21. FORMATION MICROSCANNER LOGGING RESPONSES TO LITHOLOGY IN GUYOT CARBONATE PLATFORMS AND THEIR IMPLICATIONS: SITES 865 AND 866¹

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ABSTRACT

Formation MicroScanner (FMS) images from Sites 865 and 866, Allison and Resolution guyots, respectively, were integrated with conventional logging data and core descriptions to provide detailed stratigraphic columns for those sites. First, large-scale (in hundreds of meters) lithologic units were determined from the conventional logs and compared to core descriptions; these units typically contain several major carbonate facies. Second, a "type" facies log response was established for each carbonate facies by correlating the best FMS images with the conventional log response and core information for the same depth. Finally, the remaining portions of the FMS images were interpreted using the "type" facies log responses as standards of comparison. It was immediately apparent that core recovery at both sites was highly preferential (e.g., within the alternating packstone-wackestone intervals, only small pebbles of well-cemented wackestone were recovered). This study indicated that packstone was the dominant lithology (51%) at Hole 865A and that grainstone (25%) and packstone (20%) were the dominant lithologies at Hole 866A (excluding dolomite).

The sedimentary record, as determined from core-log integration, not only confirmed shipboard conclusions regarding the gross vertical trajectories of Allison and Resolution guyots, but provided details of facies changes that resulted from small-scale fluctuations in sea level. Deposition of the Albian section of Resolution Guyot overlaps in time with deposition of sediments at Allison Guyot. This study supports Strasser's conclusion (this volume) that a hiatus probably exists near the Albian/Aptian boundary at Resolution Guyot. Correlations of bed thicknesses and logging signatures between the two holes indicates that much of the Albian section may have been removed from Allison Guyot as it emerged above sea level.

INTRODUCTION

The central and northwestern Pacific Ocean seafloor is festooned with platforms and chains or groupings of seamounts, many of which are Cretaceous guyots, with their summits at depths of about 1500 m (Menard, 1964; Matthews et al., 1974; Winterer and Metzler, 1984; McNutt et al., 1990). Many of the volcanic edifices of the guyots are capped by shallow-water carbonate sediments. The sediments, which are almost always deposited at or near sea level, record the vertical trajectory of these guyots relative to sea level, caused by subsidence or uplift as well as fluctuations in eustatic sea level. The main scientific theme of the Leg 143 drilling program was the origin and evolution of Cretaceous guyots and their relationship to sea level. Thick deposits of carbonate platform sediments were sampled at Sites 865 and 866, located in the lagoons of two drowned carbonate platforms in the Mid-Pacific Mountains, Allison and Resolution, respectively (Fig. 1).

The formation of thick carbonate deposits requires a dynamic equilibrium between carbonate deposition and subsidence. The added dimension of time leads to the development of stratigraphic sequences that reflect changes in the depositional style. Variables affecting the style of deposition include changes in sea level, sedimentation rates, type of facies at the platform margin, and tectonic activity. A continuous vertical record of the stratigraphic sequence is necessary to determine the history of the relative influences of these variables on deposition at a particular site. For that reason, the stratigraphic record is usually the primary objective of scientific drilling, but is the most difficult to obtain when using rotary (RCB) drilling in reefal carbonates. Indeed, mean core recovery at Sites 865 and 866 was less than 16% and only 1% to 2% in the shallow-water limestones. Thus, logs are particularly important at those sites to address the primary objectives of the drilling (Shipboard Scientific Party, 1993). Downhole logs provide the essential addition of a continuous time-series measurement of the physical properties of the borehole wall. The purpose of this study is to integrate the conventional logging data, FMS images, and core information to provide detailed stratigraphic information that can constrain the vertical trajectories of the guyots, allowing one to correlate directly with other guyots drilled during Legs 143 and 144.

DESCRIPTION AND INTERPRETATION OF FMS LOGS

Logging Program at Sites **865 and 866**

Ocean Drilling Program (ODP) Leg 143 obtained an extensive suite of in-situ borehole logging measurements (Sager, Winterer, Firth, et al., 1993). Hole 865 A (Fig. 1), which penetrated 870.9 m of calcareous oozes, wackestones and packstones, and basalts at Allison Guyot, was the only hole logged at Site 865; good logs were obtained within the open borehole between 100.5 and 867.0 mbsf. Five tool strings were run, including the sonic-porosity-density-gamma, the resistivitygamma, the geochemical tool string, the Japanese downhole magnetometer, and the Formation MicroScanner (FMS). Descriptions of the tools and their data characteristics are found in Shipboard Scientific Party (1993). The same five tool strings were run in Hole 866A at Resolution Guyot (Fig. 1) to cover the interval 74.5 to 1679.4 mbsf for the first run only. Subsequent runs concluded at shallower depths because of cave-ins above 1679.4 mbsf, and the FMS was not run above 253.0 mbsf because of the extremely large hole size (>17.2 in. or 43.7 cm).

