# 42. SYNTHESIS OF TERRIGENOUS ACCUMULATION RATES AND BIOSTRATIGRAPHIC STUDIES AT SITES IN THE NORTHWESTERN PACIFIC OCEAN, WITH COMPARISONS TO ADJACENT REGIONS OF THE PACIFIC GYRE<sup>1</sup>

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#### ABSTRACT

This paper presents the integrated results of quantitative sedimentologic and biostratigraphic studies at Sites 871, 872, and 873 in the Marshall Islands, Hole 880A on Takuyo-Daisan Guyot, and Hole 810C on Shatsky Rise for the late Pliocene to Holocene. Published results from studies of terrigenous mineral (eolian) accumulation at additional Ocean Drilling Program, Deep Sea Drilling Project, and piston core sites located around or within the boundaries of the (Northern Hemisphere) Pacific Subtropical Gyre are used to provide additional regional comparisons among high-quality, high-resolution eolian records from the Pacific Basin.

Two time intervals have been selected for these regional comparisons. The first interval selected is from 3 million years ago (Ma) to the present, which allows for comparisons to be made among several sites which have relatively low temporal resolution records. The second interval is from 600 ka to the present, which allows for comparisons to be made among the small number of high-quality, high-resolution, proxy records from sites in the northwest and equatorial Pacific Ocean. Additional discussions will focus on the potential for correlating several marine records of terrigenous dust accumulation at sites located downwind from Asia beneath the Pacific Subtropical Gyre with the continental climate record preserved in alternating loess and soil deposits of Central China.

#### INTRODUCTION

The goal of this paper is to combine the results of recent sedimentologic and biostratigraphic studies of cores collected during Ocean Drilling Program (ODP) Leg 144. We will discuss the sedimentologic results from two guyots in the Marshall Islands (Lo-En and Wodejebato guyots; Holes 871A and 873B; Rack, Bohrmann, and Hobbs, this volume) and from Takuyo-Daisan Guyot (Hole 880A; Rack, Lawyer, and Gee, this volume), and the results from quantitative studies of nannofossil biostratigraphy (Erba, this volume) and studies of siliceous microfossils (Fenner, this volume) from Quaternary sediments at these same sites. The sedimentologic and biostratigraphic data will be used to discuss the timing of terrigenous mineral fluxes to the northwest and southern to southwest regions of the (Northern Hemisphere) Pacific Subtropical Gyre, and the relationships between mineral fluxes, nannofloral assemblage shifts, and paleoclimatic changes.

Sediment mass accumulation rates (MAR) of different sedimentary components have been calculated at each of these sites to allow for comparisons across a wide geographic region, primarily using sites located at intermediate water depths. The calculated flux data for Hole 880A are given in Appendix A of this paper; the flux data for Holes 871A and 873B are slightly modified from Rack, Bohrmann, and Hobbs (this volume). The combined results from these Leg 144 sites are compared with the published results from sedimentologic, magnetic, and biostratigraphic studies at Hole 810C (Premoli Silva et al., 1993; Rack et al., 1993; Sager et al., 1993) and with unpublished mineral flux data from this same hole on Shatsky Rise (see Appendix B; Rack and Janecek, unpubl. data). The results of these laboratory analyses and the age model derived from an integrated magnetostratigraphy and biostratigraphy for Hole 810C (Premoli Silva et al., 1993; Sager et al., 1993) have been used to calculate MARs for sedimentary components for the time interval from late Pliocene to present (see Appendix C).

Published results from mineral (eolian) studies in the northwestern and equatorial to tropical Pacific Ocean have been utilized to produce summary figures of the weight percentages and mass accumulation rates of individual sedimentary components (for site locations, see Table 1 and Fig. 1) to allow for effective regional comparisons among sites. These summary figures show the temporal resolution of data from previous eolian studies around the Pacific Subtropical Gyre for the late Pliocene to Quaternary. The figures can be used to evaluate potential differences among the various age models for these records, and can provide a glimpse of the present state of our understanding of these proxy records, which are used to identify changes in atmospheric circulation and source region aridity around the Pacific Basin (see review by Rea, 1994).

Two discrete time intervals (0–3 Ma and 0–600 ka) are presented in this paper. These two intervals were selected to assess our current understanding of late Pliocene to Quaternary terrigenous mineral (eolian) records from two widely separated regions of the Pacific Subtropical Gyre: one in the equatorial to tropical North Pacific, and one in the northwest Pacific Ocean.

In the first part of this synthesis, two transects of sites in the tropical North Pacific Ocean are shown: one transect in the western equatorial Pacific (Hole 803D, Krissek and Janecek, 1993; Holes 871A, 872A, and 873B; Rack, Bohrmann, and Hobbs, this volume), and one transect in the eastern equatorial Pacific (Holes 848B, 849B, 852B, and 853B; Hovan, 1995). These two transects of sites are compared for the interval 0–3 Ma. In addition, these low-resolution records are compared with the high-resolution records from piston Core RC11-210 (Chuey et al., 1987; Rea et al., 1991), located at a site between these two transects in the central equatorial Pacific Ocean.

The second part of this synthesis presents a discussion of terrigenous records from the Northwest Pacific Ocean. Two new records are presented: one from Hole 880A (Rack, Lawyer and Gee, this volume; Appendix A) and one from Hole 810C (Rack et al., 1993; Appendixes B and C). The sediment record from Hole 880A is considered to be a hemipelagic record, with significant quantities of volcaniclastic material and possibly terrigenous material transported

<sup>&</sup>lt;sup>1</sup> Haggerty, J.A., Premoli Silva, I., Rack, F., and McNutt, M.K. (Eds.), 1995. Proc. ODP, Sci. Results, 144: College Station, TX (Ocean Drilling Program).

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Site	Latituda	Longituda	Water depth	Poforance to data	Site ID No.
(mole)	Lanuae	Longhude	(most)	Reference to data	(11g. 1)
KK75-02	38°37.00'N	179°20.00'E	5475	Janecek (1983)	1
V21-146	37°41.00'N	163°02.00'E	3968	Hovan et al. (1989, 1991)	2
310	36°52.00'N	176°54.00'E	2525	Rea and Janecek (1982), Janecek (1983)	3
880A	34°12.53'N	144°18.74'E	1525	Rack et al. (this volume; this paper)	4*
578	33°55.56'N	151°37.74'E	6010	Janecek (1985)	5**
810C	32°25.40'N	157°50.44'E	2623	Rack et al. (1993)	6
576	32°21.40'N	164°16.50'E	6200	Janecek (1985), Leinen (1985)	7
463	21°21.00'N	174°40.00'E	2525	Rea and Janecek (1981, 1982)	8
873B	11°53.84'N	164°55.23'E	1334	Rack et al. (this volume)	9
872A	10°05.85'N	162°51.96'E	1084	Rack et al. (this volume)	10
871A	5°33,43'N	172°20.66'E	1255	Rack et al. (this volume)	11
803D	2°25.98'N	160°32.46'E	3412	Krissek and Janecek (1993)	12
LL44-GPC3	30°20.00'N	157°49.00'W	5705	Leinen and Heath (1981), Janecek and Rea (1983)	13
853B	7°12.66'N	109°45.08'W	3716	Hovan (in press)	14
852B	5°17.57'N	110°4.58'W	3860	Hovan (in press)	15
503B	4°03.02'N	95°38.32'W	3672	Rea and Janecek (1982), Rea et al. (1986)	16
RC11-210	1°49.00'N	140°03.00'W	4420	Chuev et al. (1987), Rea et al. (1991)	17
849B	0°10.98'N	110°31.18'W	3840	Hovan (in press)	18
848B	2°59.63'S	110°28.79'W	3856	Hovan (in press)	19

Table 1. Selected sites where eolian studies have been conducted in the Pacific Ocean.

Notes: Single asterisk (\*) = Hole 880A is considered a hemipelagic site, which contains significant quantities of continentally derived and volcanic material. Caution should be exercised when comparing the results obtained from this hole with other sites on this list. Double asterisks (\*\*) = Site 578 is also considered to be a hemipelagic site (see Rea, Basov, Janecek, Palmer-Julson, et al., 1993). Site identification numbers are provided to help the reader locate these sites in Figure 1. mbsl = meters below sea level.

as runoff from Japan and/or China; therefore, we will not discuss this site in the context of eolian transport of grains.

The record of terrigenous mineral accumulation at Hole 810C on southern Shatsky Rise is combined with prior results from this hole (Natland, 1993; Polgreen et al., 1993; Premoli Silva et al., 1993; Rack et al., 1993; Sager et al., 1993), which are compared with the results of published eolian studies from piston Cores V21-146 (Hovan et al., 1989, 1991; northern Shatsky Rise) and KK75-02 (Janecek, 1983; Janecek and Rea, 1984, 1985; north central Pacific Ocean). These marine records are then compared with the continental record of climate represented by fluctuations in magnetic susceptibility measurements from the Xifeng I loess site in Central China (Kukla et al., 1988, 1990; Liu et al., 1988a, 1988b).

# BACKGROUND AND PREVIOUS WORK

During the past 20 years a large number of eolian studies have been conducted, using marine sediment samples to examine the linkages between atmospheric and oceanographic processes over millions of years, and at shorter time scales (Rea, 1994). Studies in the north Pacific Ocean have emphasized the importance of understanding changes in the relative intensity and positions of the Northern Hemisphere westerlies and equatorial trade wind belts, which control the grain size distribution of transported mineral dust from continental source regions, and the role of continental aridity in determining the magnitude of the terrigenous dust flux to the Pacific Ocean.

The complex proxy environmental signals provided by fluctuations in eolian mass accumulation rates and changes in mineral grain size are linked to a wide array of processes, some better understood than others, including: (1) seasonal or interannual variability of eolian and atmospheric transport processes (Arimoto, et al., 1985; Betzer et al., 1988; Blank et al., 1985; Duce et al., 1980, 1983, 1991; Merrill et al., 1989; Uematsu et al., 1983, 1985; among others); (2) climatic or orbital (tens to hundreds of thousands of years) fluctuations in the geographic position and intensity of the mid- and low-latitude wind belts (westerly and easterly wind systems, respectively) and the Intertropical Convergence Zone (Hovan, 1995; Janecek and Rea, 1983, 1984, 1985; Rea, 1990, 1994; Rea et al., 1985; Wyrtki, 1974; among others); (3) temporal changes in the aridity or seasonality of continental source regions where eolian dust originates (Leinen, 1989; Prospero, 1981; Phillips et al., 1993; Pye and Zhou, 1989; Rea and Leinen, 1988; Windom, 1976; Zhang et al., 1994; among others); and (4) temporal fluctuation in the atmospheric transport of volcanic

particles to these regions (Furuta, et al., 1986; Kamata, 1989; Kennett and Thunell, 1975; Kennett et al., 1977; Machida and Arai, 1983, 1988; Natland, 1993).

Some of the initial eolian studies conducted on Deep Sea Drilling Project (DSDP) cores were focused on the aseismic rises of the North Pacific Ocean (Hess Rise, Shatsky Rise, and the Mid-Pacific Mountains), which are located above the nepheloid layer, south of the latitude of ice-rafted sedimentation during both glacial and interglacial times, and which lie far from the continental sources of hemipelagic deposition in the Pacific (Rea and Janecek, 1981). Marine researchers have explained Cenozoic fluctuations in mineral mass accumulation rates at these sites as having resulted from changes in the aridity of the source regions of the mineral dust (Janecek and Rea, 1983, 1984; Leinen and Heath, 1981; Rea and Janecek, 1981, 1982; also see Rea, 1994), and they have explained changes in the median grain size of the mineral particles as having resulted from changes in the energy or strength of the transporting medium (atmospheric winds).

From the late Miocene to the Quaternary, there has been a significant increase in the accumulation of eolian (inorganic crystalline) material at several mid-Pacific sites (e.g., DSDP Sites 310 and 463, and giant piston Core LL44-GPC3; Leinen and Heath, 1981; Rea and Janecek, 1981, 1982; Janecek and Rea, 1983), which has been explained by a combination of factors, including the following: tectonic motions that have transported sites through or into the influence of various wind belts, Northern Hemisphere climatic deterioration and the subsequent intensification of atmospheric circulation as continental ice-sheets expanded, and periodic pulses of global or regional volcanism (Kennett and Thunnell, 1975; Kennett et al., 1977; Rea and Scheidegger, 1979).

This synthesis attempts to bring together the existing data from a number of sites around the North Pacific Ocean where late Pliocene to Quaternary mineral dust (eolian) studies have been undertaken. Where possible, the published data from these studies have been supplemented to construct an internally consistent set of figures with which to compare similar time intervals at various site locations around the Pacific Subtropical Gyre.

## METHODS

### **Chemical Analyses and Mineral Extractions**

Samples from Holes 871A, 873B, 880A, and 810C (Fig. 1) were freeze-dried, weighed, and used for laboratory chemical analyses to



Figure 1. Location map showing selected sites used for eolian studies in the Pacific Ocean. The numbers on this map correspond to the appropriate sites listed in Table 1.

determine the weight percentages of inorganic carbon and calcium carbonate, and the weight percentage of the fine-grained lithogenic fraction (<63-µm-sized mineral grains) of the bulk sediment (see discussion of methods and data contained in Rack et al., data reports, this volume; and Rack et al., 1993).

The inorganic carbon content of each sample was determined using a Coulometrics carbon dioxide coulometer. Calibration was performed using pure calcium carbonate as a standard. The weight percentage of carbonate was calculated from the inorganic carbon (IC) content, assuming that all carbonate was in the form of calcite: weight percentage of  $CaCO_3 = \%IC \times 8.332$ .

Following the determination of inorganic carbon, a split of each sample was weighed and then processed according to a lithogenic extraction procedure, as described by Rea and Janecek (1981) and Clemens and Prell (1990). The lithogenic component of the bulk sediment (including contributions from both terrigenous and volcanic sources) was isolated by a series of selective extractions to remove calcium carbonate, opaline silica, Fe and Mn oxides, hydroxides and zeolites. The material passing through a <63- $\mu$ m sieve after the completion of the extraction procedure was weighed to determine the weight percentage of the fine-grained terrigenous or mineral fraction (<63  $\mu$ m) of the total sample.

Throughout this paper, the extracted silt + clay-sized inorganic mineral fraction of the bulk samples (including both lithogenic and volcanic grains <63  $\mu$ m in size) are referred to as either the "finegrained terrigenous" or "fine-grained lithogenic" fraction of the bulk sediment; these designations do not necessarily indicate an eolian transport mechanism for this material. In general, the fine-grained lithogenic fraction of these samples is dominated by clays, finegrained quartz, and grains of probable volcanic origin.

#### Sediment Mass Accumulation Rate Calculations

Dry bulk density values were determined for each sample at Holes 871A, 873B, and 810C using either laboratory data, or regressions between laboratory measurements of dry bulk density and the wet-bulk density data collected by the GRAPE (gamma-ray attenuation porosity evaluator) at each site (see Rack et al., 1993; Rack, Bohrmann, and Hobbs, this volume).

Because of the scarcity of physical properties measurements at Hole 880A, estimates of dry bulk density were determined using a regression derived at Hole 810C between laboratory dry bulk densities and GRAPE bulk density values. The bulk sediment recovered in Hole 810C is lithologically similar to that recovered in Hole 880A (calcareous ooze), although Hole 810C has significantly less coarsegrained material (lithogenic and volcanic grains) than is found at Hole 880A. The abundance of coarse-grained material at Hole 880A leads us to classify it as a hemipelagic site that has been influenced by late Neogene volcanism and by possible surface runoff from Japan and/or China. A comparison of the six laboratory determinations of dry bulk density at Hole 880A and predicted dry density values, which were determined from the GRAPE bulk density record at Hole 880A and the regression equation from Hole 810C, indicate that the predicted values are about 10%-20% higher than the laboratory values. Calculated accumulation values are slightly high because of this error, but errors in linear sedimentation rate determinations are likely to be of equal or greater magnitude.

The product of the linear sedimentation rate (LSR = cm/k.y.) and the dry bulk density (DBD = g/cm<sup>3</sup>) equals the sediment mass accumulation rate (MAR = g[cm<sup>2</sup>×k.y.]<sup>-1</sup>). The MAR is multiplied by the percentage of each sedimentary component to give MARs for the carbonate (CAR), noncarbonate (NCAR), and terrigenous (TMAR) components of the total flux at each site. The values of terrigenous mass accumulation rate (TMAR) are presented in units of mg(cm<sup>2</sup>×k.y.)<sup>-1</sup>.

When possible, the weight percentage and accumulation of the "unknown" portion of the bulk sediment has been estimated. This has been accomplished by assuming that the total sample equals 100%, and then subtracting the weight percentages of various "known" components to determine what is unknown. This calculation is presented to show the possible importance of various components of the total sedimentary flux (coarse-grained lithogenic grains, biogenic opal, Fe and Mn oxides, hydroxides and zeolites, and organic carbon), which are often not accounted for in the available chemical data provided by published studies.

