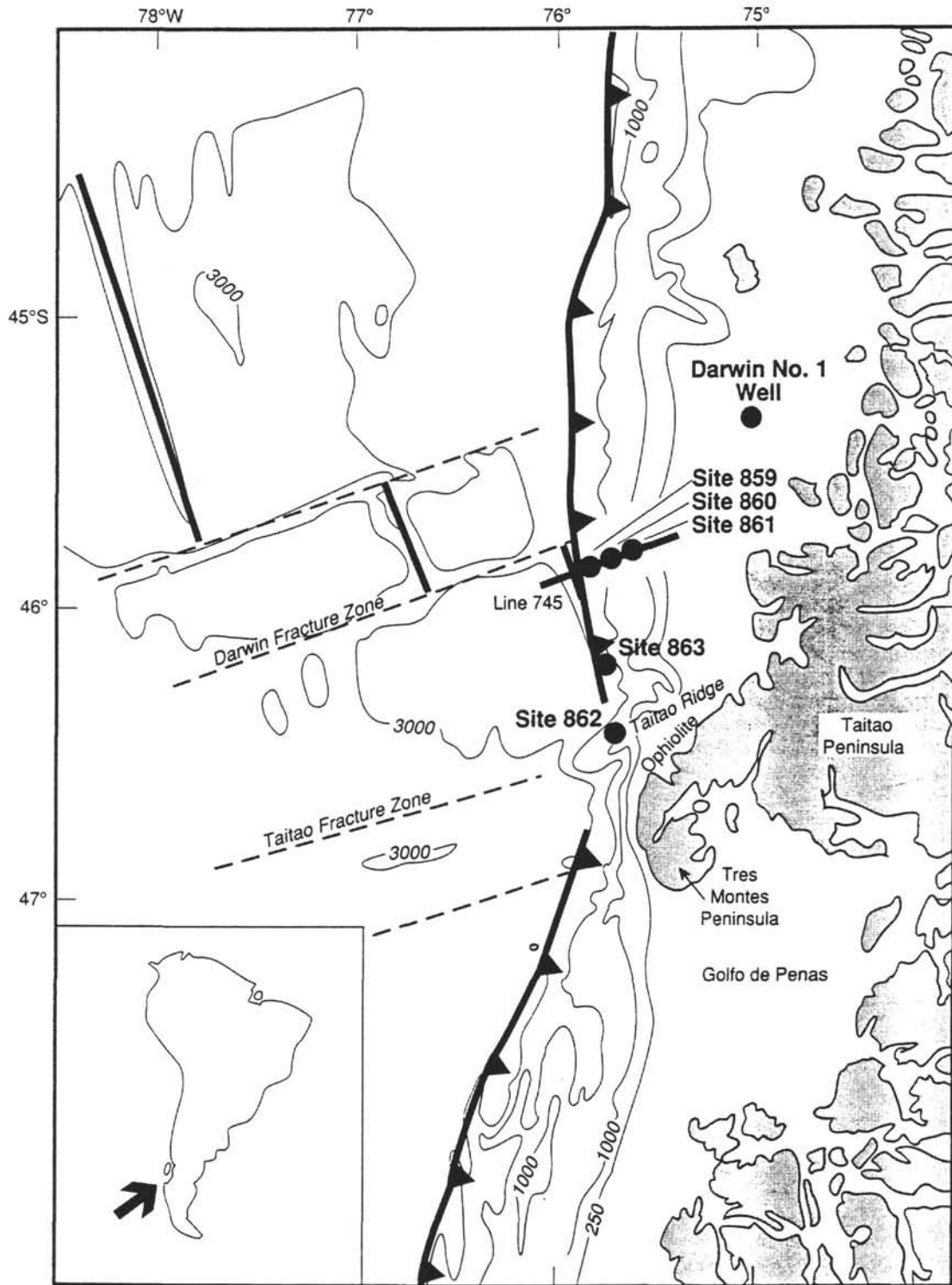


Figure 1. Location map of Sites 859-863.

time. Further on, *Pseudoemiliana lacunosa* was commonly observed, and this species is unknown from sediments older than Zone NN15.

Nannofossils are apparently more common in the sediments of the lower part of Zone NN16 indicating decreasing water temperature during the late Pliocene. Fluctuations in abundance throughout the Pliocene sequence were observed at Sites 860 and 861.

Reworked species from older strata (Eocene and lower Miocene) occur rarely throughout the series. Imbrications of the sedimentary

sequences certainly exist. However, it was not possible for us to recognize them by biostratigraphic evidence.

### Planktonic Foraminifers

Stratigraphic results and paleotemperature estimations for Leg 141 were obtained by the investigation of planktonic foraminifers from 653 samples. Most of them are poor in planktonic foraminifers,

and only the uppermost sequences contain richer assemblages. To acquire a sufficient number of specimens, the entire residue was usually picked and to get a comparable value in abundance, the number of foraminifers in 20 cm<sup>3</sup> of sediment was calculated.

Temperature is one of the habitat-controlling factors of foraminifer distribution in modern oceans. The taxa of planktonic foraminifers can be divided roughly into warm-water assemblages and cold-water forms. The majority of the Holocene planktonic foraminiferal species lives in tropical and subtropical waters, whereas only a single species, *Neogloboquadrina pachyderma* (sinistral coiling form), lives in polar waters. The regions where warm-water and cold-water assemblages overlap in distribution, and where the greatest faunal contrast occurs, are designated the Transition Zone (Bé and Tolderlund, 1971). The investigated area falls in the Southern Transition Zone of recent planktonic foraminifers and is bounded and influenced by the Subtropical Zone in the north and by the Subantarctic Zone in the south. *Gobocoinella inflata* is the most common taxon in Leg 141 sediments. It appears to be indigenous and therefore a good indicator species of the Transition Zone. However, it is possible to recognize the *Truncorotalia truncatulinoides* Zone, the *Gobocoinella inflata* Zone, and the *Globorotalia crassaformis* Zone, respectively, the *Neogloboquadrina pachyderma* (sinistral coiling form) Zone. The FO of *Truncorotalia truncatulinoides* is strongly influenced by temperature, and, therefore, the base of this zone varies. The zonations used in this paper are summarized in Figure 2. Location, position, water depth, and penetration of each hole are given in Table 1.

**Site 859**

Site 859 is the westernmost site of the east-west profile of Sites 859 through 861, located most seaward at the deepest water depth.