In both holes, cycle skipping and other noise resulting from inadequate centralization of the tool compromised the quality of the sonic logs; hence, these were not used extensively for this analysis. Because of its mechanical problems, no valid data were recorded by the geochemical tool at either hole. Large hole diameters over 20- to 35-m

¹ Winterer, E.L., Sager, W.W., Firth, J.V., and Sinton, J.M. (Eds.), 1995. *Proc. ODP, Sci. Results,* 143: College Station, TX (Ocean Drilling Program).

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Figure 1. Location of Leg 143 drill sites and principal seamount chains, western central Pacific Ocean basin. Stippled areas are shallower than 4 km. Solid line shows track of the *JOIDES Resolution.*

intervals account for questionable porosity/density data for large portions of both holes.

FMS Resistivity Measurements

The resistivity log is a combined measure of the resistivities of both the rock-forming minerals and the fluids contained in the pore spaces. The constitutive minerals of common sedimentary rocks are highly resistive; hence, water content and salinity are the two most important factors that control measured electrical resistivity. Other factors that can influence resistivity include the concentration of hydrous and metallic minerals, vesicularity, the geometry of interconnected pore spaces, and, to a lesser extent, temperature. Research has shown that, to a first-order approximation, resistivity is proportional to the inverse square root of the porosity (Archie, 1942).

In pure limestones, electrical current is conducted almost entirely by the brine contained in the pore space. High resistivities correspond to low porosities; low resistivities correspond to high porosities. Wet clays conduct electricity, and the magnitude of this conduction effect in clayey limestones will depend on the percentage of clay content and clay mineral composition (Archie, 1952).

The FMS produces high-resolution resistivity measurements of the borehole wall (Ekstrom et al., 1987) that are ideal for detailed sedimentological studies (e.g., Bourke et al., 1989; Adams et al., 1990; Luthi, 1990; Harker et al., 1990). The tool consists of 16 electrodes on each of four orthogonal pads that press against the borehole

wall. Electrodes are spaced about 2.5 mm apart in two diagonally offset rows. The focussed current that flows from the electrodes is recorded as a series of measurements of the variations in resistivity with respect to depth. Shore-based processing converts these measurements into spatially oriented images of the borehole wall. With a sampling interval of 2.5 mm, the vertical resolution is 2.5 mm (Serra, 1989). Coverage of the borehole wall for each pass of the tool varies inversely with hole diameter—about 40% of an 8.5-in. hole, less for greater diameters. Perhaps the most important limitation of the tool is the restriction to hole diameters of less than 38.1 cm (15 in.). Some information may be gleaned from the image if one or two pads make good contact with the borehole wall, as is the case with highly elliptical holes, but often no useful information is contained in FMS images from washed out sections.

The resistivity values of the surfaces sampled by the pads are converted to gray scale and plotted as a set of vertical strips with an indication of pad orientation. This oriented image (Figs. 2-7) corresponds to an unrolled cylinder with a width that is equivalent to the diameter of the borehole. Horizontal and vertical scales are the same. In general, low resistivities (high conductivities) are shown as dark tones, whereas high resistivities (low conductivities) are shown as light tones. The image tone is a qualitative representation of electrical resistivity because, when processing the data, normalization procedures are applied to optimize the contrast on the images. Images are usually interpreted with the aid of conventional resistivity logs to quantify the variation of electrical resistivity within the formation.

Figure 2. FMS image from 692.8 to 694.2 mbsf, Hole 866A, interpreted as grainstone texture showing intervals of greater (light) and lesser (dark) cemen tation. Dark spots are molds or vugs.