## **Quantitative Biostratigraphic Studies**

Erba (this volume) presents quantitative studies of calcareous nannoflora from Sites 871, 872, 873 and 880 using the informal morphometric taxonomic concepts described by Rio (1982), Rio et al. (1990), and Raffi et al. (1993) within the genus *Gephyrocapsa*. The *Gephyrocapsa* group shows very high abundances with respect to the other groups of small-sized coccoliths (non-*Gephyrocapsa* species of nannoflora) of Pleistocene age. The identification of *Gephyrocapsa* events has been shown to be highly useful for subdividing Pliocene to Quaternary age sequences for high-resolution studies of biostratigraphy and paleoceanography (Raffi et al., 1993; Wei, 1993).

Several recent studies have highlighted the use of these morphometric taxonomic concepts in nannoflora biostratigraphy, which can be integrated with correlations to oxygen isotope records and astronomical calibration (see studies by Castradori, 1993; Wei, 1993; Raffi and Flores, 1995; among others).

For the Leg 144 sites, a polarizing light microscope was used to count at least 500 specimens in a smear slide made from each sample (Erba, this volume). The percentages of single classes within the genus *Gephyrocapsa* were calculated from the original counts of total nannofloras. At Hole 810C, Premoli Silva et al. (1993) have conducted a similar, although semi-quantitative biostratigraphic study, using similar taxonomic methods. These data are used to provide age models for use in the calculation of sedimentary fluxes at Sites 871, 873, 880 and 810.

The ages that have been assigned to the biostratigraphic datums used in this study (see Table 2) are from a recent late Neogene synthesis of sites in the eastern equatorial Pacific Ocean (i.e., ODP Leg 138) by Shackleton et al. (1995). We assume that the ages of the nannofossil datums are synchronous for the geographic area studied here (Erba, this volume). The validity of this assumption will need to be examined in future studies. Studies of calcareous nannoflora and orbitally calibrated oxygen isotope records from the North Atlantic and equatorial Pacific Ocean demonstrate the range of potential errors related to this assumption of synchroneity (Raffi et al., 1993; Wei, 1993).

# **RESULTS AND DISCUSSION**

This paper presents the combined results of studies of the finegrained terrigenous (mineral dust) component of pelagic sediment from the Marshall Islands (Holes 871A, 872A, and 873B on Limalok, Lo-En, and Wodejebato guyots, respectively; Rack, Bohrmann and Hobbs, this volume) and on Takuyo-Daisan Guyot near Japan (Hole 880A; see Rack, Lawyer and Gee, this volume; Fenner, this volume), and quantitative studies of Quaternary nannofossil biostratigraphy from these same sites (see Erba, this volume). Siliceous microfossils are restricted to the uppermost few meters of the pelagic cap sediments on each of these guyots, whereas late Pliocene to Quaternary calcareous nannofloras are found to be very abundant and well preserved at these sites.

Published results of terrigenous mineral (eolian) studies from additional ODP and piston core site locations from the boundaries and interior of the (Northern Hemisphere) Pacific Subtropical Gyre are used to provide additional comparisons among several high-quality mineral flux records. The high-resolution mineral record from Hole 810C on Shatsky Rise will be highlighted, as it provides a long-term (>3 m.y.) marine eolian record that can potentially provide correlations to the continental record of climate preserved by alternating soil/loess horizons in Asia.

## **Results from Tropical Pacific Sites**

#### Sites 871, 872, and 873

Site 871 is located on the south central portion of Limalok Guyot, in the Ratak Chain of the southern Marshall Islands (Table 1 and Fig. 1). The sediments in the pelagic cap at Site 871 (0–133.7 meters below seafloor [mbsf]) consist of foraminifer nannofossil ooze and foraminifer ooze of Pleistocene to early Miocene age. The pelagic cap is divided into two subunits by a disconformity at 26.5 mbsf, where overlying light gray nannofossil foraminifer ooze of Pleistocene to latest Miocene age changes to white homogeneous foraminifer ooze of middle to early Miocene age below. The sediment in the lower portion of the pelagic cap is well-sorted and winnowed, with a medium-sand grain size. The disconformity, spanning the upper middle Miocene through most of the upper Miocene (approx. 6 m.y.), marks a major change in sediment accumulation on this guyot (Premoli Silva, Haggerty, Rack, et al., 1993), which may correspond to an intensification of intermediate water flow in response to cooling at high southern latitudes.

The pelagic cap at Site 872 on Lo-En Guyot (Table 1 and Fig. 1) consists of an upper lithologic subunit of white to very pale brown nannofossil foraminifer ooze and intercalated foraminifer ooze of Pleistocene to late Miocene age (30–32 m thick), and a lower subunit of very pale brown homogeneous foraminifer ooze, which is well sorted, winnowed, with a medium-sand grain size, and is of late Miocene to late Oligocene age. The transition between the two subunits is marked by a physically mixed and incomplete interval of nannofloras and faunas of different zones, which spans the upper middle to lower upper Miocene, overlain by a disconformity (Premoli Silva, Haggerty, Rack, et al., 1993).

Site 873 is situated on the south central summit of Wodejebato Guyot, in the Ralik Chain of the northern Marshall Islands (Table 1 and Fig. 1). The pelagic cap on Wodejebato Guyot consists of an upper lithologic subunit of light gray nannofossil foraminifer ooze of Pleistocene to latest Miocene age; whereas, the lower subunit is composed of a well-sorted and winnowed, very pale brown, mainly homogeneous foraminifer ooze with a medium-sand grain size, and it is of middle to early Miocene age. The transition between the two subunits is marked by a disconformity spanning most of the upper middle and upper Miocene (Premoli Silva, Haggerty, Rack, et al., 1993).

Diatoms at Sites 871, 872, and 873 are scarce, poorly preserved and exhibit low diversity (Fenner, this volume). Conversely, abundant nannofossils were observed throughout the Quaternary sequences at Sites 871, 872, and 873. Because of the relatively shallow water depths of the studied sites, nannofloras are well preserved to moderately preserved (Erba, this volume).

#### Sedimentologic and Biostratigraphic Results

Weight percentages and MARs of calcium carbonate and finegrained mineral material at Holes 871A and 873B are given by Rack, Bohrmann, and Hobbs (this volume). Determinations of the weight percentage of calcium carbonate were also made at Site 872, allowing the noncarbonate weight percentage of the bulk sediment to be qualitatively estimated (100% - wt% CaCO<sub>3</sub> = noncarbonate); however, no mineral extractions were made at this site.

Measurements of the weight percentage of the CH<sub>3</sub>COOH-insoluble residue (>2  $\mu$ m) were also made by Fenner (this volume) at Holes 871A, 872A, 872C, and 873B. This residue will contain mineral dust, ash, opal, and organic carbon. These measurements, although limited to the upper core at each hole, are in general agreement with the qualitative trends presented by Rack, Bohrmann and Hobbs (this volume) for the weight percentage of non-carbonate material.

Erba (this volume) reports a complete sequence of nannofossil events and biozones at Sites 871 and 872; however, at Site 873, most of the Quaternary is missing. The detection of only the lower Pleistocene nannofossil events at Site 873 suggests that the accumulation of pelagic sediments on Wodejebato Guyot is strongly affected by current activity (Premoli Silva, Haggerty, Rack, et al., 1993; Erba, this volume; Pearson, this volume). For depth correlations of the nannofloral zonal boundaries and more detailed discussion of the data from

Table 2. Sub-bottom depth of	f nannofossil datums	identified at selected	Pacific Ocean ODP	sites of relevance t	to this study
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	Event	Age (Ma)	Hole 871A (mbsf)	Hole 872A (mbsf)	Hole 873B (mbsf)	Hole 880A (mbsf)	Hole 810C (mbsf)	Hole 803D (mbsf)	Site 853 (mcd)	Site 852 (mcd)	Site 849 (mcd)	Site 848 (mcd)
7628		(1.147)	(most)	(11031)	(most)	(most)	(indat)	(incor)	(inco)	(incu)	(110.0)	(unita)
FO	Acme E. huxleyi	0.085	0.82	0.25		1.80	0.41					
FO	E. huxleyi	0.260	2.90	1.35		4.80	2.61					
LO	P. lacunosa	0.461	5.30	2.35		5.20	5.66	18.24	2.75	4.80	13.30	6.05
Reentrance	Medium Gephyrocapsa	1.028	9.25	4.75		10.20		23.63	4.03	11.68	28.81	16.37
LO	Large Gephyrocapsa	1.240	12.25	6.25		11.59	21.67	27.61		14.68	33.90	19.55
FO	Large Gephyrocapsa	1.457	13.35	7.30	3.07	13.17	24.84	29.90	5.93	18.30	39.86	23.20
LO	C. macintyrei	1.597	16.35	7.55	3 57	14.57	28.28	32.90	6.68	20.28	44.05	24.32
FO	Medium Genhyrocansa	1.660		1100	0.00		-01-0	37.40	7 32	20.28	47.05	24.47
FO	G. oceanica s.l.	1.660	17.25	8.35	4 02	15.97		27.10	1.0.2	20120		
LO	D. brouweri	1.960	18 35	8.85	4 97	16.37	28.28	43 71	7.70	21.73	59.35	26.92
LO	D. pentaradiatus	2.440	10000	0.00	7.84	10.07	act is an o	49 35	11.17	28.92	07100	30.65
LO	D. surculus	2 610			1.0.1		36.17	15100	13.05	32.83	71.00	30.65
LO	D. tamalis	2 760	1915		10.63		40.55		14.45	34 33	77.20	32.09
10	Sphenolithus spp	3 650	17.15		10.05		40.55		20.50	41 32	101.85	36.15
10	R nseudoumbilicus	3 804			17.03		57 78		20.50	42.85	105.25	37 42
FO	Circular D asymmetricus	4.030			17.05		57.70		22.70	43.80	100.000	38 70
10	A primus	4.540							25.99	51.35		40.26
10	C acutus	4.070							20.15	21.33		47.20
FO	C. maosus	5.040					60.01		29.13	56 15	157.25	50.00
FO	C. rugosus	5.040					69.01	77.07	29.14	50.45	137.23	50.90
10	C. acuius	5.540			21.21			11.91	32.68	03.24	174.30	52.45
LO	1. rugosus	5.540			21.31			70.01	22.40	66.00	174.30	57 10
LO	D. quinqueramus	5,560		21.75			75.21	79.81	32.68	00.32	189.65	57.13

Notes: Nannofossil age model for Holes 871A, 872B, 873B, and 880A by Erba (this volume); for Hole 810C by Premoli Silva et al. (1993); for Hole 803D by Takayama (1993); and for Sites 853, 852, 849, and 848 by Raffi and Flores (1995) and Shackleton et al. (1995). mcd = meters composite depth. FO = first occurrence datum, and LO = last occurrence datum.

these sites, the reader is directed to the descriptions in Erba (this volume). Table 2 provides a summary of the ages and sub-bottom depths used to construct the age models and to plot the data from these holes.

In Hole 871A on Limalok Guyot, there are three prominent increases in MAR at ~1.5 Ma, 1.2 to 1.0 Ma, and ~0.4 Ma, respectively (Fig. 2A). The terrigenous mineral flux (TMAR) is observed to be highest during the ~1.5 Ma and ~0.4 Ma increases in the overall rate, when the TMAR increases to 2-3 times above its background flux values (from ~3 mg[cm<sup>2</sup> × k.y.]<sup>-1</sup> to 6–9 mg[cm<sup>2</sup> × k.y.]<sup>-1</sup>). The nannofloral abundance plots (Fig. 2B) show that the 0.4 Ma TMAR increase is positioned between a sharp decrease in the percentage of G. caribbeanica and an increase in the abundance of E. huxlevi. A qualitative increase is observed in the weight percentage of the >2 µm-sized CH<sub>3</sub>COOH-insoluble residue within this interval (Fenner, this volume), yielding an increased overall flux of insoluble material. The TMAR increase at ~1.5 Ma is associated with an overall increase in the abundance of the Gephyrocapsa group and is positioned just below the increase in the percentage of the large Gephyrocapsa species. These quantitative measures of nannofossil species abundance fluctuations may provide the means to evaluate changes in nannoflora with respect to changes in mineral fluxes, thereby constraining the ecologic and oceanographic interpretation of pelagic sediment accumulation patterns on these guyots.

In Hole 872A on Lo-En Guyot, there are two prominent increases in the noncarbonate flux (NCAR), at ~1.7–1.5 Ma and at ~1.2–1.0 Ma, respectively. The former increase in noncarbonate flux (e.g., mineral component, plus biogenic opal, organic carbon, oxides, hydroxides and zeolites) represents an approximate doubling of the flux above background levels; the latter increase is more subtle (Fig. 3A). The increase in noncarbonate flux at ~1.7–1.5 Ma is associated with both an increase in the overall abundance of the *Gephyrocapsa* group and a subsequent increase in the percentage of large *Gephyrocapsa* species (see Fig. 3B).

In Hole 873B on Wodejebato Guyot, the upper part of the Quaternary is missing but there is a broad increase in TMAR from about 3.0–1.2 Ma. However, a subtle increase in mineral flux is again observed at ~1.7–1.5 Ma (Fig. 4A). There is an increase in the overall abundance of the *Gephyrocapsa* group since before about 2.0 Ma, but the subsequent increase in the percentage of the large *Gephyrocapsa* species is more diffuse (Fig. 4B), perhaps because of the effects of sediment reworking by currents around/over the pelagic cap of this guyot (Pearson, this volume).

## **Discussion of Tropical and Equatorial Pacific Sites**

The combined results from Holes 871A, 872A, and 873B, show increases in mineral flux to the region of the Marshall Islands at  $\sim$ 2.3–1.5 Ma, at  $\sim$ 1.2–1.0 Ma, and possibly at  $\sim$ 0.4 Ma, respectively. We now compare the Marshall Island records with other sites in the north tropical and equatorial Pacific Ocean region, such as to the south, at Hole 803D on Ontong Java Plateau (Fig. 5; Krissek and Janecek, 1993), and in the eastern equatorial Pacific (ODP Leg 138; Sites 848, 849, 852, 853; Figs. 6–9; Hovan, 1995), which form a south-to-north latitudinal transect along 110°W to 109°W longitude.

The temporal resolution of data from any of these sites depends on the sample-spacing and the linear sedimentation rate at a particular site. Data from Hole 803D exhibit the lowest resolution for the interval in question (0–3 Ma), Hole 849B exhibits the highest resolution obtained. In this paper, only general comparisons will be made among these sites because of the complex nature of the proxy signal contained in the mineral flux records. Krissek and Janecek (1993) and Hovan (1995) provide detailed discussions and more extensive paleoenvironmental interpretations of these data.

The TMARs at sites located in the equatorial Pacific region are strongly influenced by the latitudinal position of the Intertropical Convergence Zone (ITCZ), which moves north and south in response to the changing strength of the northeast and southeast trade wind systems, both seasonally and over longer time scales. The ITCZ is characterized by light and variable winds and increased rainfall, with resulting wet-depositional scavenging of aerosol particles; the ITCZ thus forms an effective barrier to latitudinal eolian transport (for discussion, see Hovan, 1995). Therefore, the records of mineral dust fluxes at these equatorial sites may not entirely reflect Asian sources of mineral dust, but the cross-Pacific comparison may nonetheless be instructive. We will also examine the high-resolution eolian record from piston Core RC11-210 (Chuey et al., 1987; Rea et al., 1991), which was recovered from the central equatorial Pacific, between the eastern and western transects of sites discussed here.

Terrigenous mineral dust transported to Hole 803D on Ontong Java Plateau may originate from different sources and transport paths, including northeasterly and southeasterly trade winds, the Northern Hemisphere westerlies, or variable local winds (Krissek and Janecek, 1993). These authors have used mineral assemblage identifications to determine that Asian source areas beneath the Northern Hemisphere westerlies have exerted a stronger influence on the mineral assem-



Figure 2. A. Summary plots of sedimentologic data and mass accumulation rates vs. age (0–3 Ma) in Hole 871A (Limalok Guyot). The plots are of linear sedimentation rate (LSR), mass accumulation rate (MAR), carbonate mass accumulation rate (CAR), terrigenous mass accumulation rate (TMAR), and the weight percentage of calcium carbonate. Note the change to units of  $mg(cm^2 \times k.y.)^{-1}$  for TMAR, relative to the other plots. Also shown are plots of the ratio of species of the *Gephyrocapsa* group and of the non-*Gephyrocapsa* group, relative to the total nannoflora, respectively. **B.** Summary of quantitative nannofossil biostratigraphy vs. age (Ma) in Hole 871A. Data are from Erba (this volume). Note: *P. lacunosa* E (E = elliptical form).

blages brought to Ontong Java Plateau than Central or South American source areas beneath the easterlies.