**Calcareous Nannoplankton**

Cores 141-859A-1H and -2H belong to the uppermost Pleistocene (Zone NN21). Nannofossils are few to common (Table 2). The assemblages consist of following species: *Helicosphaera carteri*, *Cyclococcolithus leptoporus*, *Gephyrocapsa ericsonii*, *G. oceanica*, *Syracosphaera pulchra*, and *Emiliania huxleyi*. Reworked taxa from older strata are rare.

The entire underlying sequence is dated late Pliocene (Zones NN16–NN18) and can not be further subdivided. This age is based on the presence of *Cyclococcolithus macintyreii*, *Helicosphaera sellii*, a very small *Gephyrocapsa* sp. and small *Reticulofenestra* cf. *pseudoumbilica*. Thus, a significant unconformity of more than 1 Ma exists.

The Pliocene sediments are extremely poor in nannofossils. Most of the samples taken from Hole 859B are barren because of dissolution (Table 3), which may be related to compaction of the sediments and the circulation of fluids within the accretionary wedge. Reworked specimens from older strata (mostly Miocene) were observed in several samples from throughout the entire series.

The biostratigraphic results obtained from Holes 859A and 859B show that the sediments are not older than 3.4 Ma, as indicated by the presence of the small *Gephyrocapsa* sp. and *Pseudoemiliania lacunosa*. That means that the upper Pliocene has a thickness of almost 470 m. Imbrication within this thick pile of sediments is highly probable, but it is not possible to prove this assumption because of the poor biostratigraphic resolution.

**Planktonic Foraminifers**

The biostratigraphic results obtained by the investigation of planktonic foraminifers are based on 44 samples from Hole 859A (down to 145.5 m below seafloor [mbsf]) and 93 samples (54.10 through 80.90 mbsf, and 141.20 through 476.10 mbsf) from Hole 859B. Foraminifers occurred in Hole 859A in 60% of the observed samples and in 45% of the samples from Hole 859B. No foraminifers were obtained from the micrite-cemented stones in Cores 141-859B-18R, -27R, -29R, -37R, and -38R. Sample 141-859A-9P-CC from

**Table 1. Location, position, water depth, and penetration of Leg 141 holes.**

Hole	Position	Water depth (m)	Penetration (m)
859A	45°53.761'S, 75°51.165'W	2741.2	146.5
859B	45°53.720'S, 75°51.330'W	2748.9	476.1
860B	45°51.972'S, 75°45.101'W	2145.9	617.8
861A	45°51.025'S, 75°41.531'W	1667.0	9.5
861B	45°51.025'S, 75°41.531'W	1667.0	9.5
861C	45°51.025'S, 75°41.531'W	1652.2	353.1
861D	45°51.008'S, 75°41.499'W	1652.2	496.3
862A	46°30.475'S, 75°49.603'W	1274.9	22.1
862B	46°30.475'S, 75°49.603'W	1268.6	42.9
862C	46°30.510'S, 75°49.566'W	1228.6	102.1
863A	46°14.210'S, 75°46.371'W	2564.2	297.3
863B	46°14.210'S, 75°46.371'W	2564.2	742.9

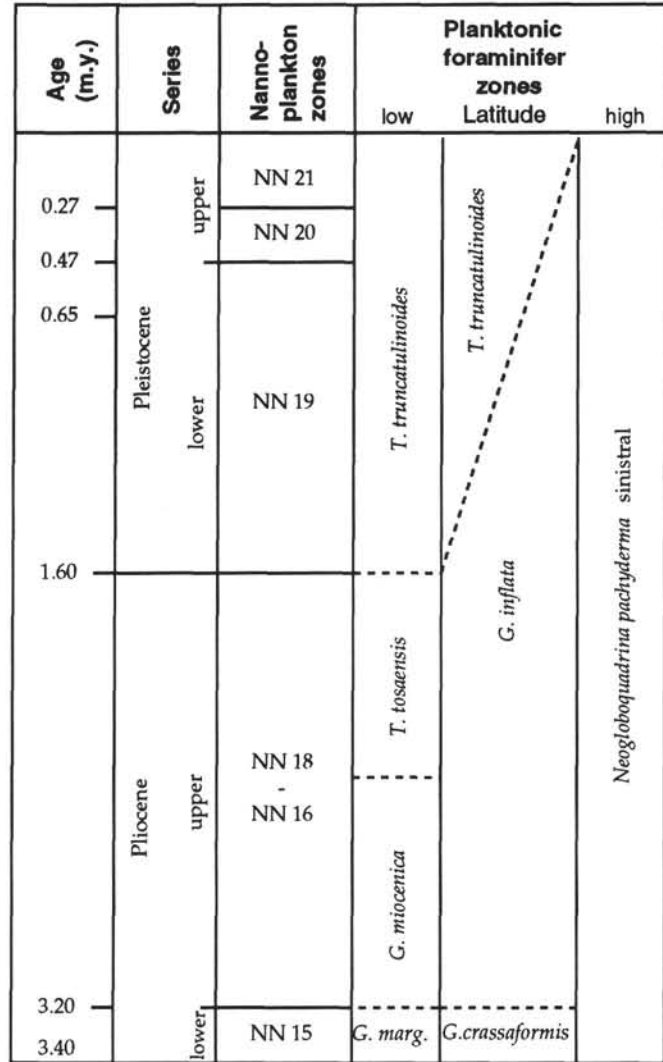


Figure 2. Correlation of nannoplankton and planktonic foraminiferal zones.

58.40 mbsf contains a mixed assemblage of taxa of the *Truncorotalia truncatulinoides* Zone together with some orange-stained *Neogloboquadrina pachyderma* (sin.) and is considered to represent in part downhole contamination.

At Site 859 two planktonic foraminiferal events to consider are the FO of *T. truncatulinoides* in Sample 141-859A-2H-CC and the FO of *G. inflata* in Sample 141-859B-29R-CC. Unfortunately, both events are followed downhole by assemblages dominated by mainly *N. pachyderma* (sin.). Consequently, these FO are not thought to



Core 141-860B-4H through Section 141-860B-17X-2 belong to the lower Pleistocene (Zone NN19), as shown by the presence of *Gephyrocapsa oceanica*, *G. caribbeanica*, and *Pseudoemiliana lacunosa*. Nannofossils are common only in Cores 141-860B-7H and -16X. All other samples of the lower Pleistocene sequences are poor in nannoplankton. These results show that the Pliocene/Pleistocene boundary lies about 60 m deeper than given in Behrmann, Lewis, Musgrave, et al. (1992).