Methodology for Integration of Conventional Logging Data, FMS Images, and Core Data

Integration of the FMS, conventional logs, and cores proceeded in the following way: First, conventional resistivity, velocity, porosity, density, gamma, and caliper logs were plotted together for 100-m intervals. This multiple-trace presentation facilitated determination of large-scale trends and patterns. The major sedimentary units were identified on the basis of baseline shifts or similar variability, mainly in the resistivity and natural gamma-ray logs (see Shipboard Scien tific Party, 1993). These major sedimentary units then were identified in the FMS images.

To do this, the FMS logs were depth-matched to the conventional logs by correlating the gamma-ray logs between runs. A linear depth shift was applied to all logs to correct for differences in the stretch of the logging cable, using the gamma log from the first logging run as reference. These depths should be accurate to within ±2 m of the drillers' depths. All logging depths (recorded as feet below rig floor) were then corrected to meters below seafloor by subtracting the depth from the rig floor to the seafloor. Short, repeat sections of the FMS for both holes were compared to differentiate between geologic real ity and processing artifact.

Second, the identification of a "type" log facies for each recovered core facies was accomplished by intercomparison of FMS, conven tional logs, and core descriptions. The FMS and conventional logs were displayed and inspected simultaneously to identify conventional log responses that corresponded to characteristic patterns observed on the FMS image. Long (>IO m) FMS and conventional logging sec tions with reasonably uniform characteristics were then compared to descriptions of core material taken from the same depth range. Spe

Figure 3. This FMS image from 271.9 to 273.4 mbsf, Hole 866A, interpreted as a packstone (top)-wackestone (bottom) texture, presents a mottled appear ance with irregular white spots on an irregular dark gray background. Dark spots are molds or vugs.

cial care was exercised because core recovery was poor and obviously preferential; typically, within the alternating packstone-wackestone intervals, only small pebbles of the well-cemented wackestones were recovered. However, we reasoned that if the logging responses were similar over a given depth interval and only one lithology was recov ered over that interval, then we could be certain that our interlog cor relation as a "type" facies was accurate. The classification of shallow water carbonate textures used here is the same as that used in Sager, Winterer, Firth, et al. (1993), namely, that of Dunham (1962). Fortu nately, the thick sections of wackestone-packstones and grainstones provided multiple "type" facies that could be checked against each other. The remaining portions of the FMS logs were interpreted by comparison to the "type" facies logging responses.

The greatest source of error in the identification of facies in the FMS images was hole size. Hole diameters for more than two-thirds of Hole 865A and one-third of Hole 866A were larger than the maximum opening diameter (38.1 cm) of the FMS calipers. The FMS data were of poor quality for the depth intervals 102-145, 198-212, 248-296, 305-322, 373-440, 500-527, 565-638, 651-702, and 746-791 in Hole 865A, and 200-262, 375-390, 455-661, 802-825, 976 993,1037-1054,1078-1106,1124-1152,1278-1292, and 1317-1340 in Hole 866A. Poor pad contact in these portions of the hole degraded the quality of the images. On some images two, sometimes only one, pads made good contact with the borehole wall; nonetheless, the images could still be interpreted based on comparisons to the type log facies. The conventional logs did not require contact with the bore hole wall; corrections for the effects of variations in hole size were applied in post-cruise data processing.

Figure 4. FMS image from 663.1 to 664.6 mbsf, Hole 866A. A mudstone presents a relatively homogeneous white to light gray image, depending on the porosity (the lower the porosity, the whiter the image). The black band above 664.0 mbsf is a clay seam.

FMS IMAGE INTERPRETATIONS

Keeping in mind that the image tone is not directly proportional to the resistivity value, the dark tones of the FMS image usually corre spond to materials that are conductive because of high intergranular porosity, fine-grained textures, or the presence of conductive miner als (e.g., clay and pyrite). The light tones correspond to materials that are resistive, either because of low intergranular porosity (e.g., mud stones) or coarse-grained texture (Serra, 1989). Identification of lithology was sometimes clear and straightforward, sometimes intui tive. For example, thick $(>1$ cm) organic- and clay-rich facies were easily identifiable in the FMS image as sharply defined black bands and in conventional logs, as a sharp, low-resistivity peak correspond ing to a high-gamma peak. Thin layers or disseminated clay and organic-rich matter (as reported in the core descriptions) produced subtle logging responses that would have been difficult to interpret without calibration by the core material. Some minor lithologies, notably boundstone, were difficult to identify accurately in the FMS images because these produced no characteristic logging response aside from low density values, which are common to several carbon ate textures. Combining the core descriptions with boundstone tex ture described by Serra (1989), we concluded that the boundstone texture was characterized by irregular, dark features (molds) in a gray background with white, somewhat vertical, interconnecting features that would result from the continuous, cemented framework, and with white spots or streaks (shell fragments).