From 0 to 3 Ma at Hole 803D, there are two broad intervals with higher mineral fluxes of approximately  $90-100 \text{ mg}(\text{cm}^2 \times \text{k.y.})^{-1}$ : one interval from >3.0 to ~2.7 Ma and one from ~2.4 to 1.6 Ma. These two intervals of higher mineral flux are followed by an apparent decline in the mineral flux uphole, before increasing again at about 0.5 Ma (Fig. 5). The general trends in mass accumulation observed at Hole 803D are qualitatively similar to those identified at Hole 873B, located 1000 km to the north. However, the sampling resolution of both of these records are generally poor and more data are needed to resolve shorter intervals.

Berger et al. (1994) note large fluctuations in sedimentation rates (primarily caused by changes in carbonate accumulation) at Hole 806B, located upslope of Hole 803D on Ontong Java Plateau, with significant increases in sediment accumulation at >2.0–1.8 Ma, 1.5– $\sim$ 1.3 Ma, ~1.2– $\sim$ 1.0 Ma, ~0.6–0.5 Ma, and 0.4– $\sim$ 0.2 Ma, respectively; these fluctuations suggest a climatic influence on carbonate accumulation on Ontong Java Plateau during the Quaternary (~413 k.y. periodicity over the past 2 m.y.). No mineral data are presently available from Hole 806B. Future improvements in the spatial and temporal resolution of mineral studies could greatly improve our ability to make cross-equatorial comparisons between sites on the Ontong Java Plateau (south of the equator) and sites on Marshall Island guyots (north of the equator) in the western Pacific.

Holes 848B, 849B, 852B, and 853B form a south-to-north transect across the equator in the eastern Pacific. Terrigenous MAR values increase northward from the southern holes (Holes 848B and 849B) to Hole 852B, but decrease at Hole 853B (see Figs. 6–9). The detailed patterns of flux variability at each of these holes are highly complex; this suggests that an increased sampling resolution is necessary to show the true variability in these records. At Hole 853B, there is a gradual, overall decrease in mineral flux from 3 to ~2.3 Ma, followed by a broad reversal in this trend (broad increase in mineral flux) from ~2.3 to ~0.6 Ma (Fig. 9). The mineral flux records from the sites to the south are more variable, with a more sawtooth pattern of variability being evident in the flux data (see Figs. 6–8).

Hovan (1995) notes that there are significant differences in the records of TMAR and grain size from north to south along this transect, and that there may have been a regional decrease in the supply of eolian material to the equatorial Pacific from eastern source regions (Central and South America) since the late Pliocene. The regional differences in records of eolian grain size and TMARs during the Quaternary may reflect a close relationship in the eastern equatorial Pacific between changes in atmospheric circulation strength, equatorial upwelling, and regional biological productivity (Murray and Leinen, 1993; Murray et al., 1993; Rea, 1994; Snoeckx and Rea, 1994; Farrell et al., 1995; Hovan, 1995; Pisias et al., 1995; among others).

The Quaternary sequence recovered in piston Core RC11-210, located just north of the equator in the central Pacific Ocean, provides high-quality, high-resolution, proxy records of atmospheric and oceanic variability (Chuey et al., 1987; Rea et al., 1991; for site location, see Table 1 and Fig.1). Core RC11-210 is the best studied core from the equatorial Pacific (Rea, 1994). Summary plots of the weight



Figure 2 (continued).

percentages of individual sedimentary components and corresponding mass accumulation rates are given in Figure 10A–B (data from Chuey et al., 1987; Rea et al., 1991). These plots illustrate the cyclical nature of these proxy records for the interval from 0 to 0.6 Ma, although there is a change in the nature of the variability of paleoclimatic proxy indicators at about 300 ka (Pisias and Rea, 1988; Rea, 1994). Changes in the median mineral grain size in this core have been interpreted as reflecting a northward latitudinal shift in the position of the ITCZ (of  $\sim 5^{\circ}$ ) after oxygen isotope stage 9 (at  $\sim 250-$ 300 ka). This shift in the position of the ITCZ results in a change in source for the terrigenous dust being transported to this site (change from Asian to South American source) and a change from a pattern of higher TMARs during glacial stages before the shift, to a pattern of higher fluxes of mineral dust during interglacials after the shift (see Chuey et al., 1987; Rea, 1994).

#### **Results from Northwest Pacific Sites**

#### Hole 880A

Hole 880A is located on Takuyo-Daisan Guyot (Table 1 and Fig. 1). Drilling at Hole 880A penetrated 18.4 mbsf. One lithologic unit consisting of four different, interlayered lithologies occurs at this site. In order of decreasing abundance, these are: volcaniclastic sand, foraminifer sand, nannofossil-foraminifer ooze, and volcanic ash (Premoli Silva, Haggerty, Rack, et al., 1993).

Recovered sediment ranges in color from shades of yellow-green and grayish brown to very dark gray. Increasing darkness typically is related to the increasing percentage of coarse-grained volcanic material, although light-colored volcanic ashes are also observed (see Rack, Lawyer, and Gee, this volume). Nonbiogenic components include feldspars, opaque minerals (possibly manganese nodules), vesicular lava fragments, volcanic lapilli, scoriaceous basalt grains, and pyrite (Premoli Silva, Haggerty, Rack, et al., 1993). Diatoms, sponge spicules, and silicoflagellates are frequent in the uppermost layers, but they decrease in abundance downhole (Premoli Silva, Haggerty, Rack, et al., 1993; Fenner, this volume). These hemipelagic sediments were deposited under the episodic influence of volcanic eruptions from the nearby convergent margin setting, which may have helped to preserve some of the more dissolution resistant biosiliceous microfossils (Fenner, this volume); continental runoff from Japan and/or China (as influenced by currents) also may have contributed to the amount of lithogenic material accumulating at this site. The recovered section at Hole 880A provides a condensed, but almost complete, Holocene to uppermost Pliocene sequence; a hiatus was detected in the upper Pleistocene where the *G. oceanica* Zone is extremely short (Erba, this volume).

# Sedimentologic and Biostratigraphic Results, Hole 880A

Eleven nannofossil biostratigraphic events and eight biozones were recognized at this site (see Table 2) and were calibrated to the magnetostratigraphy (Erba, this volume). These ages assigned to these events have been used with compositional data and dry bulk density values from Rack, Lawyer and Gee (this volume) to calculate mass accumulation rates for individual sedimentary components in Hole 880A (see Appendix A). Temporal fluctuations in MARs and in variations of the dominant nannofloral groups are shown in Figure 11A; Figure 11B shows changes in the relative abundance of individual nannofossil species (Erba, this volume). Near-continuous wholecore measurements of GRAPE bulk density and magnetic susceptibility are shown in Figure 12, along with plots of estimated dry bulk density, the weight percentages of coarse-grained and fine-grained



Figure 3. A. Summary plots of sedimentologic data and mass accumulation rates vs. age (0–3 Ma) in Hole 872A (Lo-En Guyot). NCAR = noncarbonate mass accumulation rate. B. Summary of quantitative nannofossil biostratigraphy vs. age (Ma) in Hole 872A. Data are from Erba (this volume). Note: *P. lacunosa* E (E = elliptical form).

mineral components (terrigenous and volcanic grains; Rack, et al., this volume), insoluble residues of the bulk sediment (Fenner, this volume), and the positions of ash and pumice layers in Hole 880A.

The highest MARs in Hole 880A are observed from ~1.7 to 1.5 Ma, and from 0.25 Ma to the top of the hole (Fig. 11A); the latter interval is associated with declines in the relative abundance of *G. oceanica* and small *Gephyrocapsa* species, and an increase in the abundance of *E. huxleyi* (Fig. 11B). In these two intervals, the total MAR values range from 2.0 g(cm<sup>2</sup> × k.y.)<sup>-1</sup> to 4.0 g(cm<sup>2</sup> × k.y.)<sup>-1</sup>, with the highest values occurring from ~1.7 to 1.5 Ma; fine-grained terrigenous mass accumulation rates (TMARs) range from 400 to 800 mg(cm<sup>2</sup> × k.y.)<sup>-1</sup> and ~300 mg(cm<sup>2</sup> × k.y.)<sup>-1</sup>, respectively, within these two intervals (Fig. 11A).

The high values of MARs and TMARs at ~1.7–1.5 Ma are associated with a decline in the relative abundance of *G. caribbeanica* and an increased dominance of the *P. lacunosa* E species of nannoflora (E = elliptical; 40%–60% of the total nannoflora in this interval); additional increases in the abundance of *P. lacunosa* E, from ~1.2 to 1.0 Ma and from ~0.8 to 0.5 Ma, are not associated with greatly increased MAR values. Sediment accumulation rates throughout the interval from ~1.5 to 0.4 Ma are approximately 1.0 g(cm<sup>2</sup> × k.y.)<sup>-1</sup>.

Figure 13 provides an expanded version of the data from Hole 880A for the interval from 0 to 0.6 Ma. Although Hole 880A is clearly a hemipelagic site, an understanding of the MARs of the various sedimentary components at this site near Japan may provide information about the relative intensity of western boundary ocean currents and the timing of arc-related volcanism in the general vicinity of Takuyo-Daisan Guyot.

#### Preliminary Identification of Ash Beds, Hole 880A

The close proximity of Hole 880A to sources of volcanism in and around Japan suggests that this hole might be used to develop an integrated nannofossil biostratigraphy and tephrochronology for this area of the Pacific Ocean. Hole 880A lies almost directly eastward of areas of extensive Pliocene–Pleistocene arc magmatism in Kyushu and southern to central Honshu (Kamata, 1989; Kamata and Kodama, 1994; Machida and Arai, 1983; Shimuzu and Itaya, 1993); therefore, the prevailing westerly winds would have transported material erupted from these volcanoes towards Hole 880A. Approximately 13 discrete ash layers have been identified in Hole 880A (Table 3; Premoli Silva, Haggerty, Rack, et al., 1993; Rack, Lawyer, and Gee, this volume); there may also be additional layers of disseminated ash and coarsegrained volcaniclastic particles which are not identified here.

Five widespread ashes of late Quaternary age have been identified on land and in deep-sea sediments around Japan according to Machida and Arai (1983). Three of these ashes are subalkalic tephras that have been correlated with large-scale eruptions that produced voluminous pyroclastic-flows and calderas in Kyushu, upwind of Hole 880A; these are the Kikai-Akahoya ash (K-Ah; C<sup>14</sup> age = 6.3 ka), Aira-Tn ash (AT; C<sup>14</sup> age = ~21–23 ka), and the Aso-4 ash (U-Th age = ~80 ka; Furuta et al., 1986; Machida and Arai, 1976, 1978, 1983, 1988).

An examination of the data contained in Rack, Lawyer, and Gee (this volume) and in Appendix A of this paper, suggests that at least two of these ashes are observed in Hole 880A. Based on observed increases in TMARs (and inferred quantities of coarse-grained material in the "unknown" fraction of the bulk sediment), we postulate that



Figure 3 (continued).

the AT ash lies at ~0.5–0.6 mbsf and that the Aso-4 ash lies at ~1.6–1.7 mbsf in Hole 880A. The presence of an ash layer located at approximately 50 cm from the sediment surface is confirmed by Fenner (this volume), who observed a significant increase in vesicular ash above background levels in Sample 144-880A-1H-1, 53 cm. At this time, measurements of the refractive index and mineralogy of these ashes have not been made.

Furuta et al. (1986) describe the AT ash as being fine-grained and well-sorted, with the glass fraction exceeding 90% of the entire material. The approximate age of this ash layer in Hole 880A has been estimated to be 24–28 ka, assuming constant sedimentation between the sediment surface (age = 0.0) and an age of 85 ka for the acme of *E. huxleyi* at 1.80 mbsf. This age estimate is considered to be reasonable for establishing a probable correlation of this layer with the AT ash.

The Aso-4 ash is recognized as having the longest-distance distribution of the five tephras studied by Furuta et al. (1986), occurring 1440 km away from its source (Aso Volcano in central Kyushu). Furuta et al. (1986) describe this tephra as having bubble-wall glass shards that are pale brown in color. They find this ash to be 2 cm thick, and located at 143 cm from the top of Core KH80-3-12, which is located near Hole 880A. Using the same dating criteria described above, the approximate age of the ash layer at 1.6–1.7 mbsf in Hole 880A is 76–80 ka. This age estimate is considered to be reasonable for establishing a probable correlation of this layer with the Aso-4 ash.

## Hole 810C

Hole 810C, located atop Shatsky Rise at intermediate water depths, was drilled during ODP Leg 132 (for location, see Table 1 and Fig. 1). Hole 810C penetrated 136.1 mbsf, mostly using the APC, with a total recovery of 143.8 m of nannofossil ooze. Two lithologic units

Table 3. Postulated ash or pumice layers in Hole 880A.

Depth (mbsf)	Dominant lithology	Thickness (cm)	Age (Ma)
0.50-0.53	Ash		0.026
1.60 - 1.70	Ash		0.078
2.11-2.16	Volcaniclastic sand	5	0.105
2.60-2.72	Volcaniclastic sand	12	0.136
3.50-3.54	Ash	4	0.185
3.70	Volcaniclastic sand	1	0.196
4.30	Volcaniclastic sand	1	0.231
5.4	Ash	1	0.484
6.40-6.45	Ash	5	0.599
6.85-6.95	Ash	10	0.654
7.97	Ash	1	0.776
8.95-9.80	Disseminated ash layers		
9.50	Volcanic lapilli		0.947
9.85	Volcanic lapilli		0.988
9.88-10.07	Ash	19	1.004
11.2 - 11.7	Disseminated ash layers		
11.3	Volcanic lapilli		1.194
11.5	Volcanic lapilli		1.230
12.5	Ash		1.420
12.9	Ash		1.490
14.40-14.45	Ash	5	1.592
15.65-15.70	Ash		1.647
16.47	Ash		2.035
17.30-17.60	Ash	30	2.740

were described in the upper 76 mbsf of this hole (Storms, Natland, et al., 1992). The upper unit consists of Pleistocene and younger, brown to dark gray nannofossil ooze with abundant siliceous microfossils (0–4.2 mbsf). There is evidence for cut and fill structures in the uppermost core, which suggest that significant erosional and/or sediment transport processes may have occurred in this upper unit. The lower unit consists of late Pliocene to Pleistocene light gray to white nannofossil ooze (4.2–76.0 mbsf). In the lower half of this lower unit,



Figure 4. A. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 873B (Wodejebato Guyot). B. Summary of quantitative nannofossil biostratigraphy vs. age (Ma) in Hole 873B. Data taken from Erba (this volume). Note: *P. lacunosa* E = elliptical form; C = circular form.

the light gray and white beds alternate at 40- to 80-cm intervals. Several thin ash beds and pumiceous pebbles are present in Hole 810C; these may be correlated to ash and pumice layers at other northwest Pacific sites (Natland, 1993).

#### Sedimentologic Results and Biostratigraphy, Hole 810C

The results from nearly 1000 determinations of the weight percentage of calcium carbonate in samples at approximately 5-cm sample spacing in Hole 810C have been presented by Rack et al. (1993); estimates of the noncarbonate weight percentage of the bulk samples were also made (100% - %carbonate = %noncarbonate). In this synthesis, the results from 400 chemical extractions of the weight percentage of fine-grained lithogenic material in bulk samples from Hole 810C are presented (see Appendix B and Fig. 14A).

The weight percentage of the "unknown" component in Hole 810C samples has been calculated using the "difference" method, where:  $100\% - (wt\% CaCO_3 + wt\%$  fine-grained lithogenic) = wt% unknown material. The "unknown" component includes biogenic opal, coarse-grained lithogenic particles, organic carbon, Mn and Fe oxides, hydroxides and zeolites, which have all been removed from the extracted samples by the chemical procedure (see methods). This estimate is used to provide a crude measure of the additional material fluxes to the seafloor at this site. The laboratory results, "difference" calculations, and calculated mass accumulation rates for these components are plotted vs. age in Figures 14A and 14B. Natland (1993) identified the stratigraphic position of numerous ash and pumice layers in Hole 810C; these are shown in Figure 14A. The relative contribution of volcanic material to the overall mineral flux (fine-grained and coarse-grained) is unknown at this time.

Clearly defined magnetic polarity zones provide a high-quality magnetic stratigraphy through the base of the Gilbert Chron, at approximately 76.8 mbsf in Hole 810C (Sager et al., 1993). Semi-quantitative biostratigraphic data were used to identify nannofossil zones based on the percentages of various species of the *Gephyrocapsa* group for the Pliocene to Holocene; foraminifer datums become progressively more important for age assignments with increasing sub-bottom depth in the section (see Table 2; Premoli Silva et al., 1993). Similar techniques were used at Holes 880A and 810C to identify the nannofossil boundaries (see Erba, this volume; Premoli Silva et al., 1993).