Beside changes in the abundance of nannofossils, the assemblages are the same from Section 141-860B-17X-3 down to Core 141-860B-70X. This sequence is placed in the upper Pliocene (Zones NN16–NN18). It was not possible to confirm repetitions within the Pliocene sequence as evidenced by the study of planktonic foraminifers. However, repeated changes in abundance of nannofossils throughout the sequences may indicate imbrications and repetitions. Cores 141-860B-17X, -21X, to Section -22X-2, -31X, -32X, and -58X are rich in nannoplankton with following species: *Coccolithus pelagicus*, *Cyclococcolithus leptoporus*, *C. macintyreii*, *Gephyrocapsa* sp., *Pseudoemiliana lacunosa*, *Helicosphaera carteri*, *H. sellii*, *Discolithina japonica*, *Pontosphaera pacifica*, and the small variety of *Reticulofenestra* cf. *pseudoumbilica*. The more common occurrence of *Discolithina japonica* and *Pontosphaera pacifica* has been observed within the upper Pliocene in other regions (Müller, 1979).

The lowermost part of the sequence (Cores 141-860B-62X to -70X) might belong to the lower Pliocene, owing to the presence of small *Reticulofenestra* cf. *pseudoumbilica*. But this is not a clear evidence for this age assignment.

Large parts of the Pliocene sequence are almost barren (Cores 141-860B-24X to -29X, and 141-860B-33X to -70X); only a few samples yielded rare nannofossils. However, the entire sequence is not older than 3.4 Ma as indicated by the presence of *Pseudoemiliana lacunosa* and the small *Gephyrocapsa* sp. down to at least Core 141-860B-61X.

### Planktonic Foraminifers

The biostratigraphy based on planktonic foraminifers from Site 860 is based on 157 samples from Hole 860B from the top down to 617.82 mbsf. Planktonic foraminifers occur in 85% of the investigated samples.

At Site 860 the FO of *T. truncatulinoides* was observed in Sample 141-860B-16X-3, 32–38 cm. Therefore, the sequence down to 120.12 mbsf belongs to the *T. truncatulinoides* Zone. The uppermost 6.50 m of this sequence contain rich assemblages with *T. truncatulinoides*, *Globorotaloides hexagonus*, *Neogloboquadrina dutertrei*, *Orbulina universa*, *Globigerina apertura*, and *G. bulloides*. The presence of these species and *G. inflata* indicates that tropical/subtropical taxa occurred together in transitional water masses and that warm temperatures had influenced the assemblages in the latest Pleistocene. The downhole occurrences of *N. pachyderma* indicate that more subpolar temperatures influenced the assemblages of planktonic foraminifers. The index fossil *T. truncatulinoides* occurs below 6.50 mbsf only sporadically in two samples (Samples 141-860-12X-2, 35–39 cm, and -16X-3, 32–38 cm). Therefore, the placement of the lower boundary of the *T. truncatulinoides* Zone is very uncertain. Nevertheless, in Hole 860B it corresponds well with the Zone NN19/NN18 boundary determined by calcareous nannofossils. The FO of *T. truncatulinoides* lies in a sequence common in *G. inflata*, which indicates transitional water masses.

The sediments below 24.70 mbsf contain translucent white and yellow-orange foraminifers together, indicating debris flows or other downslope sedimentation. No differences exist in the diversity of the white and the colored shells, which also show no difference in age. The calculated number of foraminifers in the samples of the *T. truncatulinoides* Zone fluctuated cyclically between zero and more than 25,000 specimens in 20 cm<sup>3</sup> of sediment. The uppermost cycle extends to 9.50 mbsf. The underlying cycle to 49 mbsf has its highest

abundances of foraminifers between 20 and 33 mbsf. The next cycle is to 91.90 mbsf, and until 116.80 mbsf most of the samples are barren in foraminifers, except that samples between 107.20 and 110.80 mbsf contain small amounts of stained, and therefore displaced, foraminifers. These fluctuations in the abundance of foraminifers may reflect different water paleotemperatures, but temperature assignments are difficult and tentative.

Below 120.12 mbsf the sequence can generally be described as an alternation of assemblages with and without *G. inflata*. This fact may indicate imbrication and repetitions. *G. inflata* was not observed between 155.50 through 192.02 mbsf, 203.70 and 242.50 mbsf, and 272.78 and 300.10 mbsf. Below 328.70 mbsf, *G. inflata* occurs sporadically at 376.90, 425.30, 451.20 and 532.49 through 535.04 mbsf, where the deepest observation of *G. inflata* was fixed. The FAD of *G. inflata* usually marks the boundary between early and late Pliocene, which is the boundary between the *Globorotalia crassaformis* Zone and the *G. inflata* Zone, respectively (Kennett and Srinivasan, 1983; Srinivasan and Kennett, 1981). The LO of *Globorotalia* cf. *spheri-comiozea* may indicate the lower Pliocene in Sample 141-860B-65X-CC and below (Table 4).

### Site 861

Site 861 is the easternmost site of the west-east transect of Leg 141 approximately parallel along to the 45°S latitude. This site is located the farthest landward and at the shallowest water depth. Sediments recovered in Holes 861A through Hole 861D from the seafloor to 496.30 mbsf represent a stratigraphic sequence ranging from Quaternary to Pliocene.

### Calcareous Nannoplankton

The sediments of Core 1H from Holes 861A and 861B are rich in nannofossils typical of the late Pleistocene *Emiliana huxleyi* Zone (NN21; Tables 5 and 6). The more diversified assemblages indicate the influence of warmer surface-water masses.

Nannofossils are few within the Pleistocene of Hole 861C, and they become rare throughout the Pliocene with exception of the lower part of Zone NN16 (Table 7). Their scarcity is the result of dissolution and dilution by a large amount of detrital material. The assemblages are of low species diversity, and the species are smaller than normal, owing probably to low surface-water temperatures. However, this may also partially reflect a relative enrichment related to selective dissolution that destroys the larger specimens first. The low diversity of the assemblages precludes conclusions about paleoclimatic changes.