Figure 5. FMS image from 283.8 to 285.4 mbsf, Hole 866A. Images of mudstones with extreme moldic porosity closely resemble those of wackestone; the two textures can be distinguished only by their relative densities.

Grainstone texture was characterized by obvious layering (Fig. 2); individual layers having a constant gray tone corresponded to wellsorted intervals. Packstones and wackestones presented a mottled appearance as a result of their highly variable grain sizes and porosities, with irregular white spots on an irregular gray background, sometimes with dark spots (molds or vugs). Wackestones (Fig. 3, bottom portion) contain more lime mud than do packstones (Fig. 3, top portion) and therefore are more resistive. Mudstones produced a relatively homogeneous white to light gray tone depending on the porosity (Fig. 4). Images of mudstone having extensive moldic porosity (Fig. 5) closely resembled packstone-wackestone, but could be distinguished on the basis of relatively higher density in the conventional logs. Rudstone texture is characterized by irregular white spots (packstone-wackestone) isolated in a dark background and in the conventional logs by low density, resistivity, and sonic velocity. Dolomite often displays a texture similar to rudstone: leached and moldic porosities combine to present a dark background (Fig. 6), but much higher resistivity values in the conventional log. Partial dolomitization is indicated in the FMS image by dark areas that correspond to zones of higher porosity that may cross bedding planes (e.g., in the grainstones, Fig. 7).

Figure 8 provides a key to patterns that indicate carbonate textures in Figures 9 and 10. Figures 9 and 10 display the conventional logs most important for interpreting the FMS images, including resistivity, gamma-ray, density, and caliper, the core number and amount recovered, together with interpretation of the lithology. The pattern indicates

Figure 6. FMS image from 1535.8 to 1537.6 mbsf, Hole 866A. Dolomites typically display a rudstone texture (i.e., irregular white spots [dense, redeposited carbonate]) in an irregular, dark gray to black background (leached and moldic porosities).

the dominant lithology for an interval; in some cases, a split pattern is shown to indicate almost equal interbeds of another lithology. Split patterns also are used to indicate partial dolomitization or a dolomitized section having relict texture. Because of the limited range of black- and-white patterns and shadings available, more detailed descriptions of the facies are listed in Tables 1 (Hole 865A) and 2 (Hole 866A). Occasionally, the core material was matched to the FMS images on the basis of such distinctive characteristics as size and amount of moldic porosity, presence of stylolites or clay seams, among others. Core numbers are placed in the comment section of the bed or beds from which they may have originated. In many cases, especially in the thin packstone-wackestone interbeds, it was not possible to identify the place of origin of the core material with confidence.

EFFECTS OF LITHOLOGY ON CORING

Overall recovery in both holes was poor, averaging only about 2% in the shallow-water limestones, despite repeated efforts to improve recovery by changing the style of coring. Instead of smoothly cut core, hard, shelly wackestone was recovered as small chunks with

Figure 7. FMS image from 1458.8 to 1460.4 mbsf, Hole 866A. Dolomitization occurs preferentially within the most porous facies, the grainstones. In this example, some grainstone texture remains, although dark areas, corresponding to zones of higher porosity, cut across bedding planes.

typical dimensions of 1 to 4 cm. These hard limestone layers constitute thin, scattered layers; the bulk of the sediment consists of limestones too porous or weakly cemented to core using the available tools. In intervals where variations in cementation were present, lower resistivity and density values directly corresponded to large hole diameters; soft, poorly indurated material was preferentially removed during the drilling process. Recovery was lowest in strata having intergranular and extreme moldic porosity (packstones and grainstones) and highest in clay-rich or dolomitized strata. For example, as clay content and dolomitization increased downhole at Site 866 from about 600 mbsf, recovery increased to about 50% in the lowest 50 m of the hole.

The FMS log reveals more well-cemented intervals than were recovered as core. Recovery in wackestone-packstone (hard-soft) intervals having thicknesses on the order of 1 to 2 m was particularly poor. In some cored intervals, it was possible to locate the recovered material at its source depth (Tables 1 and 2); a pattern of recovery is suggested such that when coring begins or ends in a well-lithified layer, some of that layer is usually recovered.