The results of laboratory analyses on samples from Hole 810C are used to calculate linear sedimentation rates and sediment mass accumulation rates for the various components (Appendix C and Fig. 14B). Hole 810C provides a high-resolution record of biogenic and lithogenic sediment accumulation from almost directly downwind of Hole 880A and Asian sources of eolian materials. The character of both the carbonate and fine-grained terrigenous weight percentage fluctuations in Hole 810C provide clear evidence of cyclicity in these records (Fig. 14A); higher frequency fluctuations for the late Pliocene change to longer period undulations for the Quaternary. Two intervals of very low sediment accumulation or nondeposition are clearly evident in the plot of linear sedimentation rate (Fig. 14B); the first is from ~1.9 to ~1.6 Ma, and the second is from ~1.25 to ~1.0 Ma. A large and seemingly abrupt increase in linear sedimentation rate and mass accumulation occurs at approximately 0.8 Ma in Hole 810C (Fig. 14B). This increase is followed by two to three, large-amplitude, undulations in the accumulation of both carbonate and mineral components in the younger record (Fig. 14B).

Although there are still significant fluctuations in the weight percentages of sedimentary components in the upper part of Hole 810C



Figure 4 (continued).

from 0.6 to ~0.3 Ma, the quality of the record above this level is poor (Fig. 15A); it is condensed and probably incomplete as indicated by the presence of cut-and-fill structures in Core 132-810C-1H (Storms, Natland, et al., 1992). MARs decrease abruptly at ~0.45 Ma, from values of  $3.0-4.0 \text{ g}(\text{cm}^2 \times \text{k.y.})^{-1}$  to values of  $1.0-1.5 \text{ g}(\text{cm}^2 \times \text{k.y.})^{-1}$  (Fig. 15B). Studies are presently underway to improve the age model and sampling resolution in the uppermost sections of this hole.

#### **Discussion of Northwest Pacific Sites**

The sedimentologic results from Hole 810C provide a highresolution record of biogenic and lithogenic sediment accumulation from a locality at intermediate water depths on southern Shatsky Rise. The record exhibits clear evidence of cyclicity in the character of both carbonate and eolian weight percentage fluctuations, with higher frequency fluctuations in the late Pliocene changing to longer period undulations in the Quaternary.

Piston Core V21-146 provides the best high-resolution record from northern Shatsky Rise for the past ~530 k.y. (Hovan et al., 1989, 1991). In the following paragraphs, these two records from Shatsky Rise and a third record from piston Core KK75-02, located just west of Hess Rise will be discussed as being the best marine colian records presently available from the northwest Pacific. With future adjustments to their individual age models, they provide the opportunity for improved long-term correlations between marine records of paleoclimate and the locss/soil records of continental paleoclimate from central China.

## Piston Core V21-146

One of the best examples of a continuous record of Quaternary climatic and oceanic conditions in the northwest Pacific over the past 530,000 years is the pelagic sediment record provided by Core V21-146 (Hovan et al., 1989, 1991; for location, see Table 1 and Fig. 1). Studies of this core have allowed direct comparisons to be made between the record of mineral dust transport to the Pacific Ocean, the SPECMAP oxygen isotope record, and the magnetic susceptibility record from continental loess deposits at Xifeng, China (Hovan et al., 1989, 1991; Kukla et al., 1990).

The primary data from Core V21-146 are listed in Table 2 of Hovan et al. (1991); however, in this manuscript, additional values have been calculated in order to estimate the weight percentage and mass accumulation rates of the noncarbonate fraction of the bulk sediment (100% - %carbonate = %noncarbonate), and the weight percentage of the "unknown" component of the bulk sediment ( $100\% - [wt\% CaCO_3 + wt\% Eolian (ash free) + \%ash] = \%unknown component). The "unknown" components are those materials that are dissolved during the extraction procedure and have not been otherwise measured (e.g., biogenic opal, sand-sized volcanic grains, organic carbon, Fe and Mn oxides, hydroxides, and zeolites; see methods); we expect that the weight percentage of biogenic opal, from diatom and radiolarian skeletons, makes up a significant percentage of the unknown fraction at this site.$ 

The summary plots of the weight percentages and mass accumulation rates of individual components (Fig. 16A–B) show that the combined volcanic ash and "unknown" components of the bulk sediment (20%–50% of the total) make a potentially large contribution to the noncarbonate accumulation at this site, in addition to the mineral dust flux. A better understanding of the composition of these unknown (undetermined) components of the bulk sediment may help to provide improved linkages between continental records of paleoclimate, atmospheric processes, and biological or chemical processes in the ocean.





A number of recent biological and chemical oceanographic studies have highlighted the potential role of iron in limiting phytoplankton primary production in the northern and central Pacific (DiTullio and Laws, 1991; Martin and Fitzwater, 1988). These studies suggest that temporal fluxes of particulate iron from continental sources may play an important role in phytoplankton bloom dynamics in the open ocean. Thus, an improved understanding of the export of mineral dust from Asian source regions and their accumulation in the sediments of the Pacific Basin may provide clues to understanding long-term carbon flux variations in the northern and central Pacific. This can only be accomplished by attempting to separate and quantify the relative contributions of volcanic and mineral dust sources of atmospheric input to the ocean (see Olivarez et al., 1991; Nakai et al., 1993; Rea, 1994), and also quantify the various contributions to the biogenous flux (measure the percentages of opal and organic carbon in the sediment, in addition to percentage of carbonate).

The stratigraphic positions of identified ash horizons in Core V21-146 may provide interesting comparisons with magnetic susceptibility records from this site (John King, URI, personal communication) and with the combined magnetic susceptibility records and identified ash layers at Hole 810C, DSDP Site 577, and possibly at Hole 880A. Rock magnetic measurements of sediments from Core V21-146 (northern Shatsky Rise), DSDP Site 577, and Hole 810C (southern Shatsky Rise; Doh, 1987; Polgreen et al., 1993), and Hole 880A (Takuyo-Daisan Guyot; Rack, Lawyer, and Gee, this volume), form a roughly east to west transect along a N-S trending arc towards Pliocene to Holocene sources of arc volcanism; these sites also extend from the northern boundary of the Pacific Subtropical Gyre toward the region of increased hemipelagic sediment processes along its western boundary (e.g., west of the Kuroshio Current; southeast of Japan). The combined use of sedimentological and geochemical tools combined with measurements of rock magnetic properties correlated to ash layers may provide a significantly improved stratigraphic framework for this region.

#### Piston Core KK75-02

Piston Core KK75-02 was collected from deep water to the east of the Hess Rise (Table 1 and Fig. 1). The predominant sedimentary components of Core KK75-02 are terrigenous clay, radiolarians, and diatoms, with subordinate amounts of quartz, sponge spicules, and manganese micronodules. The stratigraphic control for this core was provided by paleomagnetic reversal stratigraphy and radiolarian extinction levels (Janecek, 1983; Janecek and Rea, 1984, 1985; Rea et al., 1985).

A summary of the weight percentages and mass accumulation rates of principal sedimentary components (from the seafloor to 0.6 Ma) is shown in Figure 17 (data from Janecek, 1983). The weight percentage of noneolian material was determined by subtracting the weight percentage of eolian material from 100%; no further quantitative breakdown of sedimentary components is presently available for this core. As noted by Janecek and Rea (1984), there are several relative maxima in eolian accumulation, with reduced rates of accumulation in two intervals centered at ~0.55 and ~0.12 Ma, respectively (also see Rea, 1994).

#### Links Between Continental and Marine Eolian Records

A large amount of effort has been put into dating and understanding the climatic significance of the widespread loess and soil deposits of central China. These studies have focused on establishing



Figure 6. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 848B (eastern equatorial Pacific). Data are from Hovan (1995).

the magnetostratigraphy at the various loess localities (Luochuan, Lanzhou, Xifeng, Baoji, etc.) and on understanding the magnetic susceptibility signal, which varies significantly between loess and soil horizons in these sequences (Ding et al., 1993; Heller and Liu, 1982, 1984, 1986; Heller et al., 1991, 1994; Kukla, 1987; Kukla and An, 1989; Kukla et al., 1988, 1990; Liu et al., 1988a, 1988b; Rutter et al., 1990, 1991).

The fluctuations in magnetic susceptibility have been attributed to (1) changes in the atmospheric transport and accumulation of finegrained magnetite, (2) the degree of pedogenesis in different loess or soil horizons, and/or (3) other environmental causes (Evans and Heller, 1994; Eyre and Shaw, 1994; Kukla et al., 1988, 1990; Liu et al., 1992, 1993; Maher and Thompson, 1991) which are linked to fluctuations in aridity, rainfall, temperature and climate over the past 2.5 Ma (An et al., 1991; Ding et al., 1992; Liu et al., 1989; among many others).

Attempts to correlate the magnetic susceptibility record of the loess sequences with the marine oxygen isotope record have met with varying degrees of success, with best results coming from sequences of Brunhes age (Heller and Liu, 1982, 1986; Hovan et al., 1989, 1991; Kukla et al., 1990). At present, the marine eolian record from Core V21-146 (Fig. 16B) is the only one that has been used to provide a direct link between the China loess sequences and mineral fluxes to the open Pacific Ocean.

Figure 18 provides a visual comparison of the weight percentages of mineral components and the mineral accumulation rates from the three northwest Pacific eolian sites discussed in this study (Hole 810C, Core V21-146, and Core KK75-02). Also shown is the magnetic susceptibility record of loess and soil horizons at the Xifeng I locality in Central China (data from Kukla et al., 1990), where high susceptibility values correspond to soil horizons and low susceptibility values correspond to loess beds. The point of this visual comparison is to suggest that improved correlations between marine eolian records and the Chinese loess sequences are possible for extended stratigraphic intervals, with further improvements to the existing age models and correlations among marine sites with available mineral data. The long high-resolution record from Hole 810C, at present, provides the best opportunity to accomplish these types of land-marine correlations for the older (>0.5 Ma) portions of the loess record (Rack and Janecek, unpubl. data).

#### SUMMARY AND CONCLUSIONS

This paper has presented the integrated results of quantitative sedimentologic and biostratigraphic studies at Sites 871, 872, and 873 in the Marshall Islands, in Hole 880A on Takuyo-Daisan Guyot, and in Hole 810C on Shatsky Rise for the late Pliocene to Quaternary. Mass accumulation rates of different sedimentary components have been calculated at each site to compare biogenic and mineral fluxes across a wide geographic region at intermediate and deeper water depths.

Published results from mineral dust studies at additional ODP, DSDP, and core sites located within or near the boundaries of the Pacific Subtropical Gyre have been used to provide regional comparisons among several high-quality eolian records. Additional discussions have addressed the present state of research aimed at understanding the linkages between continental records of paleoclimate from central China and marine eolian records located downstream, beneath the Pacific Ocean.

Based on observed increases in the weight percentage and accumulation rate of lithogenic sediments (terrigenous and volcanic grains), and age estimates from nannofossil stratigraphy, we have tentatively identified the presence of two widespread marker tephra of late Quaternary age in Hole 880A. These two ashes, which have been identified



Figure 7. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 849B (eastern equatorial Pacific). Data are from Hovan (1995).

on land and in deep-sea sediments around Japan (Machida and Arai, 1983), are identified in Hole 880A as the Aira-Tn (AT) ash at ~0.5-0.6 mbsf, and the Aso-4 ash at ~1.6-1.7 mbsf.

The sedimentologic results from Hole 810C provide a high-resolution record of biogenic and lithogenic sediment accumulation from a locality at intermediate water depth on Shatsky Rise. The record exhibits clear evidence of cyclicity in the character of both carbonate and eolian weight percentage fluctuations, with higher frequency fluctuations for the late Pliocene changing to longer period undulations for the Quaternary.

The combined use of quantitative nannofossil stratigraphic studies, with sedimentologic studies, and the systematic identification of ash (tephra) and pumice layers at sites in the northwest Pacific should improve our understanding of the timing and significance of paleoclimatic and paleoceanographic events in this region. Additional emphasis on oxygen isotope determinations and the use of orbital tuning strategies at these sites will help to improve the stratigraphic resolution.

Improved correlations between marine eolian records, such as the records at Hole 810C and Core V21-146, and the classic sections of Chinese loess, should make possible higher resolution comparisons than were previously available, especially for older stratigraphic intervals, with further improvements to the age models and through the use of astronomic tuning of the records from these sites.

Fluctuations in mass accumulation rates and terrigenous mineral fluxes seem to be generally coherent between sites in the Marshall Islands, but the low resolution of the data and the presence of current-influenced reworking of sediments in some intervals prevent detailed correlations between sites. Significant increases in both MARs and TMARs are observed for the intervals from >1.7 to 1.5 Ma and from 1.2 to 1.0 Ma in the records from Limalok and Lo-En guyots (Holes

871A and 872A), and possibly at ~0.4 Ma at Hole 871A. Hole 873B shows a broad increase in TMAR from ~3.0 to ~1.2 Ma, with erosion caused by possible current activity around the summit of Wodejebato Guyot limiting the accumulation of younger sediments. Distinctive fluctuations in nannofloral assemblages are associated with temporal changes in both carbonate and mineral fluxes; these comparisons may provide a linkage between long-term proxy records of paleoclimatic conditions in the western tropical North Pacific.

Quantitative measures of nannofloral abundance fluctuations, both within specific nannofossil groups and among individual species, can be useful for constraining ecologic and oceanographic interpretations for these ODP sites. Major changes in the composition of Pleistocene nannofloras appear to be coeval on a regional basis, across a wide latitudinal range within the subtropical to equatorial belts (Raffi and Flores, 1995; Erba, this volume). However, the nannofossil records from Sites 871, 872, 873, and 880 indicate that some taxa were locally and temporarily influenced by the specific oceanographic conditions around these guyots. Future investigations of the distribution of specific taxa, such as the smallest specimens of *Gephyrocapsa* (2–2.5  $\mu$ m) and *Florisphaera profunda*, may help in separating global from local or regional effects.

The lack of adequate late Pliocene to Quaternary mineral flux records from Ontong Java Plateau prevents, at present, the effective development of a north-south latitudinal transect of sites across the equator in the western Pacific. The timing of observed fluctuations in sedimentation rates at Hole 806B on Ontong Java Plateau (Berger et al., 1994) suggests that comparisons between the Ontong Java sites and data from the pelagic caps of Leg 144 guyots could lead to an improved understanding of intermediate-water-depth ocean circulation patterns and the geochemical variability of paleowater masses in the western equatorial to tropical Pacific Ocean. Improvements in the sampling resolution of eastern equatorial Pacific mineral dust records, their combination with measurements of the flux of other sedimentary components (e.g., %carbonate, %opal, %organic carbon), and the use of near-continuous proxy records (GRAPE bulk density and magnetic susceptibility data) from these cores may significantly improve our understanding of past changes in atmospheric and oceanic conditions in the equatorial Pacific (see Farrell et al., 1995; Pisias et al., 1995). Additional studies on transects of piston cores along 110°W longitude (Snoeckx and Rea, 1994) and 135°W longitude (Murray et al., 1993) should improve the spatial resolution of paleoclimatic and paleoceanographic proxy records of variability across the equatorial Pacific, and will allow for more rigorous comparisons with the existing records from Core RC11-210.

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<sup>\*</sup> Abbreviations for names of organizations and publications in ODP reference lists follow the style given in *Chemical Abstracts Service Source Index* (published by American Chemical Society).

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Figure 8. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 852B (eastern equatorial Pacific). Data are from Hovan (1995).



Figure 9. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 853B (eastern equatorial Pacific). Data are from Hovan (1995).



Figure 10. A. Summary plots of sedimentologic data vs. age (0-0.6 Ma) for piston Core RC11-210 (central equatorial Pacific). Data are from Chuey et al. (1987) and Rea et al., (1991). B. Summary plots of mass accumulation rates vs. age (0-0.6 Ma) for piston Core RC11-210. Above about 0.3 Ma, there is an inferred change in the source of terrigenous material, from Asian sources below to South American sources above (Chuey et al., 1987; Rea et al., 1991). The percentage of "unknown" corresponds to the weight percentage of noncarbonate material that has not been identified in the total percentage (100% - wt% [carbonate + opal + organic carbon + terrigenous components]).



Figure 10 (continued).



Figure 11. A. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-3 Ma) in Hole 880A (Takuyo-Daisan Guyot). Note that TMAR is given in units of mg $(\text{cm}^2 \text{ k.y.})^{-1}$ . B. Summary plots of quantitative nannofossil biostratigraphy vs. age (Ma) in Hole 880A. Data are from Erba (this volume).



Figure 11 (continued).



Figure 12. Summary plots of physical properties, sedimentologic data, and low-field magnetic susceptibility vs. age (Ma) in Hole 880A. The percentage of insoluble residue in bulk samples is obtained from data in Fenner (this volume). The percentage of "unknown" material equals the percentage of noncarbonate material minus the fine-grained lithogenic material. Much of the "unknown" material is suspected to be sand-sized volcanic or hemipelagic grains (see Rack, Lawyer, and Gee, this volume).



Figure 13. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-0.6 Ma) in Hole 880A. Data are from Rack, Lawyer, and Gee (this volume).