Cores 141-861C-1H through -4H belong to nannoplankton Zone NN21 as shown by the presence of *Emiliana huxleyi*. Nannofossils are few to common, and they show signs of dissolution in Core 141-861C-1H. The assemblages consist of the following species: *Helicosphaera carteri*, *Coccolithus pelagicus*, *Gephyrocapsa ericsonii*, *G. oceanica*, *Cyclococcolithus leptoporus*, and *Emiliana huxleyi*. Other species such as *Syracosphaera pulchra*, *S. mediterranea*, *Rhabdosphaera claviger*, *Braarudosphaera bigelowi*, *Pontosphaera syracusana*, and *Scapholithus fossilis* were observed only sporadically.

The sediments belonging to nannoplankton Zone NN20 contain only rare nannofossils. The assemblages are very poor, containing only *Helicosphaera carteri*, *Cyclococcolithus leptoporus*, *Coccolithus pelagicus*, and *Gephyrocapsa ericsonii*. Core 141-861C-7H to Section 141-861C-15X-4 is almost barren of nannoplankton. Nannofossils become more common within the lowermost part of Core 141-861C-15X. The boundary between nannoplankton Zones NN19 and NN20 lies probably at the base of Core 141-861C-15X, where the last occurrence of *Pseudoemiliana lacunosa* was observed. Nannofossils are more frequent within the sediments attributed to Zone NN19.

The upper part of the Pliocene section is very poor in nannoplankton. Apart from core-catcher Sample 141-861C-24X, the sediments from Cores 141-861C-21X through -30X are barren. Nannofossils

Table 3. Distribution of calcareous nannoplankton and planktonic foraminifers in Hole 859B.

ODP 141-859B	Depth (mbsf)	Calcareous nannoplankton										Planktonic foraminifers																										
		Abundance	Preservation	<i>B. bigelowi</i>	<i>C. pelagicus</i>	<i>C. leptoporus</i>	<i>C. macintyreii</i>	<i>D. japonica</i>	<i>G. sp.</i>	<i>H. carteri</i>	<i>H. sellii</i>	<i>P. pacifica</i>	<i>P. lacunosus</i>	<i>R. cf. pseudoumbilica</i>	<i>S. pulchra</i>	Nannoplankton zone	Number of foraminifers in 20 cm <sup>3</sup> sediment	Planktonic foraminifers %	Preservation	<i>G. inflata</i>	<i>G. scitula</i>	<i>G. aff. bulloides</i>	<i>G. quinqueloba</i>	<i>G. crassaformis</i>	<i>G. glutinata</i>	<i>G. puncticulata</i>	<i>G. ruber</i>	<i>G. crassula</i>	<i>G. cf. inflata</i>	<i>N. pachyderma</i> sim.	<i>N. pachyderma</i> dex.	Remarks	Planktonic foraminifer zones					
1R-2, 66-71	54.16	R	G	F	R	R	F	F	R	R	R	R	R	R	NN16 - NN 18	250	95	C																				
1R-3, 63-68	55.63	R	G	F	R	R	F	F	R	R	R	R	R	R			50	95	C	R	R	R	R	R	R	R	R	R	R	R	R	R	R	R				
1R-CC	61.60	R	G	F	R	R	F	F	R	R	R	R	R	R			0																					
2R-1, 38-42	62.27	R	G	R	R	R	R	R	R	R	R	R	R	R			10	90	M									R										
2R-2, 70-72	63.48	T	M				R										25	90	M			R																
2R-4, 62-67	66.72	R	G	R	R	R	R	R	R	R	R	R	R	R			0																					
2R-CC	71.30	R	G	R	R	R	R	R	R	R	R	R	R	R			10		M						R													
3R-2, 66-71	73.46	R	G	R	R	R	R	R	R	R	R	R	R	R			200	95	M																			
3R-4, 77-82	75.29	T	M	R	R	R	R	R	R	R	R	R	R	R			200	95	M					R														
3R-6, 68-73	77.98	R	G	R	R	R	R										100	95	M			R	R															
3R-CC	80.90	R	G	F	R	R	F	F	R	R	R	R	R	R			50	95	M			R	R	R	R	R												
4R-1,120-122	141.20	B															1																					
4R-CC	148.50	R	G	R	R	R	R	R	R	R	R	R	R	R			250	80	M	R		R	R	R														
5R-CC	158.10	B															100	95	M			R	R	R														
9R-CC	196.70	B															10		P			R																
10R-2,68-70	198.88	B															0																					
10R-3,119-125	200.89	B															0																					
10R-5, 68-74	203.38	B															0																					
10R-CC	206.40	R	G	R	R	R	R	R	R	R	R	R	R	R			0																					
11R-1,58-62	206.98	B	G														0																					
11R-CC	216.00	T	M	R	R	R	R									0																						
12R-2, 70-72	217.50	B														0																						
12R-3, 91-93	219.21	B														0																						
12R-5, 66-67	221.52	B														0																						
12R-CC	225.60	B														0																						
13R-2, 57-62	226.47	B														0																						
13R-3,140-145	228.80	B														0																						
13R-5, 70-75	230.25	B														0																						
13R-CC	235.30	B														0																						
14R-2, 70-72	237.50	B														0																						
14R-3,130-134	239.60	B														0																						
14R-4, 70-74	240.50	B														0																						
14R-CC	245.00	B														0																						
15R-2, 27-33	246.77	B														0																						
15R-3, 27-33	248.27	B														0																						
15R-5, 28-33	250.78	B														0																						
15R-CC	254.60	B														0																						
16R-1, 70-72	255.30	B														0																						
16R-2, 20-22	256.30	B														1																						
16R-CC	264.30	R	G	R	R	R	R	R	R	R	R	R	R	R		2		C	P							R												
17R-1,122-124	265.52	B														1			C																			
17R-4,120-122	270.00	B														0																						
17R-6, 70-72	272.50	B														0																						
17R-CC	274.00	B														0																						
18R-1, 0-1	274.10	B														0																						
18R-4, 28-30	278.78	B														0																						
18R-CC	283.50	B														1		P																				
19R-2,120-122	286.20	B														0																						
19R-4, 55-57	288.55	B														0																						
19R-5, 4-7	289.54	R	M							T				T		0																						

become more common in Core 141-861C-34X, but decrease again in Core 141-861C-35X through -37X. The assemblages are of very low diversity containing *Cyclococcolithus macintyreii*, *Helicosphaera carteri*, *H. sellii*, *Pseudoemiliania lacunosa*, and *Gephyrocapsa* sp. The Pliocene/Pleistocene boundary is determined tentatively between Cores 141-861C-20X and -21X by the FO of *Gephyrocapsa oceanica*. Cores 141-21X through -25X were barren of calcareous nannoplankton. Small specimens of *Reticulofenestra* cf. *pseudoumbilica* are present from Core 141-861C-34X. This lowermost part of the sequence most probably belongs to the lower part of Zone NN16 (2.6–3.2 Ma.). This age agrees with the results obtained by foraminifers.