Table 3 summarizes the statistical distribution of lithologies interpreted from the core-log integration at Holes 865A and 866A. All beds containing predominantly grainstone (e.g., grainstone-packstone, grainstone-rudstone) were grouped under grainstone; similarly, all beds containing predominantly rudstone (e.g., rudstone-packstone,

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rudstone-grainstone) were grouped under rudstone. Algal-mat-rich intervals and laminite were grouped together. Inaccurate estimations of abundance arise from the minimum practical resolution of the FMS tool (0.5 cm). Clay in recovered core material is distributed as individual fine (millimeter-thickness) seams, intervals of closely spaced fine seams, and as sharply defined centimeter-scale beds. The presence of millimeter-scale clay seams can be distinguished in the best quality FMS images because of the large resistivity contrast with their surroundings, but because of the resolution limitations of the FMS tool, the thicknesses of such seams are falsely imaged as 0.5 to 1.0 cm. The same is probably true of the less common, thin, highly resistive, wackestone intervals and hard grounds. Bed thicknesses of 1 cm or less (as measured in the FMS images) were not included in the statistical analysis.

Packstone is the dominant lithology at Hole 865A (51%), whereas wackestone, which constitutes only 3% of the logged section, was the dominant lithology recovered. Similarly, except for the dolomites at the base of the section, wackestone, wackestone-mudstone, mudstone, and laminite dominated recovery at Hole 866A. The grainstones and packstones that make up 25% and 20% of the section, respectively, were poorly represented in the recovered core material.

SUMMARY OF LOG-CORE INTEGRATION FOR HOLE 865A

Both core samples identified as "speleothem" and the broad range of porosities that characterize the depth interval from 140.0 to 162.0 mbsf suggest that fresh water may have been present in the formation; however, carbon isotope data are inconclusive in this regard (Sager, Winterer, Firth, et al., 1993).

The depth interval from 162.0 to 600.0 mbsf consists almost exclusively of upward-fining sequences of packstone, wackestone, and mudstone. Typically, the packstones are extremely porous and have extensive moldic porosity, and the wackestones and mudstones, with and without moldic porosity, are burrowed. Alternating intervals of rhythmic variations in resistivity and homogeneous zones are common throughout this interval. The character of the gamma, density, and velocity (not shown) logs indicates the presence of many thin layers in the variable zones. The FMS shows numerous thin layers of alternating high and low resistivities on the order of 10 cm, with some as thin as 1 cm.

The first traces of clays were detected in cores at about 600 mbsf, and confirmed by elevated gamma readings. Clay seams of measur-

Table 2. Description of beds at Hole 866A (Resolution Guyot).

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able width are not visible in the FMS images above about 657 mbsf; however, thin (below resolution) clay seams may have caused the stratification observed in the mudstones and wackestones-mudstones beginning at 603 mbsf. Concentrations of clay seams appear to be spaced at 20- to 50-cm intervals in the recovered core, but are spaced farther apart in the FMS images; again, probably because some thicknesses are below the resolution of the FMS tool. Clay seams may be the result of autocyclic processes, such as changing influx paths for terrigenous material or allocyclic fluctuation of sea level, rainfall intensity, carbonate production, or some combination of these (e.g., James, 1989).

Although the depth interval from 600 to 835.5 mbsf also consists of sequences of packstone, wackestone, and mudstone, the relatively greater abundance of mudstones, clay seams, and marls indicates a more restricted lagoonal setting. Grainstone intervals are rare and together with occurrences of rare planktonic species in the core material are probably the result of occasional washover.

In summary, the sedimentary record of Allison Guyot shows small-scale, shallow-water carbonate sequences of (late) Albian age. Facies imply a clay-, organic-, and pyrite-rich, marshlike environment with a nearby volcanic landmass in the lower part of the hole, and restricted, lagoonal settings in its upper part, similar (but not identical) to the Albian section at Hole 866A. The absence of clays above 600 mbsf suggests an expanding lagoon and submergence of the volcanic edifice.