Figure 14. **A.** Summary plots of GRAPE bulk density (GBD), sedimentologic data, the positions of ash and pumice layers (Natland, 1993), and low-field magnetic susceptibility at Hole 810C on Shatsky Rise. The "unknown" sediment component is suspected to be largely composed of biogenic opal, coarse-grained lithogenic grains, oxides and hydroxides, and organic carbon. **B.** Summary plots of linear sedimentation rates, dry bulk density, and mass accumulation rates vs. age (0–3 Ma) in Hole 810C.



Figure 14 (continued).



Figure 15. A. Summary plots of GRAPE bulk density (GBD), sedimentologic data, and low-field magnetic susceptibility vs. age (0–0.6 Ma) in Hole 810C. Data are from Rack et al. (1993) and Rack, Lawyer, and Gee (this volume). The magnetic susceptibility record from the Xifeng I site in Central China for the same time interval is also shown; susceptibility data are from Kukla et al. (1990). **B.** Summary plots of mass accumulation rates vs. age (0–0.6 Ma) in Hole 810C.



Figure 15 (continued).



Figure 16. A. Summary plots of sedimentologic data, median colian grain size (in phi units), and oxygen isotope stratigraphy vs. age (0–0.6 Ma) for piston Core V21-146 (northwest Pacific). Data are from Hovan et al. (1991). The percentage of "unknown" material is suspected to be largely composed of biogenic opal, coarse-grained lithogenic grains, oxides, hydroxides, zeolites, and organic carbon. **B.** Summary plots of mass accumulation rates vs. age (0–0.6 Ma) for piston Core V21-146.





721



Figure 17. Summary plots of sedimentologic data and mass accumulation rates vs. age (0-0.6 Ma) for piston Core KK75-02. Data are from Janecek (1983). The weight percentage of noneolian material equals the difference of 100% - wt% eolian material.



Figure 18. Comparisons between selected Pacific "eolian" records and the magnetic susceptibility record from Xifeng I, in Central China. Magnetic susceptibility data from the loess record at Xifeng I are from Kukla et al. (1990). The marine data from piston Core V21-146 represent the only record that has been correlated to the continental record (Hovan et al., 1989, 1991). The record from Hole 810C is poorly resolved over the upper 0.2 m.y., but it provides the opportunity to correlate older intervals (>0.5 Ma; see Fig. 14) with the continental loess records from Central China (Rack and Janecek, unpubl. data).

# APPENDIX A

Linear Sedimentation and Mass Accumulation Rates for Samples from Hole 880A (Takuyo-Daisan Guyot)

Core, section,	Depth	Age	LSR	GBD.3	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )	(g/cm <sup>3</sup> )		g	$(cm^2 \times k.y.$	.)~1		$mg(cm^2 \times k.y.)^-$
44-880A-	0.20	0.000	2.119	1.60	1.01	214	0.74	1.40	1.12	0.076	276
1H-1, 20-22 1H-1, 30-32	0.20	0.009	2.118	1.73	1.01	2.14	0.74	1.40	1.12	0.276	320
1H-1, 40-42	0.40	0.019	2.118	1.73	1.10	2.33	0.89	1.44	1.20	0.239	239
1H-1, 50-52	0.50	0.024	2.118	1.76	1.16	2.46	0.65	1.80	1.45	0.353	353
1H-1, 59-61	0.59	0.028	2.118	1.80	1.25	2.65	0.79	1.86	1.50	0.356	356
1H-1, 70–72	0.70	0.033	2.118	1.79	1.22	2.58	0.83	1.75	1.38	0.374	374
1H-1, 81-85 1H-1 00_02	0.81	0.038	2.118	1.90	1.53	3.24	0.36	2.88	2.44	0.446	446
1H-1, 100-102	1.00	0.043	2 118	1 74	1.40	2.97	0.34	2.02	1.89	0.227	227
1H-1, 110-112	1.10	0.052	2.118	1.74	1.11	2.35	0.19	2.16	1.89	0.265	265
1H-1, 120-122	1.20	0.057	2.118	1.78	1.21	2.56	0.17	2.39	2.16	0.234	234
1H-1, 130-132	1.30	0.061	2.118	1.80	1.24	2.63	0.72	1.90	1.64	0.266	266
1H-1, 140-142	1.40	0.066	2.118	1.80	1.23	2.61	0.94	1.66	1.52	0.147	147
1H-2, 1-3	1.51	0.071	2.118	1.52	0.78	1.65	0.73	0.92	0.72	0.202	202
1H-2, 10-12 1H-2, 20-22	1.00	0.070	2.118	1.74	1.10	2.55	0.78	1.55	1.20	0.330	374
1H-2, 30-32	1.80	0.085	1.714	1.76	1.16	1.99	0.95	1.04	0.83	0.209	209
1H-2, 38-40	1.88	0.090	1.714	1.77	1.17	2.01	0.68	1.33	1.08	0.253	253
1H-2, 50-52	2.00	0.097	1.714	1.72	1.07	1.83	0.39	1.44	1.15	0.290	290
1H-2, 61-63	2.11	0.103	1.714	1.80	1.26	2.16	0.47	1.69	1.34	0.351	351
1H-2, 68-70	2.18	0.107	1.714	1.81	1.27	2.18	0.53	1.65	1.42	0.227	227
1H-2, 80-82 1H-2, 90-92	2.50	0.114	1.714	1.88	1.40	2.50	0.20	2.30	2.05	0.248	248
1H-2, 100-102	2.50	0.120	1.714	1.86	1.40	2.54	0.09	2.45	2.04	0.246	246
1H-2, 110-112	2.60	0.132	1.714	1.77	1.18	2.02	0.07	1.95	1.69	0.264	264
1H-2, 119-121	2.69	0.137	1.714	1.68	0.99	1.70	0.00	1.70	1.51	0.188	188
1H-2, 130-132	2.80	0.143	1.714	1.81	1.26	2.16	0.29	1.87	1.57	0.300	300
1H-2, 140–142	2.90	0.149	1.714	1.80	1.25	2.14	0.10	2.04	1.75	0.289	289
H-3, 1-3	3.10	0.150	1./14	1.72	1.07	1.85	0.54	1.50	1.29	0.211	211
1H-3, 20-22	3.20	0.167	1.714	1.75	1.08	2.28	0.62	1.24	1.05	0.207	201
IH-3, 30-32	3.30	0.172	1.714	1.85	1.37	2.35	0.65	1.70	1.42	0.277	277
H-3, 40-42	3.40	0.178	1.714	1.79	1.23	2.11	0.24	1.86	1.56	0.305	305
1H-3, 51–53	3.51	0.185	1.714	1.85	1.36	2.33	0.16	2.17	1.56	0.612	612
H-3, 59-61	3.59	0.189	1.714	1.98	1.78	3.05	0.60	2.45	2.14	0.318	318
IH-3, /1-/3	3.71	0.196	1.714	1.80	1.25	2.14	0.17	1.98	1.75	0.222	222
1H-3, 76-60 1H-3, 90-92	3.90	0.200	1.714	1.00	1.47	2.52	0.18	1.77	1 49	0.245	245
1H-3, 100-102	4.00	0.213	1.714	1.81	1.27	2.18	0.67	1.50	1.28	0.228	228
1H-3, 110-112	4.10	0.219	1.714	1.89	1.50	2.57	0.47	2.10	1.79	0.310	310
1H-3, 118-120	4.18	0.224	1.714	1.89	1.49	2.55	0.39	2.17	1.76	0.404	404
IH-3, 130–132	4.30	0.231	1.714	1.85	1.39	2.38	0.21	2.17	1.90	0.269	269
1H-3, 140–142	4.40	0.237	1.714	1.84	1.35	2.31	0.17	2.15	1.87	0.273	273
1H-4, 0-2 1H-4, 10-12	4.50	0.245	1.714	1.74	1.11	1.90	0.38	1.55	1.33	0.195	195
1H-4, 20-22	4.70	0.254	1.714	1.79	1.23	2.11	0.65	1.46	1.26	0.192	192
1H-4, 30-32	4.80	0.260	0.199	1.79	1.22	0.24	0.08	0.16	0.14	0.028	28
1H-4, 40-42	4.90	0.310	0.199	1.76	1.15	0.23	0.07	0.15	0.13	0.022	22
1H-4, 48–50	4.98	0.350	0.199	1.82	1.29	0.26	0.09	0.17	0.13	0.039	39
1H-4, 60-62	5.10	0.411	0.199	1.84	1.34	0.27	0.04	0.23	0.18	0.047	4/
1H-4, 80-82	5.30	0.401	0.882	1 79	1.10	1.04	0.18	0.80	0.72	0.138	181
1H-4, 90-92	5.40	0.484	0.882	1.74	1.11	0.98	0.11	0.87	0.76	0.117	117
1H-4, 98-100	5.48	0.493	0.882	1.74	1.09	0.96	0.27	0.69	0.57	0.114	114
1H-4, 110–112	5.60	0.506	0.882	1.79	1.22	1.08	0.34	0.74	0.61	0.128	128
IH-4, 120–122	5.70	0.518	0.882	1.77	1.18	1.04	0.16	0.88	0.70	0.175	175
H-4, 130–152	5.80	0.529	0.882	1.78	1.19	1.05	0.24	0.81	0.63	0.180	180
IH-5, 1-3	6.01	0.553	0.882	1.09	1.01	0.89	0.51	0.58	0.49	0.192	192
IH-5, 10-12	6.10	0.563	0.882	1.83	1.32	1.16	0.12	1.04	0.70	0.336	336
IH-5, 19-21	6.19	0.573	0.882	1.78	1.20	1.06	0.16	0.90	0.69	0.216	216
H-5, 30–32	6.30	0.586	0.882	1.71	1.04	0.92	0.09	0.83	0.55	0.277	277
H-5, 39-41	6.39	0.596	0.882	1.79	1.21	1.07	0.06	1.01	0.75	0.261	261
H-5, 50-52	6.50	0.608	0.882	1.85	1.33	1.17	0.22	0.96	0.80	0.160	160
H-5, 70-72	6.70	0.631	0.882	1.80	1.30	1.15	0.10	0.99	0.82	0.176	176
H-5, 80-82	6.80	0.642	0.882	1.85	1.38	1.22	0.56	0.65	0.54	0.114	114
H-5, 90-92	6.90	0.654	0.882	1.85	1.38	1.22	0.00	1.22	1.02	0.196	196
H-5, 98-100	6.98	0.663	0.882	1.74	1.10	0.97	0.47	0.50	0.37	0.131	131
H-5, 110-112	7.10	0.677	0.882	1.78	1.19	1.05	0.46	0.59	0.44	0.144	144
H-5, 120-122	7.20	0.688	0.882	1.76	1.15	1.01	0.43	0.58	0.43	0.152	152
H-5, 130-132	7.30	0.099	0.882	1.75	1.12	0.99	0.42	0.57	0.42	0.149	149
H-6, 0-2	7.50	0.722	0.882	1.78	1.20	1.00	0.29	0.77	0.30	0.208	173
1H-6, 10-12	7.60	0.733	0.882	1.83	1.31	1.16	0.44	0.72	0.52	0.195	195
1H-6, 20-22	7.70	0.744	0.882	1.77	1.17	1.03	0.26	0.77	0.58	0.192	192
1H-6, 30–32	7.80	0.756	0.882	1.80	1.26	1.11	0.36	0.75	0.60	0.154	154
IH-6, 40-42	7.90	0.767	0.882	1.78	1.20	1.06	0.00	1.06	0.58	0.479	479
1H-6, 48-50	7.98	0.776	0.882	1.86	1.40	1.23	0.29	0.94	0.74	0.200	200
1H-6, 60-62	8.10	0.790	0.882	1.76	1.16	1.02	0.09	0.94	0.71	0.231	231
1H-6 80_82	8.20	0.801	0.882	1.83	1.32	1.10	0.07	0.82	0.93	0.100	100
	0.00	0.010	0.002	1.0.5	1.34	1.10	0.55	0.02	0.09	0.147	147

						MAD	CAR	NCAD	OMAD	TMAD	TMAD2
Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	LSR (cm/k.v.)	GBD.3 (g/cm <sup>3</sup> )	DBD (g/cm <sup>3</sup> )	MAK	CAR	$(cm^2 \times k.y.$	) <sup>-1</sup>	TWAK	$mg(cm^2 \times k.y.)^{-1}$
111.6 00 02	8.40	0.824	0.882	1.93	1 22	1.17	0.33	0.85	0.69	0.154	154
1H-6, 100–102	8.50	0.835	0.882	1.77	1.18	1.04	0.24	0.85	0.57	0.240	240
1H-6, 110–112 2H-1 5–7	8.60	0.847	0.882	1.61	0.89	0.78	0.20	0.58	0.42	0.167	167
2H-1, 3-7 2H-1, 11-13	9.01	0.893	0.882	1.70	1.03	0.91	0.10	0.78	0.52	0.205	205
2H-1, 20-22	9.10	0.903	0.882	1.68	0.99	0.87	0.34	0.54	0.36	0.175	175
2H-1, 29-31 2H-1, 40-42	9.19	0.914	0.882	1.62	0.89	0.78	0.36	0.42	0.26	0.167	231
2H-1, 50-52	9.40	0.937	0.882	1.72	1.07	0.94	0.33	0.61	0.40	0.213	213
2H-1, 59-61	9.49	0.947	0.882	1.71	1.06	0.93	0.25	0.68	0.45	0.230	230
2H-1, 80-82.	9.70	0.901	0.882	1.77	1.18	1.04	0.25	0.33	0.58	0.287	287
2H-1, 90-92	9.80	0.983	0.882	1.63	0.91	0.80	0.12	0.68	0.53	0.147	147
2H-1, 100–102 2H-1, 110–112	9.90	1.005	0.882	1.72	1.07	0.94	0.27	0.67	0.49	0.176	1/6
2H-1, 120-122	10.10	1.017	0.882	1.78	1.21	1.07	0.15	0.92	0.78	0.140	140
2H-1, 130–132 2H-2, 0–2	10.20	1.028	0.656	1.73	1.08	0.71	0.09	0.61	0.52	0.092	92
2H-2, 10-12	10.37	1.054	0.656	1.81	1.27	0.83	0.13	0.71	0.60	0.110	110
2H-2, 20-22	10.47	1.069	0.656	1.80	1.26	0.83	0.31	0.51	0.40	0.111	111
2H-2, 27-29 2H-2, 40-42	10.54	1.100	0.656	1.83	1.31	0.86	0.27	0.59	0.47	0.177	177
2H-2, 50-52	10.77	1.115	0.656	1.76	1.15	0.75	0.29	0.46	0.34	0.118	118
2H-2, 57-59	10.84	1.126	0.656	1.75	1.13	0.74	0.19	0.55	0.41	0.134	134
2H-2, 77-79	11.04	1.145	0.656	1.77	1.25	0.77	0.10	0.63	0.49	0.128	145
2H-2, 90-92	11.17	1.176	0.656	1.85	1.38	0.91	0.24	0.67	0.55	0.120	120
2H-2, 100–102 2H-2, 110–112	11.27	1.191	0.656	1.86	1.40	0.92	0.17	0.75	0.58	0.171	171
2H-2, 120-122	11.47	1.222	0.656	1.72	1.08	0.71	0.11	0.60	0.49	0.102	102
2H-2, 132-134	11.59	1.240	0.515	1.73	1.08	0.56	0.15	0.41	0.32	0.082	82
2H-2, 140–142 2H-3, 1–3	11.67	1.255	0.515	1.75	0.83	0.57	0.14	0.43	0.36	0.067	61
2H-3, 10-12	11.87	1.294	0.515	1.81	1.27	0.65	0.09	0.57	0.49	0.075	75
2H-3, 20-22 2H-3, 27-20	11.97	1.314	0.515	1.76	1.15	0.59	0.10	0.49	0.43	0.064	64 50
2H-3, 40-42	12.04	1.353	0.515	1.80	1.24	0.64	0.08	0.56	0.48	0.075	75
2H-3, 50-52	12.27	1.372	0.515	1.75	1.12	0.58	0.09	0.48	0.44	0.048	48
2H-3, 58-60 2H-3, 70-72	12.35	1.388	0.515	1.75	1.09	0.56	0.09	0.48	0.59	0.084	60
2H-3, 80-82	12.57	1.430	0.515	1.77	1.17	0.60	0.17	0.43	0.36	0.077	77
2H-3, 90-92 2H-3, 97-99	12.67	1.450	0.515	1.82	1.29	0.66	0.16	0.50	0.44	0.062	62
2H-3, 110-112	12.87	1.489	0.515	1.82	1.29	0.66	0.06	0.61	0.48	0.130	130
2H-3, 120-122	12.97	1.508	0.515	1.77	1.17	0.60	0.19	0.41	0.37	0.038	38
2H-3, 130–132 2H-3, 140–142	13.07	1.528	2.800	1.88	1.47	3.33	0.15	2.96	2.38	0.572	572
2H-4, 0-2	13.27	1.551	2.800	1.79	1.22	3.42	0.50	2.92	2.34	0.582	582
2H-4, 10-12 2H-4, 20-22	13.37	1.554	2.800	1.84	1.34	3.75	0.92	2.83	2.34	0.491	491
2H-4, 30-32	13.57	1.561	2.800	1.83	1.33	3.72	0.79	2.94	2.24	0.702	702
2H-4, 40-42	13.67	1.565	2.800	1.82	1.31	3.67	1.42	2.25	1.74	0.513	513
2H-4, 57–59	13.84	1.508	2.800	1.78	1.19	3.33	0.33	3.00	2.66	0.328	341
2H-4, 72-74	13.99	1.576	2.800	1.80	1.25	3.50	0.74	2.76	2.39	0.375	375
2H-4, 80–82 2H-4, 90–92	14.07	1.579	2.800	1.74	1.12	3.14	0.55	2.59	2.21	0.375	535
2H-4, 100-102	14.27	1.586	2.800	1.73	1.08	3.02	0.15	2.88	2.25	0.632	632
2H-4, 110–112 2H-4, 120–122	14.37	1.590	2.800	1.80	1.25	3.50	1.14	2.36	1.76	0.597	597
2H-4, 130–132	14.57	1.597	2.222	1.77	1.17	2.60	0.80	1.80	1.50	0.304	304
2H-4, 140-142	14.67	1.602	2.222	1.81	1.26	2.80	0.76	2.04	1.65	0.386	386
2H-5, 10-12	14.77	1.610	2.222	1.74	1.39	3.09	0.30	2.10	2.16	0.639	639
2H-5, 20-22	14.97	1.615	2.222	1.80	1.25	2.78	0.77	2.01	1.56	0.448	448
2H-5, 30-32 2H-5, 40-42	15.07	1.620	2.222	1.86	1.41	3.13	0.58	2.55	2.18	0.368	368
2H-5, 50-52	15.27	1.628	2.222	1.83	1.34	2.98	0.38	2.59	1.97	0.620	620
2H-5, 62-64	15.39	1.634	2.222	1.86	1.42	3.16	0.39	2.77	2.37	0.402	402
2H-5, 72-74 2H-5, 80-82	15.49	1.638	2.222	1.87	1.44	3.20	0.72	2.48	2.09	0.398	398
2H-5, 88-90	15.65	1.646	2.222	1.79	1.23	2.73	0.04	2.69	1.64	1.047	1047
2H-5, 100-102	15.77	1.651	2.222	1.87	1.43	3.18	0.71	2.46	2.08	0.389	389
2H-5, 120–122	15.97	1.660	0.133	1.84	1.35	0.18	0.05	0.13	0.11	0.022	22
2H-5, 129-131	16.06	1.727	0.133	1.85	1.39	0.18	0.05	0.14	0.10	0.038	38
2H-5, 136–138 2H-6, 0–2	16.13	1.780	0.133	1.72	1.10	0.14	0.04	0.10	0.07	0.023	23
2H-6, 10-12	16.37	1.960	0.133	1.83	1.31	0.17	0.05	0.13	0.11	0.019	19
2H-6, 20-22	16.47	2.035	0.133	1.70	1.03	0.14	0.03	0.11	0.08	0.027	27
2H-6, 40-42	16.67	2.185	0.133	1.74	1.08	0.15	0.03	0.12	0.09	0.032	23
2H-6, 50-52	16.77	2.260	0.133	1.80	1.25	0.17	0.05	0.11	0.09	0.025	25
2H-6, 60-62 2H-6, 69-71	16.87	2.335	0.133	1.81	1.26	0.17	0.04	0.13	0.11	0.024	24
2H-6, 81-83	17.08	2.493	0.133	1.77	1.16	0.15	0.02	0.13	0.11	0.024	24
2H-6, 90-92	17.17	2.560	0.133	1.77	1.16	0.15	0.03	0.13	0.11	0.019	19
2H-6, 97-99 2H-6, 110-112	17.24	2.012	0.133	1.80	1.25	0.17	0.03	0.14	0.12	0.019	21