Hole 861D was washed down to 342.30 mbsf. The entire sequence is assigned to the upper Pliocene (Zones NN16–NN18; Table 8). Further subdivision is impossible because of the scarcity of nannofossils.

Core 141-861D-1R belongs to the upper Pliocene (Zone NN16–NN18) with the following assemblage: *Coccolithus pelagicus*, *Gephyrocapsa* sp., *Cyclococcolithus macintyreii*, *C. leptoporus*, *Pseudoemiliania lacunosa*, *Discolithina japonica*, *Helicosphaera carteri*, and *H.*

Table 3 (continued).

ODP 141-859B	Depth (mbsf)	Calcareous nannoplankton							Planktonic foraminifers																											
		Abundance	Preservation	<i>B. ligulata</i>	<i>C. pedregalis</i>	<i>C. leptopus</i>	<i>C. macintyri</i>	<i>D. japonica</i>	<i>G. sp.</i>	<i>H. carteri</i>	<i>H. sellii</i>	<i>P. pacifica</i>	<i>P. lacunosa</i>	<i>R. cf. pseudumbilica</i>	<i>S. pulchra</i>	Nannoplankton zone	Number of foraminifers in 20 cm <sup>3</sup> sediment	Planktonic foraminifers %	Preservation	<i>G. inflata</i>	<i>G. scitula</i>	<i>G. aff. bulloides</i>	<i>G. quinqueloba</i>	<i>G. crassaformis</i>	<i>G. glutinata</i>	<i>G. puncticulata</i>	<i>G. ruber</i>	<i>G. crassula</i>	<i>G. cf. inflata</i>	<i>N. pachyderma sin.</i>	<i>N. pachyderma dex.</i>	Remarks	Planktonic foraminifer zones			
19R-CC	293.20	B	B													0																				
20R-3, 63-38	296.83	R	G													0																				
20R-4, 7-12	297.77	R	G		R											0																				
20R-CC	302.80	B														10		P																		
21R-2, 79-84	305.09	B														25	1	P			R															
21R-4, 59-64	307.89	T	M		R			R								1		G																c		
21R-CC	312.20	B														0																				
22R-1, 56-60	312.76	B														0																				
22R-CC	321.90	B														10																				
23R-CC	331.50	B														50																				
25R-2, 27-30	342.97	B														100		P																d		
25R-CC	350.80	B														50	20	P																d		
26R-2, 0-2	352.30	B														0	80	P	R		R													d		
26R-CC	360.40	T	M		R											15	85	P			R														d	
27R-2, 52-57	362.42	R	G		R			R								50																				
27R-3, 44-48	363.84	B														1000		P																		
27R-4, 27-31	365.17	T	M					R								250	90	P							R											
27R-CC	370.10	R	M		F	R	T	F	R		R	F				0	80	P		R	R														d	
28R-3, 128-130	374.38	B														2	90	P																		
28R-CC	379.50	B														50																				
29R-3, 87-93	383.38	B														100		P																		
29R-5, 18-22	385.69	R	G		F			R	R		R					0		P																		
29R-CC	389.20	B														150	95	P	R	R																
30R-2, 70-75	391.40	B														25																				
30R-3, 130-135	393.50	T	M		R											100	80	P				R		R												
30R-5, 70-75	395.90	B														3	90	P																		
30R-CC	398.90	B														3	90	P				R														
32R-1, 70-75	409.30	B														0		P																		
32R-CC	418.30	T	M		R											50		P																		
33R-2, 21-26	420.01	T	M		R											2																				
33R-CC	427.90	T	M					R								0	70	P																		
34R-2, 68-70	430.08	B														3		P																		
34R-3, 134-136	432.24	B														1																				
34R-CC	437.60	B														0		P																		
35R-1, 90-92	438.50	B														0																				
35R-3, 90-92	441.50	B														0																				
35R-CC	447.10	B														0																				
36R-2, 69-74	449.29	B														4																				
36R-2, 74-77	449.34	B														0																				
36R-CC	456.70	T	M		R			R								0		P						R												
37R-1, 55-59	457.25	B														3																				
37R-CC	466.40	R	M		R			R	R		R	R				0																				
38R-CC	476.10	B														0		P																		

Note: c = different colors, d = deformed specimen.

**Planktonic Foraminifers**

The biostratigraphic results obtained from planktonic foraminifers are based on investigation of nine samples from Hole 861A and six samples from Hole 861B, which penetrates only the upper 10 m of the sequence, as well as on studies of 131 samples from Hole 861C (353.10 m total depth) and 35 samples from Hole 861D (344.48 to 496.30 mbsf).

At Site 861 two recognized planktonic foraminifer events are the FO of *T. truncatulinoides*, which defines the lower boundary of the *T. truncatulinoides* Zone in Sample 141-861C-16X-4, 30-35 cm, at 133.70 mbsf, and the FO of *G. inflata*, which marks the lower boundary of the *G. inflata* Zone in Sample 141-861D-8R-1, 46-51 cm, at 410.26 mbsf. On the basis of planktonic foraminifers three zones are distinguishable in the Quaternary and Pliocene series: the *T. truncatulinoides* Zone, the *G. inflata* Zone, and the *G. crassaformis* Zone.

Paleotemperatures are apparently reflected by the foraminiferal assemblages. The uppermost 3 m of the *T. truncatulinoides* Zone contain well-preserved and rich assemblages with *Globorotaloides hexagonus*, *Beella digitata*, *Neogloboquadrina dutertrei*, *Globigerina apertura*, *G. bulloides*, *Orbulina universa*, and *T. truncatulinoides* accompanied by *G. inflata*. This indicates the occurrence of

tropical/subtropical taxa during the late Pleistocene in transitional water masses. Down to 69.50 mbsf most of the samples contain *T. truncatulinoides* and *N. pachyderma*, which indicate lower temperatures. The FO of *T. truncatulinoides* is documented only by one whole and a broken specimen, which shows the uncertainty of the zonal boundary. Cyclic fluctuations in abundances of foraminifers as observed in the *T. truncatulinoides* Zone of Hole 860B were also recognized within the sequences from Site 861. Low number of foraminifers (Tables 5, 6, and 7) coincides roughly with even isotope stages, corresponding to cool intervals (e.g. at 7.15 mbsf = oxygen isotope event 2.2 according Schönfeld and Spiegler (this volume), between 12 and 14 mbsf = event 4.2, at 27.59 mbsf = event 6.2, and around 48 mbsf = oxygen isotope event 8). Below 42.80 mbsf most of the samples contain yellow- to red-colored foraminifers together with white specimens, which documents downslope redeposition also described from Site 860.