SUMMARY OF LOG-CORE INTEGRATION FOR HOLE 866A

The depth interval from 206 to 466 mbsf consists of sediments derived from a restricted, shallow, subtidal inner shelf. From 206 to 346 mbsf, rudstones and wackestone-packstone sequences are common, with minor mudstone and grainstone. Extensive dissolution is evident at the base of this pile, indicating possible emergence and erosion of the platform. The mudstone-rudstone-wackestone facies common throughout the interval from 346 to 351 mbsf probably indicate a upward-deepening transition. Wackestones, packstones, and especially mudstones are common throughout the depth interval from 351 to 449 mbsf. Another upward-deepening transition is indicated by the packstones and marls that occur within the depth interval from 449 to 466 mbsf.

Tidal flats typified the depositional environment of sediments from the depth interval from 466 to 692 mbsf. The major lithologies consist of wackestone-packstone alternations with scattered abundant grainstones. Several distinct intervals can be distinguished on the basis of the presence or absence of grainstone or mudstone: from 466 to 479 mbsf, wackestone-packstone, grainstones abundant (slight progradation); 479 to 528 mbsf, wackestone-packstone, no mudstones; 528 to 543 mbsf, wackestone-packstone, grainstones abundant (slight progradation); 543 to 587 mbsf, wackestone-packstone, some grainstone and rudstone (period of maximum flooding; aggradation); 587 to 648 mbsf, wackestone-packstone, grainstones abundant (slight progradation); clay seams at 587 to 590 mbsf; 648 to 676 mbsf, wackestonemudstone and laminites (inner platform facies; aggradation); from 676 to 692 mbsf is a transition facies, wackestone-mudstone, rudstone, grainstone (possibly a late highstand stage).

Deepening occurs upward over the interval from 692 to 798 mbsf. The grainstones recovered within the interval from 692 to 791 mbsf reflect a foreshore environment during a highstand stage. The basal contact with underlying well-cemented mudstone is sharp. Slight backstepping is indicated by the mudstones of from 791 to 798 mbsf.

| | | | | | Hole 865A | | | | | | |
|--|--|---|--|---|---|--|--|---|---|--|----------------------------|
| | Wackestone Packstone | | Packstone -wackestone | Mudstone | Mudstone- wackestone Grainstone | | Rudstone | Clay | Marl | Laminite | |
| Min. Max. Sum No. Mean G. | 0.09 5.55 22.56 19 1.19 3 | 0.08 35.98 349.01 92 3.79 51 | 0.34 8.73 59.37 26 2.28 8 | 0.09 4.26 15.26 22 0.69 $\overline{2}$ | 0.32 16.08 198.88 87 2.29 29 | 0.13 7.74 22.18 10 2.22 3 | **** **** **** 米米米米 **** **** | 0.05 1.05 2.13 7 0.30 <1 | 0.43 2.76 11.23 1.60 $\mathbf{2}$ | **** **** **** **** **** **** | |
| | | | | | Hole 866A | | | | | | |
| | Wackestone Packstone | | Packstone -wackestone | Mudstone | Mudstone- wackestone Grainstone | | Rudstone | Clay | Marl | Laminite | Dolomite |
| Min. | 0.03 0.09 | 0.08 0.02 | 0.15 **** | 0.09 0.06 | 0.11 | 0.10 | | | | | 0.38 9.64 |
| Max. Sum No. Mean $\%$ | 5.96 1113.73 108 1.05 8 9 | 7.52 272.95 149 1.83 20 <1 | 6.04 136.34 81 1.68 10 **** | 3.40 49.4 64 0.77 4 | 16.64 103.69 46 2.25 8 | 24.60 339.32 167 2.03 25 | 7.93 122.48 70 1.75 | 1.32 5.12 21 0.24 | **** **** **** **** | 4.73 45.96 51 0.90 | 171.28 75 2.28 13 |

Table 3. Bed thickness statistics for Holes 865A and 866A.

The strata deeper than 798 mbsf reflect the initial emergence and later flooding of the carbonate platform. The general flooding of the platform and evolution of a tidal-flat environment are revealed in the sediments of the interval from 798 to 1144 mbsf. From 798 to 853 mbsf, sediments are lagoonal (restricted shallow subtidal to intertidal) in origin (i.e., wackestone-packstone intervals with much clay). Algal mat disappears at top of this section, indicating a lowstand stage. The grainstones, abundant wackestone, and rudstone of the depth interval from 853 to 932 mbsf indicate a highstand stage, with maximum flooding at about 890 mbsf. From 932 to 1144 mbsf, mudstone, packstone, laminites, and clay reflect a restricted and shallow intertidal to supratidal environment. A grainstone at 1018 mbsf having ripup clasts at its base (visible in the FMS image) is probably a storm deposit; similarly, other minor grainstones in this interval are probably washover.