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	Core, section.	Depth	Age	LSR	GBD 3	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>	
interval (cm) (mbsf) (Ma) (cm/k.y				(cm/k.y.)	(g/cm <sup>3</sup> )	(g/cm <sup>3</sup> )	$g(cm^2 \times k.y.)^{-1}$					$mg(cm^2 \times k.y.)^{-1}$	
	2H-6, 120-122	17.47	2.785	0.133	1.76	1.15	0.15	0.03	0.13	0.10	0.029	29	
	2H-6, 130-132	17.57	2.860	0.133	1.72	1.07	0.14	0.03	0.12	0.08	0.038	38	
	2H-6, 139-141	17.66	2.928	0.133	1.74	1.12	0.15	0.04	0.11	0.10	0.015	15	
	2H-7, 0-2	17.77	3.010	0.133	1.76	1.15	0.15	0.03	0.12	0.10	0.017	17	
	2H-7, 10-12	17.87	3.085	0.133	1.71	1.05	0.14	0.03	0.11	0.09	0.021	21	
	2H-7, 20-22	17.97	3.160	0.133	1.71	1.05	0.14	0.04	0.10	0.08	0.025	25	
	2H-7, 30-32	18.07	3.235	0.133	1.78	1.21	0.16	0.02	0.14	0.11	0.031	31	
	2H-7, 40-42	18.17	3.310	0.133	1.80	1.25	0.17	0.02	0.15	0.12	0.023	23	
	2H-7, 52-54	18.29	3.400	0.133	1.77	1.18	0.16	0.04	0.12	0.09	0.027	27	

Notes: Depth (mbsf) = sub-bottom depth in meters below seafloor; age (Ma) = millions of years ago; LSR = linear sedimentation rate (cm per thousand years); GBD.3 = GRAPE bulk density (3 point smoothed, g/cm<sup>3</sup>); DBD = dry bulk density (g/cm<sup>3</sup>); MAR = mass (bulk) accumulation rate (grams of sediment per square cm per thousand years; same units for other rates); CAR = carbonate accumulation rate; NCAR = noncarbonate accumulation rate; OMAR = noncarbonate minus eolian accumulation rate; TMAR<sup>1</sup> = terrigenous accumulation rate (lithogenic material <63  $\mu$ m) in units of g(cm<sup>2</sup> × k.y.)<sup>-1</sup>; TMAR<sup>2</sup> = terrigenous accumulation rate in units of mg(cm<sup>2</sup> × k.y.)<sup>-1</sup>. The carbonate and noncarbonate weight percentage data used to calculate these accumulation rates are contained in Rack et al. (Data Report, this volume).

# APPENDIX B

# Results of Chemical Extractions of Lithogenic Material in Samples from Hole 810C (Shatsky Rise)

Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	Fine-grained lithogenic (wt%)	Noncarb. – fine-grained lithogenic (wt%)	Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	Fine-grained lithogenic (wt%)	Noncarb. – fine-grained lithogenic (wt%)
132-810C-					2H-3, 134.6	6.75	0.490	9.66	11.97
1H-1, 11.5	0.12	0.025	13.40	15.26	2H-3, 140.7	6.81	0.491	9.48	13.38
1H-1, 17.5 1H-1, 50.5	0.17	0.035	23.58	21.09	2H-3, 140.5 2H-4, 59.5	7.50	0.493	8.71	30.19
1H-1, 80.5	0.81	0.117	22.46	18.82	2H-4, 74.5	7.64	0.513	6.61	10.02
1H-1, 116.6	1.17	0.146	30.85	18.23	2H-4, 95.6	7.86	0.519	6.22	9.11
1H-1, 134.5 1H-2, 5.5	1.55	0.176	22.85	19.18	2H-4, 116.5 2H-4, 146.5	8.36	0.524	5.15	7.26
1H-2, 41.5	1.91	0.204	18.37	24.07	2H-5, 35.5	8.75	0.542	7.20	15.60
1H-2, 71.5 2H-1 5 5	2.21	0.228	25.34	19.60	2H-5, 56.6	8.97	0.548	12.43	8.20
2H-1, 8.5	2.49	0.248	33.94	18.58	2H-5, 104.6	9.45	0.561	33.97	26.09
2H-1, 14.5	2.55	0.255	22.59	18.28	2H-5, 125.6	9.66	0.566	20.52	12.42
2H-1, 20.5 2H-1, 26.5	2.61	0.260	21.95	16.40	2H-5, 146.5 2H-6, 14,5	9.86	0.571	34.38	26.24
2H-1, 29.6	2.70	0.266	22.82	19.24	2H-6, 44.5	10.34	0.584	26.05	14.40
2H-1, 35.5	2.76	0.270	30.24	22.63	2H-6, 80.5	10.70	0.593	27.94	16.40
2H-1, 38.5 2H-1, 44.5	2.79	0.272	30.41	16.17	2H-6, 119.7 2H-6, 146.6	11.10	0.604	17.04	12.26
2H-1, 50.6	2.91	0.280	16.22	15.92	2H-7, 14.5	11.55	0.616	17.33	11.18
2H-1, 53.5	2.94	0.282	12.94	19.00	2H-7, 44.5	11.84	0.623	27.04	14.47
2H-1, 59.0 2H-1, 65.5	3.00	0.286	16.28	15.21	3H-1, 14.5 3H-1, 50.6	12.05	0.629	23.14	14.96
2H-1, 68.5	3.09	0.292	17.10	13.84	3H-1, 86.5	12.76	0.648	6.21	9.40
2H-1, 74.5	3.15	0.296	17.13	15.92	3H-1, 101.6	12.92	0.652	7.93	20.50
2H-1, 89.5	3.30	0.306	8.95	13.25	3H-1, 146.6	13.37	0.656	11.63	14.68
2H-1, 98.5	3.39	0.311	17.59	16.74	3H-2, 5.7	13.46	0.666	9.13	11.10
2H-1, 104.5	3.45	0.315	19.96	19.71	3H-2, 11.7	13.52	0.668	6.17	9.20
2H-1, 113.7	3.54	0.321	22.18	21.42	3H-2, 14.5 3H-2, 20.5	13.60	0.670	5.91	10.05
2H-1, 119.5	3.60	0.325	23.94	24.56	3H-2, 26.5	13.66	0.671	4.92	9.30
2H-1, 125.5 2H-1, 128.5	3.66	0.329	18.87	24.81	3H-2, 29.7	13.70	0.672	4.86	9.75
2H-1, 120.5 2H-1, 134.5	3.75	0.335	17.40	13.09	3H-2, 35.5 3H-2, 41.5	13.81	0.675	8.60	10.01
2H-1, 140.6	3.81	0.339	20.05	13.98	3H-2, 44.5	13.84	0.676	9.07	10.01
2H-1, 143.5 2H-1, 149.5	3.84	0.341	21.87	13.46	3H-2, 50.5	13.90	0.678	9.54	11.69
2H-2, 5.5	3.96	0.349	20.51	12.83	3H-2, 59.5 3H-2, 59.5	13.90	0.680	6.07	11.11
2H-2, 8.5	3.99	0.351	19.95	12.55	3H-2, 65.5	14.05	0.682	7.66	9.37
2H-2, 14.5 2H-2, 20.5	4.05	0.355	18.85	17.37	3H-2, 71.6	14.12	0.683	7.99	12.06
2H-2, 26.5	4.16	0.362	16.16	13.14	3H-2, 80.5	14.20	0.685	9.70	8.53
2H-2, 29.5	4.20	0.365	18.04	17.18	3H-2, 89.6	14.30	0.688	14.96	7.05
2H-2, 35.5 2H-2, 38.5	4.26	0.369	19.67	14.29	3H-2, 101.6 3H-2, 104.6	14.42	0.691	14.73	11.52
2H-2, 53.6	4.44	0.381	5.53	10.11	3H-2, 110.7	14.51	0.693	15.86	10.65
2H-2, 59.5	4.50	0.385	5.03	10.01	3H-2, 119.5	14.59	0.696	21.35	11.85
2H-2, 65.5 2H-2, 68.6	4.50	0.389	3.39	9.33	3H-2, 125.5 3H-2, 131.5	14.65	0.697	17.57	13.61
2H-2, 74.5	4.64	0.394	4.40	10.42	3H-2, 134.5	14.74	0.700	17.90	12.95
2H-2, 83.5	4.74	0.400	5.00	10.06	3H-2, 140.5	14.80	0.701	19.27	10.49
2H-2, 89.5 2H-2, 95.5	4.80	0.404	5.41	9.57	3H-2, 140.5 3H-3, 5.7	14.80	0.705	20.89	15.89
2H-2, 101.6	4.92	0.412	5.36	13.91	3H-3, 11.5	15.01	0.707	19.28	15.40
2H-2, 104.5 2H-2, 110.5	4.95	0.414	6.05	12.04	3H-3, 20.5	15.10	0.709	15.80	11.84
2H-2, 116.6	5.07	0.422	5.22	10.32	3H-3, 29.5	15.19	0.711	24.39	13.76
2H-2, 119.7	5.10	0.424	5.31	11.76	3H-3, 35.5	15.25	0.713	26.35	15.61
2H-2, 125.5 2H-2, 131.5	5.22	0.428	5.95	10.62	3H-3, 71.5 3H-3, 74.6	15.61	0.723	20.96	16.93
2H-2, 134.5	5.25	0.434	9.14	13.86	3H-3, 80.6	15.71	0.725	24.69	16.19
2H-2, 140.7	5.31	0.438	9.84	12.91	3H-3, 89.5	15.80	0.727	26.88	19.32
2H-2, 140.0 2H-3, 5.5	5.45	0.442	16.20	11.21	3H-3, 101.5	15.85	0.729	27.31	22.28
2H-3, 11.5	5.51	0.451	18.34	15.58	3H-3, 104.6	15.95	0.731	27.33	20.94
2H-3, 14.5 2H-3, 20.5	5.55	0.454	19.24	19.40	3H-3, 110.5	16.00	0.733	24.21	18.92
2H-3, 26.5	5.66	0.461	22.97	11.91	3H-3, 119.5	16.09	0.735	23.60	16.73
2H-3, 29.5	5.70	0.462	21.53	18.59	3H-3, 131.6	16.22	0.738	23.27	12.78
2H-3, 35.5 2H-3, 41.5	5.70	0.464	21.51	20.35	3H-3, 134.5 3H-3, 140.5	16.24	0.739	22.24	14.10
2H-3, 50.5	5.91	0.468	20.66	16.57	3H-3, 146.5	16.36	0.742	20.45	16.73
2H-3, 56.5	5.97	0.469	19.31	17.76	3H-4, 5.5	16.45	0.744	23.83	16.37
2H-3, 59.5 2H-3, 65.5	6.00	0.470	13.05	13.48	3H-4, 11.5 3H-4, 14, 5	16.51	0.746	24.00	16.73
2H-3, 71.5	6.12	0.473	10.96	12.36	3H-4, 20.6	16.61	0.749	26.24	17.82
2H-3, 74.6	6.15	0.474	10.77	19.60	3H-4, 26.5	16.66	0.750	25.13	14.89
2H-3, 89.6	6.30	0.475	12.75	15.43	3H-4, 29.5	16.82	0.754	24.36	15.11
2H-3, 95.5	6.36	0.479	18.50	18.20	3H-4, 44.6	16.85	0.755	19.98	18.43
2H-3, 110.6 2H-3, 116.5	6.51	0.483	13.30	12.29	3H-4, 65.5 3H-4, 71, 7	17.06	0.761	15.88	12.24
2H-3, 119.5	6.60	0.485	15.89	14.92	3H-4, 74.6	17.15	0.763	21.58	15.47
2H-3, 131.6	6.72	0.489	13.40	11.85	3H-4, 80.5	17.20	0.764	23.41	14.27