**Site 862**

Site 862 was drilled to determine the basement lithology of the Taitao Ridge which was reached in Hole 862C.

















Table 9. Distribution of calcareous nannoplankton and planktonic foraminifers in Holes 862A, 862B, and 862C.

ODP 141-862A	Samples (Core, section, interval in cm)	Depth (mbsf)	Calcareous nannoplankton					Planktonic foraminifers																							
			Abundance	Preservation	<i>C. pelagicus</i>	<i>C. leptopus</i>	<i>C. macintyrei</i>	<i>G. sp.</i>	<i>H. carteri</i>	<i>P. lacunosa</i>	Nannoplankton zone	Numbers of Forams in 20 cm <sup>3</sup> sediment	Plankt. foram. %	Preservation	<i>G. hexagonus</i>	<i>B. digitata</i>	<i>N. duterrei</i>	<i>T. truncatulinoides</i>	<i>G. apertura</i>	<i>G. bulloides</i>	<i>G. cf. bulloides</i>	<i>G. inflata</i>	<i>G. scitula</i>	<i>O. universoni</i>	<i>N. pachyderma</i> dex.	<i>N. pachyderma</i> sin.	<i>G. crassaformis</i>	Remarks	Planktonic foraminifer zones		
	1H-1, 14-18	0.14	B							1000	95	C	R	R	R	F	C	R	R	R	R	R	R	R							
	1H-1, 18-21	0.18	B							1000	95	C	R	R	R	R	F	A	R	R	R	R	R	R	R						
	1H-1, 62-64	0.62	B							100	80	M	P												R						
	1H-1, 88-91	0.88	B							8	-	P													R						
	1H-1, 101-104	1.01	B							50	80	M				R		R	R	R					R						
	1H-2, 73-76	2.23	T	M	T	T				250	90	C							R						F						
	1H-2, 131-134	2.81	B							500	90	M							R		R				F	C	R		c		
	1H-4, 31-34	4.81	B							25	10	M													F						
	1H-CC	5.40	B							5	-	M							R		R				R						
	2H-1, 42-43	5.82	B							100	60	P													C				c		
	2H-1, 95-97	6.35	B							3	-	C													R						
	2H-4, 103-105	10.93	B							50	95	P													R	F			c		
	2H-5, 80-82	12.20	B							0	-																				
	2H-6, 111-113	14.01	T	M	T	T				15	95	P													R						
	2H-CC	14.90	B							0	-																				
	3H-1, 38-41	15.28	B							1	-														R						
	3H-2, 60-63	17.00	B							5	-														R						
	3H-4, 48-51	19.88	B							1	0																				
	3H-4, 85-88	20.25	B							0	-																				
	3H-4, 106-109	20.46	B							0	-																				
	3H-CC	20.60	F	M	F	R	T	R	R	100	90	M							R						R	R			c		
	4H-CC	21.10	B						NN16-18	0	-																				
<b>ODP 141-862B</b>																															
	2X-1,96-98	18.46	B							0	-															R					
	2X-CC	27.10	B							15	90	M																			
<b>ODP 141-862C</b>																															
	1W-1.10-12	-	B							25	10	G							R							R					
	1W-1.32-39	-	B							50	5	C																			
	6R-1, 37-39	Piece 6	B							0	-																				

Note: c = different colors.

generally poorly diversified assemblages can also be explained by cold surface-water. However, the fossil assemblages observed within the uppermost meters of the sequences indicate the influence of warmer surface-water off southern Chile in further times. Planktonic foraminifers from Leg 141 sites are documented on Plate 1.

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\* Abbreviations for names of organizations and publications in ODP reference lists follow the style given in *Chemical Abstracts Service Source Index* (published by American Chemical Society).

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Table 11. Distribution of calcareous nannoplankton and planktonic foraminifers in Hole 863B.