The depth interval from 1144 to 1442 mbsf consists of shallow subtidal to intertidal facies showing a general shallowing upward of the platform. These facies are dolomitized. From 1144 to 1244 mbsf, wackestone-packstone sequences and abundant grainstone intervals indicate alternating progradation and aggradation. Throughout the interval from 1244 to 1442 mbsf, common grainstone with abundant wackestone-packstone indicate a highstand stage.

Two major intercalations of oolitic limestone are developed, one of latest Hauterivian age (Jenkyns et al., this volume), which rests on the basaltic edifice, and a second of Aptian age (Sager, Winterer, Firth, et al., 1993; Jenkyns et al., this volume) that is sandwiched between lagoonal-peritidal sediments (692-798 mbsf). The Hauterivian coarse, Unsorted grainstones at the base of the section (>1442 mbsf) were deposited on an open-marine shelf or ramp.

In summary, FMS and geophysical logs of porosity and porosityrelated parameters (Fig. 10) confirm that the sedimentary record is composed of approximately 1- to 10-m sequences. Where resolution is good, smaller-scale sequences can be defined in the FMS logs, as well; often, however, the contrast is insufficient to identify them properly. These sequences are well developed at this site in the platform interior, where they probably reflect the cyclic deepening and shallowing of the depositional environment. From the lithostratigraphic interpretation of the logs and core information, we hypothesize that sediments of the upper 600 mbsf of Hole 866 A originated in a somewhat restricted, tidal or subtidal environment, with oxygenated waters suitable for supporting a varied fauna. The predominant lithologies are upward-fining packstone-wackestone-mudstone sequences that have some significant grainstone intervals. Strasser (this volume) suggests a potential hiatus or condensed section at 480 mbsf, based on occurrence of calcrete horizons; the FMS logs place these horizons at a slightly different depth, about 465 mbsf. From about 600 mbsf downward,

lithologies indicate a restricted lagoonal environment, punctuated by intervals of slightly more open lagoonal environment of up to 100 m thick (750-850, 880-930,1180-1230, and 1260-1340 mbsf). Normal open-marine conditions prevail below 1340 mbsf.

COMPARISON OF ALBIAN SECTIONS AT SITES 865 AND 866

Beds for sections of Albian age were grouped into upward-fining (e.g., packstone-wackestone-mudstone) or upward-coarsening sequences. A plot of the thicknesses of upward-fining sequences vs. their centered depths (halfway between base and top) for both holes is shown in Figure 11. The plots should be regarded as qualitative and are not meant to provide the basis for a quantitative comparison; some beds contain numerous interbeds of clay or laminite, indicating numerous minor cycles, for example, and no attempt was made to measure beds thinner than 1 cm. Cycle thicknesses for Hole 865A are generally greater than those for Hole 866A; this is in agreement with the greater sedimentation rates suggested for Hole 865A (Sager, Winterer, Firth, et al., 1993). The thickest sequence measured for Hole 865A, located from 364.1 to 409.2 mbsf (packstone grading up to wackestone with extreme moldic porosity), is close in depth to that of Hole 866A, 411.9 to 442.0 mbsf (packstone grading up to mudstone with extreme moldic porosity). Could these two sequences be equal in time, especially because sequences above this level are generally longer for both holes, and sequences below this level are generally shorter for both holes? Evidence in support of this contention is found in logging "events" at corresponding depths at both sites. These events consist of high-amplitude excursions from prevailing values, usually in the resistivity log, occasionally seen in the natural gamma-ray log. At Site 865 (Fig. 9), such events occur at 240, 272, 305, and 320 mbsf, whereas at Site 866 (Fig. 10), similar events occur at 280,325,350, and 363 mbsf. Thus, deposition of the lagoonal sediments that make up the Albian section at Sites 866 and 865 probably overlapped somewhat. If this is the case, then, from Figure 11, it appears that much of the late Albian section above the 400-mbsf level is missing at Hole 865A, perhaps removed as the guyot emerged above sea level. At Hole 866A, a hiatus probably exists at the base of the Albian section (Strasser, this volume).