726

APPENDIX B (continue
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Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	Fine-grained lithogenic (wt%)	Noncarb. – fine-grained lithogenic (wt%)
3H-4, 89.5	17.30	0.767	31.33	21.04
3H-4, 95.5	17.35	0.768	35.60	19.98
3H-4, 101.6	17.42	0.770	33.16	18.19
3H-4, 104.7 3H-4, 110.5	17.45	0.771	39.61	14.67
3H-4, 119.5	17.59	0.775	14.69	12.81
3H-4, 125.5	17.65	0.776	15.13	14.41
3H-4, 131.5	17.72	0.778	15.79	12.52
3H-4, 140.5	17.81	0.781	17.48	14.39
3H-4, 146.5	17.86	0.784	21.41	13.49
3H-5, 11.5	18.01	0.795	24.10	21.29
3H-5, 41.5	18.32	0.805	19.93	14.78
3H-5, 71.5	18.61	0.837	13.49	10.38
3H-5, 101.5	18.91	0.858	26.16	14.13
3H-6, 11.5	19.19	0.900	2.15	32.10
3H-6, 41.7	19.82	0.921	16.46	11.74
3H-6, 71.5	20.11	0.942	24.31	14.88
3H-6, 119.5	20.59	0.905	24.23	11.92
3H-6, 140.7	20.81	0.993	7.43	9.10
3H-7, 5.5 3H-7, 26,5	20.95	1.033	21.97	12.62
3H-7, 44.6	21.35	1.148	21.30	10.70
4H-1, 14.5	21.55	1.206	22.89	16.11
4H-1, 56.5 4H-1, 74.6	21.97	1.260	13.17	14.04
4H-1, 101.5	22.41	1.291	6.23	10.07
4H-1, 131.5	22.72	1.312	13.70	10.76
4H-1, 146.5 4H-2, 14, 5	22.86	1.321	20.40	3 74
4H-2, 41.6	23.32	1.353	9.79	11.77
4H-2, 44.5	23.34	1.354	9.38	11.05
4H-2, 50.5 4H-2, 56.6	23.40	1.358	9.33	10.05
4H-2, 59.5	23.49	1.365	11.62	9.09
4H-2, 65.5	23.56	1.369	11.63	14.06
4H-2, /1./ 4H-2, 74.7	23.62	1.375	21.53	8.47
4H-2, 80.5	23.70	1.379	19.84	14.89
4H-2, 89.5	23.80	1.386	11.02	18.07
4H-2, 95.6 4H-2, 101.5	23.86	1.390	6.83	10.08
4H-2, 104.5	23.94	1.395	4.93	9.89
4H-2, 110.6	24.01	1.400	5.66	9.56
4H-2, 110.5 4H-2, 119.5	24.07	1.404	5.97	10.84
4H-2, 125.5	24.15	1.410	7.03	8.35
4H-2, 131.5	24.22	1.415	6.19	10.13
4H-2, 134.5 4H-2, 140.5	24.24	1.410	5.69	9.86
4H-2, 146.6	24.37	1.425	9.00	7.39
4H-2, 149.5	24.40	1.427	13.31	9.86
4H-3, 11.5	24.40	1.431	22.41	18.28
4H-3, 14.6	24.55	1.437	22.97	15.02
4H-3, 20.5	24.60	1.441	28.15	8.52
4H-3, 29.5	24.69	1.447	26.32	15.40
4H-3, 35.7	24.76	1.452	28.71	16.10
4H-3, 41.6 4H-3, 44.5	24.82	1.456	27.72	14.47
4H-3, 50.5	24.90	1.459	11.03	10.56
4H-3, 56.5	24.97	1.462	12.42	7.97
4H-3, 59.5 4H-3, 65.5	24.99	1.463	7.83	10.39
4H-3, 71.5	25.11	1.468	7.25	14.17
4H-3, 74.6	25.15	1.470	7.31	15.21
4H-5, 80.6 4H-3, 86.5	25.21	1.472	6.70	8.16
4H-3, 89.6	25.30	1.476	6.78	10.06
4H-3, 95.6	25.36	1.478	6.27	10.08
4H-3, 101.5 4H-3, 104.5	25.41	1.480	9.14	11.13
4H-3, 110.5	25.50	1.484	18.19	13.09
4H-3, 116.6	25.57	1.487	22.91	17.51
4H-3, 119.5 4H-3, 125.5	25.59	1.487	25.61	11.92
4H-3, 134.5	25.74	1.494	32.59	11.05
4H-3, 140.7	25.81	1.497	27.87	14.42
4H-3, 146.5 4H-3, 149.7	25.86	1.498	19.86	17.69
4H-4, 5.5	25.95	1.502	3.61	16.64
4H-4, 11.5	26.01	1.505	6.99	11.43
4H-4, 14.5 4H-4, 20.5	26.05	1.506	5.91	9.91

Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	Fine-grained lithogenic (wt%)	Noncarb. – fine-grained lithogenic (wt%)
44-4 29.6	26.20	1.512	717	12.26
4H-4, 35.5	26.25	1.514	6.88	12.69
4H-4, 41.5	26.32	1.517	7.44	11.22
4H-4, 44.5	26.34	1.518	9.05	12.69
4H-4, 50.5	26.40	1.520	13.63	13.43
4H-4, 59.6	26.50	1.525	16.18	18.23
4H-4, 65.5	26.56	1.527	17.03	19.13
4H-4, 71.6	26.62	1.529	21.22	17.52
4H-4, /4.5	26.65	1.531	25.07	8.87
4H-4, 89.5	26.80	1.537	14.81	24.39
4H-4, 95.7	26.86	1.539	20.45	14.67
4H-4, 101.6	26.92	1.542	25.02	18.48
4H-4, 104.5	26.94	1.545	23.32	22 37
4H-4, 116.6	27.07	1.548	10.44	12.45
4H-4, 119.5	27.09	1.549	8.87	14.36
4H-4, 125.7	27.16	1.551	9.24	10.74
4H-4, 131.0 4H-4, 134.5	27.22	1.554	13.20	22.00
4H-4, 140.6	27.31	1.558	9.71	15.28
4H-4, 146.5	27.36	1.560	10.29	11.19
4H-5, 14.6	27.55	1.567	11.66	10.32
4H-5, 29.5 4H-5, 41.5	27.82	1.578	29.93	18.52
4H-5, 56.6	27.97	1.584	14.72	14.77
4H-5, 74.6	28.15	1.592	13.31	9.25
4H-5, 101.0 4H-5, 110.5	28.42	1.010	8.53	11.18
4H-5, 125.5	28.65	1.969	15.86	13.14
4H-5, 134.5	28.74	1.976	27.50	20.49
4H-5, 149.5	28.90	1.987	8.15	6.67
4H-6, 44,5	29.19	2.016	12.36	13.10
4H-6, 59.6	29.50	2.027	10.46	8.86
4H-6, 71.5	29.61	2.035	15.39	21.08
4H-6, 89.5 4H-6, 125.5	30.15	2.048	9.12	13.64
4H-7, 11.5	30.51	2.096	11.62	1.000.000
4H-7, 26.5	30.66	2.106	8.38	0.05
5H-1, 20.5 5H-1, 59.5	31.10	2.130	4.11	12.26
5H-1, 80.5	31.70	2.177	27.20	17.84
5H-1, 110.5	32.01	2.197	5.99	9.54
5H-1, 134.5	32.24	2.212	14.72	12.80
5H-2, 14.5 5H-2, 44.5	32.85	2.253	25.03	13.39
5H-2, 74.5	33.15	2.272	7.85	10.26
5H-2, 104.6	33.45	2.292	21.43	18.79
5H-2, 119.5 5H-2, 143.5	33.80	2.302	9.44	12.22
5H-3, 14.5	34.05	2.332	15.89	17.03
5H-3, 29.6	34.20	2.342	11.52	14.51
5H-3, 44.5 5H-3, 74.6	34.55	2.352	10.22	12.07
5H-3, 104.5	34.94	2.391	4.72	9.92
5H-3, 131.5	35.22	2.409	19.01	16.22
5H-3, 146.5	35.37	2.419	17.88	14.60
5H-4, 26.5	35.66	2.438	8.24	11.64
5H-4, 50.5	35.90	2,454	20.50	20.85
5H-4, 74.5	36.14	2.470	5.80	11.02
5H-4, 107.5	36.80	2.514	27.71	29.00
5H-5, 14.5	37.05	2.530	5.76	9.14
5H-5, 50.5	37.41	2.554	10.88	15.39
5H-5, 02.5	37.65	2.562	13.00	8.64
5H-5, 110.5	38.01	2.594	12.32	14.53
5H-5, 119.5	38.10	2.600	36.08	19.08
5H-5, 140.5 5H-6, 5,5	38.31	2.614	8.05	6.92
5H-6, 29.5	38.69	2.638	21.86	12.58
5H-6, 56.5	38.97	2.657	4.45	7.61
5H-6, 95.5 5H-6, 125.5	39.36	2.682	4.50	9.84
5H-6, 146.6	39.87	2.716	7.64	10.73
5H-7, 20.5	40.11	2.731	9.83	13.13
5H-7, 41.5	40.32	2.745	30.05	20.00

Notes: Depth (mbsf) = sub-bottom depth in meters below seafloor; age (Ma) = millions of years ago; fine-grained lithogenic (wt%) = weight percentage of chemically extracted lithogenic material (<63  $\mu$ m) relative to bulk sample weight; Noncarb. – fine-grained lithogenic (wt%) = weight percentage of noncarbonate material (100% – wt% CaCO<sub>3</sub>; data from Rack et al., 1993) minus the weight percentage of fine-grained lithogenic material (<63  $\mu$ m).

APPENDIX C
Linear Sedimentation and Mass Accumulation Rates for Samples from Hole 810C (Shatsky Rise)

Core, section,	Depth	Age	LSR	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(c	$cm^2 \times k.y$	.)-1		$mg(cm^2 \times k.y.)^{-1}$
132-810C-			51 M.S.		00000		-	10-64-5	0.00110744	
1H-1, 17.5	0.17	0.035	0.482	0.80	0.386	0.253	0.132	0.075	0.057	57
1H-1, 23.5 1H-1, 29.5	0.23	0.048	0.482	0.78	0.376	0.230	0.146			
1H-1, 35.7	0.36	0.075	0.482	0.73	0.352	0.219	0.133			
1H-1, 41.6	0.42	0.086	1.257	0.73	0.918	0.522	0.396			
1H-1, 44.5	0.44	0.087	1.257	0.73	0.918	0.506	0.412	0.106	0.210	210
1H-1, 50.5 1H-1, 53.5	0.50	0.092	1.257	0.74	0.930	0.515	0.410	0.196	0.219	219
1H-1, 59.5	0.60	0.100	1.257	0.73	0.918	0.534	0.384			
1H-1, 65.5	0.65	0.104	1.257	0.81	1.018	0.582	0.436			
1H-1, 68.6	0.69	0.107	1.257	0.72	0.905	0.562	0.343			
1H-1, 74.5 1H-1, 80.5	0.75	0.112	1.257	0.86	1.081	0.619	0.462	0 194	0.232	232
1H-1, 83.5	0.83	0.118	1.257	0.88	1.106	0.656	0.450	0.194	0.202	2012
1H-1, 89.5	0.89	0.123	1.257	0.80	1.006	0.606	0.399			
1H-1, 95.5	0.95	0.128	1.257	0.81	1.018	0.605	0.413			
1H-1, 98.5	1.05	0.131	1.257	0.75	1.207	0.559	0.358			
1H-1, 110.6	1.11	0.141	1.257	0.85	1.068	0.554	0.515			
1H-1, 116.6	1.17	0.146	1.257	0.84	1.056	0.538	0.518	0.192	0.326	326
1H-1, 119.5	1.20	0.148	1.257	0.90	1.131	0.672	0.459			
1H-1, 125.5 1H-1, 128.5	1.25	0.152	1.257	0.83	1.043	0.657	0.380			
1H-1, 128.5	1.35	0.160	1.257	0.94	1.182	0.778	0.404	0.134	0.270	270
1H-1, 140.5	1.40	0.164	1.257	0.80	1.006	0.608	0.398			12011
1H-1, 146.5	1.47	0.169	1.257	0.83	1.043	0.525	0.519			
1H-1, 149.5	1.50	0.172	1.257	0.80	1.006	0.416	0.590	0.100	0.200	200
1H-2, 5.5 1H-2, 8.5	1.55	0.176	1.257	0.79	0.993	0.513	0.480	0.190	0.290	290
1H-2, 17.5	1.67	0.185	1.257	0.75	0.943	0.496	0.447			
1H-2, 20.5	1.71	0.188	1.257	0.76	0.955	0.517	0.439			
1H-2, 23.5	1.74	0.191	1.257	0.74	0.930	0.546	0.384			
1H-2, 29.5 1H-2, 35,5	1.79	0.195	1.257	0.80	1.006	0.615	0.391			
1H-2, 41.5	1.91	0.204	1.257	0.90	1.131	0.651	0.480	0.272	0.208	208
1H-2, 44.5	1.95	0.207	1.257	0.73	0.918	0.502	0.415			
1H-2, 65.5	2.15	0.223	1.257	0.83	1.043	0.533	0.510	0.000	0.407	107
2H-1, 5.5 2H-1, 8.5	2.46	0.248	1.257	0.83	1.043	0.413	0.630	0.203	0.427	427
2H-1, 14.5	2.55	0.255	1.257	0.87	1.031	0.609	0.421	0.188	0.233	233
2H-1, 20.5	2.61	0.260	1.517	0.84	1.274	0.786	0.489	0.209	0.280	280
2H-1, 26.5	2.67	0.264	1.517	0.78	1.183	0.698	0.486	0.242	0.244	244
2H-1, 29.6	2.70	0.266	1.517	0.81	1.229	0.712	0.517	0.236	0.280	280
2H-1, 35.5 2H-1, 38.5	2.70	0.270	1.517	0.82	1.244	0.580	0.658	0.282	0.376	360
2H-1, 44.5	2.85	0.276	1.517	0.93	1.411	0.721	0.689	0.426	0.264	264
2H-1, 50.6	2.91	0.280	1.517	0.90	1.365	0.926	0.439	0.217	0.222	222
2H-1, 53.5	2.94	0.282	1.517	0.98	1.487	1.012	0.475	0.282	0.192	192
2H-1, 59.0 2H-1, 65.5	3.00	0.280	1.517	0.90	1.305	0.963	0.403	0.180	0.222	222
2H-1, 68.5	3.09	0.292	1.517	0.93	1.411	0.974	0.432	0.195	0.241	241
2H-1, 74.5	3.15	0.296	1.517	0.94	1.426	0.955	0.471	0.227	0.244	244
2H-1, 80.5	3.21	0.299	1.517	0.93	1.411	1.107	0.304	0.005	0.100	122
2H-1, 83.5 2H-1, 89.5	3.24	0.301	1.517	0.94	1.420	1.069	0.357	0.225	0.132	132
2H-1, 95.6	3.36	0.309	1.517	0.88	1.335	0.994	0.341	0.101	0.122	122
2H-1, 98.5	3.39	0.311	1.517	0.83	1.259	0.827	0.432	0.211	0.222	222
2H-1, 104.5	3.45	0.315	1.517	0.83	1.259	0.760	0.499	0.248	0.251	251
2H-1, 110.7	3.51	0.319	1.517	0.82	1.244	0.712	0.532	0.314	0.218	218
2H-1, 119.5	3.60	0.325	1.517	0.75	1.138	0.586	0.552	0.279	0.272	272
2H-1, 125.5	3.66	0.329	1.517	0.86	1.305	0.735	0.570	0.324	0.246	246
2H-1, 128.5	3.69	0.331	1.517	0.87	1.320	0.702	0.618	0.370	0.247	247
2H-1, 134.5 2H-1, 140.6	3.75	0.335	1.517	0.79	1.198	0.833	0.365	0.157	0.209	209
2H-1, 143.5	3.84	0.341	1.517	0.90	1.365	0.883	0.490	0.184	0.292	299
2H-1, 149.5	3.90	0.345	1.517	0.88	1.335	0.840	0.495	0.215	0.280	280
2H-2, 5.5	3.96	0.349	1.517	0.84	1.274	0.849	0.425	0.163	0.261	261
2H-2, 8.5	3.99	0.351	1.517	0.84	1.274	0.860	0.414	0.160	0.254	254
2H-2, 14.5 2H-2, 20.5	4.05	0.355	1.517	0.87	1.320	0.842	0.478	0.229	0.249	249
2H-2, 26.5	4.16	0.362	1.517	0.88	1.335	0.944	0.391	0.175	0.216	216
2H-2, 29.5	4.20	0.365	1.517	0.92	1.396	0.904	0.492	0.240	0.252	252
2H-2, 35.5	4.26	0.369	1.517	0.90	1.365	0.902	0.464	0.195	0.269	269
2H-2, 38.5 2H-2, 44.5	4.29	0.371	1.517	0.88	1.335	0.867	0.468	0.253	0.215	215
2H-2, 50.6	4.41	0.379	1.517	0.85	1.305	1.094	0.211			
2H-2, 53.6	4.44	0.381	1.517	0.86	1.305	1.101	0.204	0.132	0.072	72
2H-2, 59.5	4.50	0.385	1.517	0.90	1.365	1.160	0.205	0.137	0.069	69
2H-2, 65.5 2H-2, 68.6	4.56	0.389	1.517	0.91	1.380	1.202	0.179	0.129	0.050	50
2H-2, 74.5	4.64	0.394	1.517	0.84	1.274	1.085	0.189	0.133	0.049	56
2H-2, 80.5	4.70	0.398	1.517	0.87	1.320	1.119	0.200			
2H-2, 83.5	4.74	0.400	1.517	0.91	1.380	1.173	0.208	0.139	0.069	69
2H-2, 89.5 2H-2, 05.5	4.80	0.404	1.517	0.83	1.259	1.039	0.220	0.156	0.064	64
2H-2, 101.6	4.92	0.412	1.517	0.85	1.289	1.041	0.248	0.179	0.062	69