ODP 141-863B	Samples (core, section, interval in cm)	Depth (mbsf)	Calcareous nannoplankton						Nannoplankton zone	Planktonic foraminifers																	
			Abundance	Preservation	<i>C. pelagicus</i>	<i>C. leptopus</i>	<i>C. caribbeana</i>	<i>C. ericsonii</i>		<i>G. azanica</i>	<i>H. carteri</i>	<i>S. mediterranea</i>	Numbers of foraminifer in 20 cm <sup>3</sup> sediment	Plankt. foram. %	Preservation	<i>G. inflata</i>	<i>G. crassaformis</i>	<i>G. crassula</i>	<i>N. pachydermus</i> sin.	<i>N. pachydermus</i>	<i>G. cf. bulloides</i>	<i>G. quinqueloba</i>	<i>O. irripes</i>	<i>G. glauca</i>	Remarks	Planktonic foraminifer zones	
	4X-2, 70-73	328.10	F	C	T						50	85	M	F	R	R											
	4X-3, 52-55	329.42	R	C			R				50	60	M	R				R	R	R							
	4X-5, 48-51	332.38	F	C		R	R	F			10	-	P			R	R	R									
	4X-CC	335.60	F	C							25	-	P	R	R	R											
	5X-CC	345.20	F	C		R		F	R		20	-	P			R				R							
	6X-CC	354.90	F	C	R	R	R	F			300	90	M	R	R	F		F	R				s				
	7X-1, 71-73	355.61	F	C	F	R		C	T		500	95	M	F		R	F		F		R		s				
	7X-CC	357.40	B								10	-	P			R											
	8X-CC	361.40	C	G		R		C		R	0																
	9X-CC	371.00	C	G		R		C		R	0																
	10R-1, 45-48	371.45	R	G		R		R			25	-	M	R	R												
	10R-CC	376.60	B								0																
	11R-1, 47-49	377.07	F	C	R	R		F			2500	95	G	F		A		R	R	R		c					
	11R-CC	386.30	B								0																
	12R-CC	395.80	B								2	-	P	R									s				
	13R-1, 46-48	396.26	R	G		R		R			3	-											s				
	13R-CC	405.50	B								0																
	14R-CC	415.20	B								4		P	R									s				
	15R-CC	424.70	F	G		R		C			10			R									d				
	16R-CC	434.40	F	G	R	R		F	R		20		M	F	R	R							s				
	17R-2, 46-49	436.36	F	G	R	R		F			13000	90	P	C	F	R		C				s,d					
	17R-4, 14-16	439.04	B								1500	90	P	R	R	R		C				s					
	17R-6, 3-7	441.93	C	G	F	R	R	C			6000	80	P	F	R	F		C				s,d					
	17R-CC	444.00	C	G		R		F	C		0																
	18R-1, 100-104	445.00	C	G	F	R		C	R		10000	90	P	C	F	R	R	A				s,d					
	18R-3, 82-86	447.82	C	G	F	R		C			2500	90	P	F	R	R	R	A				s,d					
	18R-CC	453.70	F	C	R	R	F	F			80	100	P	R	R	R						s					
	19R-3, 80-83	457.50	C	C	R	R	C	F	R		250	90	M	F		R		R				s,d					
	19R-CC	463.40	C	G	R	R	F	C	R		200	90	M	F	R	R						s,d					
	20R-2, 83-85	465.59	F	G	R	F	F	F			300	90	P	F				R	R			s					
	20R-CC	473.00	R	M			F	R			10	90	P	R	R							s					
	21R-CC	482.70	F	G	R	R	F	F			25	20	P	R								s					
	22R-CC	492.40	C	G	R	R	F	C	R		50	80	P	F	R	R						s					
	23R-4, 79-83	497.69	C	G	R	R		F	R	R	50	80	P	R	R	R		R	R			s					
	23R-CC	502.00	B								3	-	P	R	R			R				s					
	24R-CC	511.70	B								1	-	P	R								s					
	25R-CC	521.40	F	G	R	R	F	F	R		20	50	P	R	R		R	R				s					
	26R-1, 7-9	521.47	C	G	C		F	R	F		10	-	P	R	R		R	R				s					
	26R-2, 50-53	522.45	C	G	C	R	C		F		200	80	P	R		R		C				s,d					
	26R-CC	530.60	C	G	F	R	F	C			25	50	P	R			R	R				s,d					
	27R-01, 71-74	531.31	B								25	80	P	R	R	R						s					
	27R-CC	540.20	B								30	50	P	R	R		F					s					
	28R-CC	549.70	C	G	R	R	F	C	R	R	1	-	C	R								s					
	29R-CC	559.40	C	G	R	R	F	C	R	R	50	90	G	F	R			R				s					
	30R-CC	569.10	C	G	F	R	F	C	R		1000	90	C	F	R	R		C	R	R		s					
	31R-CC	578.80	C	G	R	R	F	C	R	R	25	90	P	R	R							s					
	32R-CC	588.40	B	G							10	0	P									s					
	33R-CC	598.10	F	G		R	R	F	F	R	10	-	P	R		R						s					
	34R-CC	607.70	B	G							0			R													
	35R-6, 45-47	615.65	F		R	R	F	F	R		5	-	P					R				s					
	35R-CC	617.40	B	G							0																
	36R-CC	627.00	F	G	R	R	F	F	R		2	-	P	R		R						s					
	37R-CC	636.60	R	G				R			2	-	P	R													
	38R-CC	646.30	C	G	R	R	F	C	R		0																
	39R-CC, 4-10	650.34	B								100	80	P	R				R				s,d					
	39R-CC	655.90	F	G	R	R		F			50	-	P	R	R	R	R	R				s					
	40R-1, 45-48	656.35	R	G		R		R			200	80	P	F	R	F	F	F				s,d					
	40R-3, 84-88	659.44	R	M				R			200	100	M	F	R	F		F				s,d					
	40R-CC	665.50	F	G				F			100	80	P	R	R	R						s					
	41R-2, 92-95	667.92	R	G		R		F	R		150	80	P	F		R						s					
	41R-4, 32-36	670.32	F	G	R	R		F			250	90	P	C				R				s,d					
	41R-CC	675.20	C	G	R	R	F	C			100	95	P	F	R	R						s					
	42R-CC	684.80	C	G	R	R	C	C	R	R	100	95	P	F	R							s					
	43R-3, 47-52	687.77	F	G		R		F			50	90	P	R		R						s,d					
	43R-CC	694.50	C	G	R	R	C	C	R		150	95	M	C				R				s					
	44R-CC	704.20	C	G	R	R	C	C			5	-	P	R				R				s					
	45R-CC	706.80	C	G	R	F	C	C			4	-	P	R								s					
	46R-CC	713.80	C	G	R	R	C	C	R	R	0																
	47R-CC	723.50	C	G	R	R	C	C			0																
	48R-CC, 20-22	725.69	C	G	F	F	C	C	R		100	90	P	R		R		R	R			s					
	48R-CC	733.20	C	G	R	R	C	C			50	90	P	F				R				s					
	49R-1, 19-22	733.39	C	G	F	F	C		R		100	90	P	F	R	F		R				s					
	49R-CC	739.40	C	G	R	R	C	A	R		100	80	M	F	R			R				s					

Note: c = different colors, d = deformed, s = silicified.

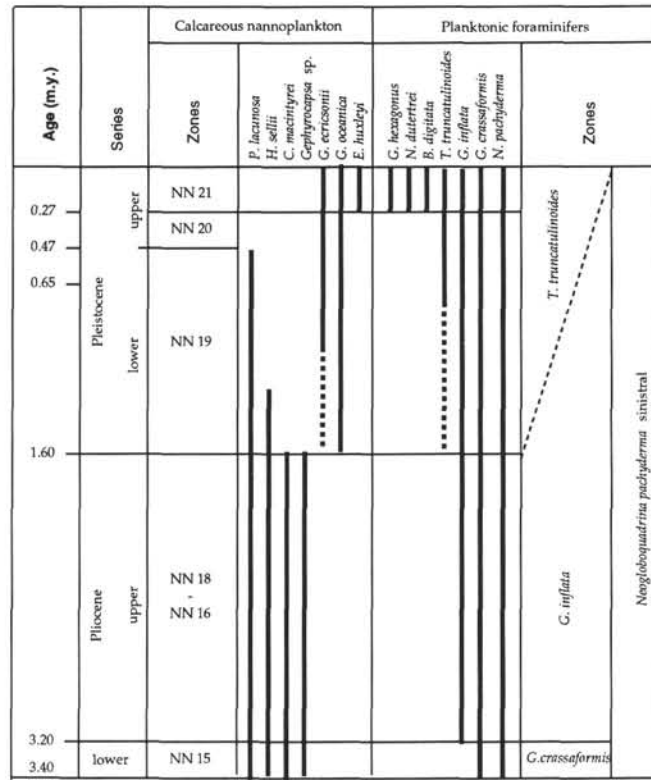


Figure 5. Range chart of Leg 141 index species.