CONCLUSIONS

Deep holes drilled into the Cretaceous lagoonal facies of Allison (Site 865) and Resolution (Site 866) guyots yielded samples of thick sections of shallow-water limestones that record the histories of the

Figure 9. Profiles of (from left) medium electrical resistivity (ILM), natural gamma-ray intensity (NGT), formation density (RHOB), and caliper (HLDT), determined from downhole logs at Hole 865A. Cored intervals and recovery (in black) are shown in the right columns. The exact position of the cored material within the cored intervals is not constrained; recovered core material was pushed to the top of each cored interval for consistency. Also shown are lithologic interpretations based on integration of the FMS and geophysical logs with core descriptions; the pattern legend is shown in Figure 8.

guyots from initial submergence of the volcanic pedestal through the final drowning of the carbonate platform. However, core recovery was poor, and we suspect that not all lithologies were sampled and, further, that studies of widely spaced core samples would not provide a sufficiently detailed record of the vertical trajectories of the guyots for correlating directly between them. As downhole logs provide a more continuous record of the physical properties of the borehole, the integration of logging data with core descriptions presents a lithologic record that is both detailed and accurate.

Facies that evoked similar electrical responses in the conventional and FMS logs were calibrated using laboratory descriptions and photographs of recovered core materials to produce "type" examples for the various carbonate facies. The type examples then were used to convert the remaining logging data into lithologic columns. We saw immediately that core recovery at both sites was highly preferential (e.g., within the alternating packstone-wackestone intervals, only small pebbles of well-cemented wackestone were recovered). This study indicated that packstone was the dominant lithology (51%) at Hole 865A, and that grainstone (25%) and packstone (20%) were the dominant lithologies at Hole 866A (excluding dolomite). Recovery of packstone and grainstone was poor.

The sedimentary record from Allison Guyot, as determined from core-log integration, confirmed the existence of numerous smallscale shallow-water carbonate sequences of (late) Albian age. Facies from the lower part of the hole imply a clay- and organic-rich, marshy environment with a nearby landmass, gradually opening uphole into

Figure 9 (continued).

a somewhat restricted lagoonal setting, similar to the Albian (200- 550 mbsf) section at Hole 866A.

Quality of logs was better at Resolution Guyot because of better hole conditions. The sedimentary record at Hole 866A was composed of sequences 1 to 10 m long. Where resolution was particularly good, smaller-scale sequences could be defined in the FMS logs as well, although the contrast was often insufficient to identify them properly. Sediments in the upper 600 mbsf of Hole 866A originated in a somewhat restricted tidal or subtidal environment. Lithologies at depths greater than 600 mbsf imply a restricted lagoonal environment with occasional periods of more open lagoonal conditions. Lithologies below 1340 mbsf suggest normal open-marine conditions.

Lithologies were grouped into upward-fining sequences at both localities.

Sequence thicknesses for Hole 865A generally are greater than those for Hole 866A; this is in agreement with the greater sedimenta-

tion rates suggested for Hole 865A (Sager, Winterer, Firth, et al., 1993). The thickest sequence measured for Hole 865A, located at 364.1 to 409.2 mbsf (packstone grading upward to wackestone with extreme moldic porosity), corresponds in depth to that of Hole 866A at 411.9 to 442.0 mbsf (packstone grading upward to mudstone with extreme moldic porosity). Support for this contention is found in the existence of several smaller-scale, but significant, logging "events" found at similar depths in both holes. At both sites, this thick sequence defines a change in either depositional style or subsidence rate as sequences above this level are generally thicker, and sequences below this level are generally thinner, for both holes. This study supports Strasser's conclusion (this volume) that a hiatus probably exists near the Albian/Aptian boundary (~480 mbsf). Correlations of bed thicknesses and logging signatures between the two holes indicate that much of the Albian section may have been removed from Allison Guyot as it emerged above sea level.

Figure 9 (continued).

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Figure 10. Profiles of (from left) medium electrical resistivity (ILM), natural gamma-ray intensity (NGT), formation density (RHOB), and caliper (HLDT), determined from downhole logs at Hole 866A. Cored intervals and recovery (in black) are shown in the right columns. The exact position of the cored material within the cored intervals is not constrained; recovered core material was pushed to the top of each cored interval for consistency. Also shown are lithologic interpretations based on integration of the FMS and geophysical logs with core descriptions; the pattern legend is shown in Figure 8.

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Figure 11. Comparison of upward-fining sequence thicknesses within the Albian portion of Hole 866A (top of logged section to 500 mbsf) and Hole 865A, which is entirely Albian in age.

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