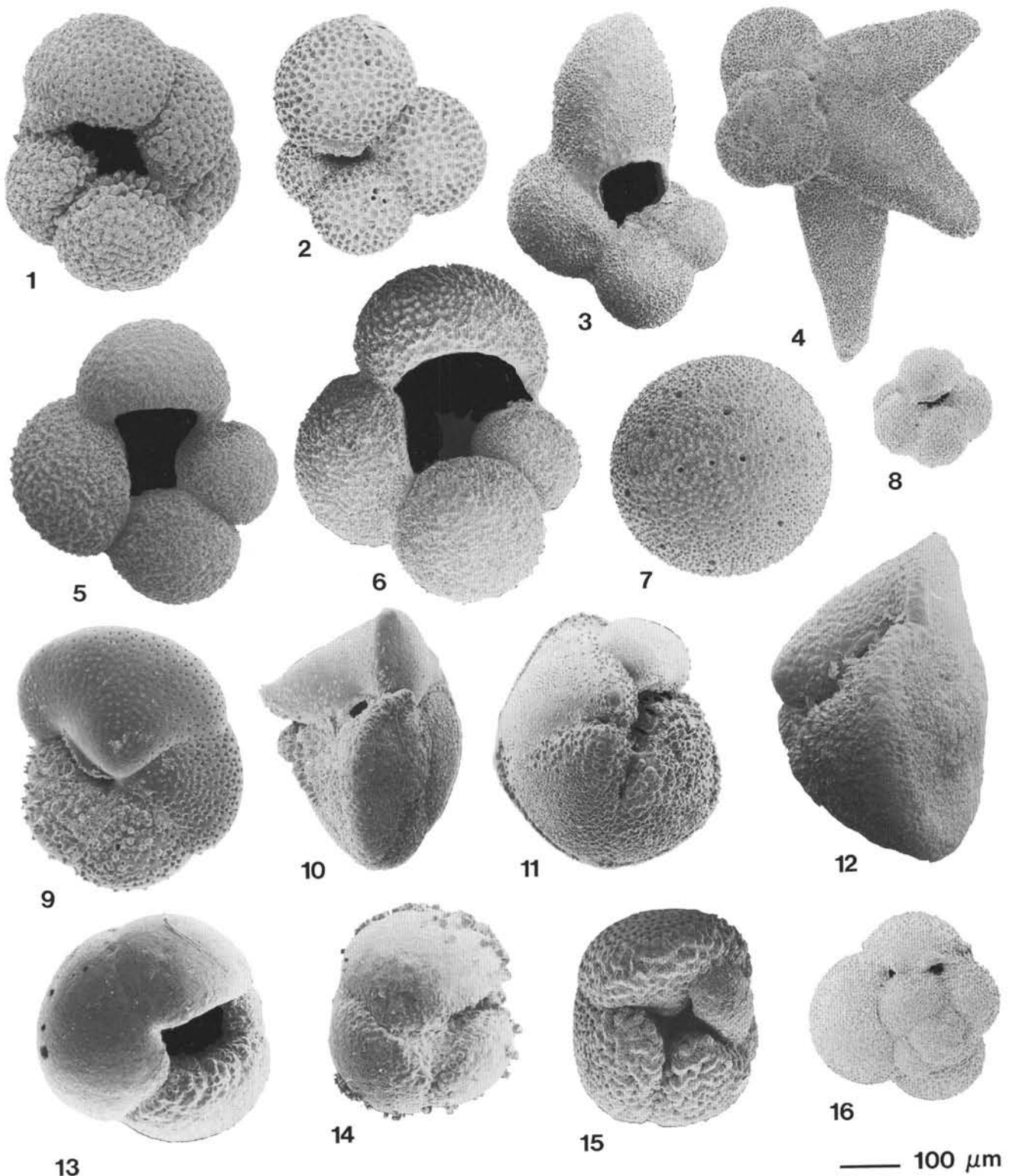


Plate 1. Pleistocene planktonic foraminifers from Leg 141 sites. **1.** *Neogloboquadrina dutertrei* (Orbigny, 1839); Sample 141-861C-1H-1, 45–49 cm. **2.** *Globorotaloides hexagonus* (Natland, 1938); Sample 141-862A-1H-1, 14–18 cm. **3, 4.** *Beella digitata* (Brady, 1879). (3) Sample 141-861C-1H-CC, (4) Sample 141-862A-1H-1, 18–21 cm. **5.** *Globigerina bulloides* Orbigny, 1826; Sample 141-861C-1H-1, 45–49 cm. **6.** *Globigerina apertura* Cushman, 1918; Sample 141-861C-1H-CC. **7.** *Orbulina universa* Orbigny, 1839; Sample 141-861C-15X-2, 6–10 cm. **8.** *Turborotalia quinqueloba* (Natland, 1938); Sample 141-861C-3H-CC. **9.** *Globorotalia scitula* Brady, 1882; Sample 141-8612A-1H-1, 14–18 cm. **10–12.** *Truncorotalia truncatulinoides* Orbigny, 1839; (10) Sample 141-862A-1H-1, 18–21 cm, (11–12) Sample 141-861C-1H-CC. **13–14.** *Globoconella inflata* (Orbigny, 1839), (13) Sample 141-861C-4H-4, 109–114 cm, (14) Specimen covered with carbonate appositions, Sample 141-862A-1H-1, 18–21 cm. **15.** *Globorotalia crassaformis* (Galloway and Wiesner, 1927); Sample 141-861C-16X-4, 30–35 cm. **16.** *Globigerinita glutinata* (Egger, 1893) with two supplementary apertures on the spiral sutures of the final chamber as shown at tropical *Globigerinita parkerae* (Bermudez, 1961); Sample 141-861C-3H-2, 70–75 cm.