					MAD	CAR	NCAD	OMAD	TMAD	TMAP2
Core, section,	Depth	Age	LSR	DBD	MAK	CAR	NCAR	1-1	IMAK	$ma(cm^2 \times k \times )^{-1}$
Interval (cm)	(mbsr)	(Ma)	(cm/k.y.)	(g/cm <sup>-</sup> )		g(c	:ш хк.у.	,		mg(cm × k.y.)
2H-2, 104.5	4.95	0.414	1.517	0.89	1.350	1.106	0.244	0.163	0.082	82
2H-2, 110.5 2H-2, 116.6	5.01	0.418	1.517	0.85	1.289	1.268	0.201	0.185	0.073	78
2H-2, 119.7	5.10	0.424	1.517	0.83	1.259	1.044	0.215	0.148	0.067	67
2H-2, 125.5 2H-2, 131.5	5.16	0.428	1.517	0.93	1.411	1.177	0.234	0.150	0.084	84 93
2H-2, 134.5	5.25	0.434	1.517	0.83	1.259	0.970	0.290	0.175	0.115	115
2H-2, 140.7	5.31	0.438	1.517	0.80	1.214	0.938	0.276	0.157	0.119	119
2H-2, 140.0 2H-3, 5.5	5.45	0.442	1.517	0.81	1.229	0.892	0.337	0.139	0.202	202
2H-3, 11.5	5.51	0.451	1.517	0.81	1.229	0.812	0.417	0.191	0.225	225
2H-3, 14.5 2H-3, 20.5	5.61	0.454	1.517	0.78	1.183	0.726	0.457	0.230	0.228	228
2H-3, 26.5	5.66	0.461	3.806	0.81	3.083	2.008	1.075	0.367	0.708	708
2H-3, 29.5	5.70	0.462	3.806	0.80	3.045	1.823	1.222	0.566	0.656	656
2H-3, 41.5	5.82	0.465	3.806	0.80	3.045	1.711	1.334	0.620	0.715	715
2H-3, 44.6	5.85	0.466	3.806	0.82	3.121	1.848	1.273	0.525	0.770	((0)
2H-3, 50.5 2H-3, 56,5	5.91	0.468	3.806	0.85	3.235	2.031	1.204	0.536	0.668	595
2H-3, 59.5	6.00	0.470	3.806	0.84	3.197	2.015	1.182	0.586	0.596	596
2H-3, 65.5	6.06	0.472	3.806	0.85	3.235	2.370	0.865	0.436	0.429	429
2H-3, 71.5 2H-3, 74.6	6.12	0.473	3.806	0.88	3.349	2.568	1.110	0.414	0.307	394
2H-3, 80.5	6.20	0.475	3.806	0.85	3.235	2.450	0.785	0.370	0.415	415
2H-3, 86.5	6.27	0.477	3.806	0.98	3.730	2.728	1.002	0.623	0.634	634
2H-3, 101.6	6.42	0.481	3.806	0.88	3.349	2.147	1.202	0.025	0.054	0.54
2H-3, 104.6	6.45	0.482	3.806	0.90	3.425	2.442	0.983	0.104	0.471	101
2H-3, 110.6 2H-3, 116.5	6.51	0.483	3.806	0.91	3.463	2.577	0.886	0.426	0.461	524
2H-3, 119.5	6.60	0.486	3.806	0.84	3.197	2.212	0.985	0.477	0.508	508
2H-3, 125.7	6.66	0.487	3.806	0.84	3.197	2.151	1.046	0.406	0.450	450
2H-3, 131.6 2H-3, 134.6	6.75	0.489	3.806	0.90	3.425	2.804	0.865	0.408	0.439	346
2H-3, 140.7	6.81	0.491	3.806	1.00	3.806	2.936	0.870	0.509	0.361	361
2H-3, 146.5 2H-4, 2, 5	6.87	0.493	3.806	0.94	3.578	2.804	0.773	0.333	0.440	440
2H-4, 5.5	6.95	0.495	3.806	0.88	3.349	2.425	0.924			
2H-4, 11.5	7.01	0.496	3.806	0.91	3.463	2.126	1.338			
2H-4, 14.6 2H-4, 20.5	7.05	0.498	3.806	0.93	3.540	2.338	1.201			
2H-4, 26.5	7.16	0.500	3.806	0.91	3.463	2.356	1.107			
2H-4, 29.5	7.20	0.502	3.806	0.93	3.540	2.472	1.068			
2H-4, 33.5 2H-4, 41.5	7.32	0.505	3.806	0.90	3.425	2.995	0.349			
2H-4, 44.6	7.35	0.505	3.806	0.98	3.730	3.169	0.561			
2H-4, 50.6 2H-4, 56, 5	7.41	0.507	3.806	0.95	3.616	3.023	0.593			
2H-4, 59.5	7.50	0.509	3.806	0.88	3.349	2.046	1.303	1.011	0.292	292
2H-4, 65.5	7.56	0.511	3.806	0.93	3.540	2.936	0.603			
2H-4, 08.5 2H-4, 74.5	7.64	0.512	3.806	0.98	3.730	3.110	0.620	0.374	0.247	247
2H-4, 80.6	7.71	0.515	3.806	1.01	3.844	2.843	1.001			
2H-4, 86.6 2H-4, 89.6	7.77	0.516	3.806	0.98	3.730	3.148	0.582			
2H-4, 95.6	7.86	0.519	3.806	0.98	3.730	3.158	0.572	0.340	0.232	232
2H-4, 101.5	7.91	0.520	3.806	0.96	3.654	2.968	0.686			
2H-4, 104.5 2H-4, 110.7	8.01	0.523	3.806	0.93	3.463	2.738	0.725			
2H-4, 116.5	8.07	0.524	3.806	0.99	3.768	2.927	0.841	0.384	0.457	457
2H-4, 119.0 2H-4, 125.5	8.10	0.525	3.806	0.93	3.540	2.792	0.747			
2H-4, 131.5	8.22	0.528	3.806	0.98	3.730	3.154	0.576			
2H-4, 134.5	8.24	0.529	3.806	1.05	3.996	3.411	0.585			
2H-4, 140.0 2H-4, 146.5	8.36	0.531	3.806	0.94	3.768	3.300	0.399	0.274	0.194	194
2H-4, 149.5	8.40	0.533	3.806	0.91	3.463	2.901	0.562			
2H-5, 5.5 2H-5, 11, 5	8.45	0.534	3.806	0.93	3.540	2.902	0.637			
2H-5, 14.5	8.55	0.537	3.806	0.98	3.730	2.941	0.789			
2H-5, 20.5	8.60	0.538	3.806	0.98	3.730	3.001	0.728			
2H-5, 20.5 2H-5, 29.5	8.69	0.540	3.806	1.05	3.768	3.027	0.741			
2H-5, 35.5	8.75	0.542	3.806	1.01	3.844	2.968	0.876	0.600	0.277	277
2H-5, 41.6	8.82	0.544	3.806	1.01	3.844	3.073	0.772			
2H-5, 50.5	8.90	0.546	3.806	0.98	3.730	2.999	0.731			
2H-5, 56.6	8.97	0.548	3.806	1.01	3.844	3.051	0.793	0.315	0.478	478
2H-5, 59.6 2H-5, 65.5	9.00	0.549	3.806	0.96	3.654	2.703	0.951			
2H-5, 71.6	9.12	0.552	3.806	1.01	3.844	2.775	1.069			
2H-5, 74.5	9.14	0.552	3.806	0.98	3.730	2.712	1.018			
2H-5, 80.6 2H-5, 86.5	9.21	0.554	3.806	0.90	3.425	2.441	0.984			
2H-5, 89.5	9.30	0.557	3.806	0.91	3.463	2.453	1.011	0.384	0.627	627
2H-5, 95.6 2H-5, 101.5	9.36	0.558	3.806	0.93	3.540	2.292	1.247			
2H-5, 104.6	9.45	0.561	3.806	0.95	3.616	1.444	2.172	0.943	1.228	1228
2H-5, 110.5	9.51	0.562	3.806	0.88	3.349	2.063	1.286			
2H-5, 116.5	9.56	0.563	3.806	0.93	3.540	2.439	1.100			

APPENDIX C (continued).

Core, section.	Depth	Age	LSR	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(c	$cm^2 \times k.y$	.) <sup>-1</sup>		$mg(cm^2 \times k.y.)^{-1}$
2H-5, 119.5	9.59	0.564	3.806	0.96	3.654	2.525	1.129			
2H-5, 125.6	9.66	0.566	3.806	0.93	3.540	2.374	1.166	0.440	0.726	726
2H-5, 131.5	9.72	0.568	3.806	0.93	3.540	2.119	1.420			
2H-5, 154.5 2H-5, 140.5	9.74	0.508	3.806	1.08	3.425	2.125	2.030			
2H-5, 146.5	9.86	0.571	3.806	0.90	3.425	1.349	2.076	0.899	1.178	1178
2H-5, 149.5	9.89	0.572	3.806	0.93	3.540	1.567	1.973			
2H-6, 5.5 2H-6, 11, 5	9.95	0.574	3.806	0.91	3,235	2 036	1.029			
2H-6, 14.5	10.05	0.576	3.806	0.80	3.045	1.914	1.131	0.365	0.766	766
2H-6, 20.5	10.10	0.578	3.806	0.85	3.235	1.968	1.267			
2H-0, 20.0 2H-6, 20.5	10.17	0.580	3,806	0.83	3.159	1.818	1.341			
2H-6, 35.6	10.26	0.582	3.806	0.84	3.197	1.896	1.301			
2H-6, 41.6	10.32	0.583	3.806	0.84	3.197	1.946	1.251			
2H-6, 44.5	10.34	0.584	3.806	0.84	3.197	1.904	1.293	0.460	0.833	833
2H-6, 56.6	10.40	0.587	3.806	0.83	3.121	1.766	1.355			
2H-6, 59.5	10.49	0.588	3.806	0.81	3.083	1.793	1.290			
2H-6, 65.5	10.55	0.590	3.806	0.90	3.425	1.770	1.655			
2H-6, 71.5 2H-6, 74, 5	10.61	0.591	3.806	0.88	3.349	2.097	1.252			
2H-6, 80.5	10.70	0.593	3.806	0.85	3.235	1.801	1.434	0.531	0.904	904
2H-6, 86.5	10.76	0.595	3.806	0.94	3.578	2.248	1.329			
2H-6, 89.5	10.80	0.596	3.806	1.03	3.920	2.290	1.630			
2H-6, 101.6	10.80	0.598	3.806	0.99	3.578	2.420	1.407			
2H-6, 104.6	10.95	0.600	3.806	0.89	3.387	2.121	1.266			
2H-6, 110.6	11.01	0.602	3.806	0.94	3.578	2.022	1.555			
2H-6, 110.7 2H-6, 119.7	11.07	0.603	3.806	0.78	2.969	2 502	1.002	0.434	0.603	603
2H-6, 125.5	11.15	0.605	3.806	0.84	3.197	2.251	0.946	0.454	0.005	005
2H-6, 134.6	11.25	0.608	3.806	0.88	3.349	1.974	1.375			
2H-6, 140.5	11.30	0.609	3.806	1.05	3.996	2.249	1.748	0.506	1.111	
2H-6, 140.0 2H-6, 149.5	11.37	0.611	3.806	0.93	3,540	1.855	1.935	0.596	1.111	1111
2H-7, 5.5	11.45	0.613	3.806	0.85	3.235	1.547	1.688			
2H-7, 8.5	11.48	0.614	3.806	0.93	3.540	2.381	1.159	0.100		(22
2H-7, 14.5 2H-7, 20,5	11.55	0.616	3.806	0.96	3.654	2.612	1.042	0.408	0.633	633
2H-7, 26.5	11.66	0.619	3.806	0.90	3.425	2.445	0.980			
2H-7, 29.5	11.69	0.619	3.806	0.91	3.463	2.087	1.376			
2H-7, 35.5	11.75	0.621	3.806	0.90	3.425	2.191	1.235			
2H-7, 41.5 2H-7, 44.5	11.84	0.623	3.806	0.89	3.273	1.914	1.359	0.474	0.885	885
2H-7, 50.5	11.90	0.625	3.806	0.88	3.349	2.007	1.343	and the second	0100000	0.00.0000
3H-1, 5.5	11.95	0.626	3.806	0.87	3.311	2.345	0.966			
3H-1, 11.5 3H-1, 14.5	12.01	0.628	3.806	0.88	3,349	2.406	0.943	0 304	0.473	473
3H-1, 20.5	12.10	0.630	3.806	0.93	3.540	2.825	0.715	0.021	0.110	
3H-1, 26.5	12.16	0.632	3.806	0.96	3.654	2.949	0.705			
3H-1, 29.5	12.19	0.633	3.806	0.93	3.540	2.680	0.860			
3H-1, 41.5	12.31	0.636	3.806	0.87	3.311	2.346	0.966			
3H-1, 44.5	12.34	0.637	3.806	0.83	3.159	1.993	1.166		The contract of	
3H-1, 50.6	12.41	0.638	3.806	0.72	2.740	1.696	1.044	0.410	0.634	634
3H-1, 65.5	12.55	0.642	3.806	0.90	3.121	2.370	0.751			
3H-1, 71.5	12.61	0.644	3.806	0.83	3.159	2.639	0.520			
3H-1, 74.6	12.65	0.645	3.806	0.77	2.931	2.508	0.423			
3H-1, 80.5 3H-1, 86.5	12.70	0.646	3.806	0.82	3.121	2.696	0.425	0.311	0.206	206
3H-1, 95.6	12.86	0.650	3.806	0.87	3.311	2.440	0.871	0.011	01800	
3H-1, 101.6	12.92	0.652	3.806	0.91	3.463	2.479	0.985	0.710	0.275	275
3H-1, 104.7	12.95	0.653	3.806	0.96	3.654	3.033	0.620			
3H-1, 116.6	13.07	0.656	3.806	0.93	3.197	2.688	0.509			
3H-1, 119.5	13.09	0.656	3.806	0.99	3.768	3.180	0.588	0.344	0.244	244
3H-1, 125.5	13.15	0.658	3.806	1.04	3.958	3.378	0.581			
3H-1, 131.7	13.22	0.660	3.806	0.90	3.425	2.828	0.598			
3H-1, 140.5	13.30	0.662	3.806	0.96	3.654	2.877	0.776			
3H-1, 146.6	13.37	0.664	3.806	0.83	3.159	2.328	0.831	0.464	0.367	367
3H-2, 5.7	13.46	0.666	3.806	0.98	3.730	2.975	0.755	0.414	0.341	341
3H-2, 14.5	13.55	0.668	3.806	1.01	3.844	3.234	0.610	0.343	0.230	233
3H-2, 20.5	13.60	0.670	3.806	0.99	3.768	3.167	0.601	0.379	0.223	223
3H-2, 26.5	13.66	0.671	3.806	0.99	3.768	3.232	0.536	0.350	0.185	185
3H-2, 35.5 3H-2, 41 5	13.75	0.674	3,806	1.01	3.844	3.129	0.715	0.385	0.193	331
3H-2, 44.5	13.84	0.676	3.806	0.98	3.730	3.018	0.712	0.373	0.338	338
3H-2, 50.5	13.90	0.678	3.806	0.90	3.425	2.698	0.727	0.400	0.327	327
3H-2, 56.5	13.96	0.679	3.806	0.98	3.730	3.011	0.719	0.389	0.330	330
3H-2, 65.5	14.05	0.682	3.806	1.01	3,844	3.189	0.655	0.360	0.295	294
3H-2, 71.6	14.12	0.683	3.806	0.93	3.540	2.830	0.710	0.427	0.283	283
3H-2, 74.7	14.15	0.684	3.806	1.03	3.920	3.077	0.843	0.477	0.366	366
3H-2, 80.5 3H-2, 86.6	14.20	0.685	3.806	1.03	3.920	3.050	0.765	0.318	0.362	502
3H-2, 89.6	14.30	0.688	3.806	0.94	3.578	2.790	0.787	0.252	0.535	535
3H-2, 95.5	14.35	0.689	3.806	0.94	3.578	2.620	0.958			

						commute				
Core, section.	Depth	Age	LSR	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(c	$m^2 \times k.y.$	.) <sup>-1</sup>		$mg(cm^2 \times k.y.)^{-1}$
				1000	212/24		-	e 		
3H-2, 101.6	14.42	0.691	3.806	1.01	3.844	2.835	1.009	0.443	0.566	566
3H-2, 1104.0	14.51	0.693	3.806	1.01	3.844	2.825	1.030	0.420	0.610	610
3H-2, 116.6	14.57	0.695	3.806	1.05	3.996	2.918	1.078	01107	0.010	
3H-2, 119.5	14.59	0.696	3.806	0.96	3.654	2.441	1.213	0.433	0.780	780
3H-2, 125.5 3H-2, 131.5	14.65	0.697	3.806	0.90	3.425	2.369	1.056	0.454	0.602	602
3H-2, 134.5	14.74	0.700	3.806	0.94	3.578	2.401	1.104	0.463	0.640	640
3H-2, 140.5	14.80	0.701	3.806	0.87	3.311	2.326	0.985	0.347	0.638	638
3H-2, 146.5	14.86	0.703	3.806	0.94	3.578	2.465	1.113	0.296	0.817	817
3H-3, 5.7	14.90	0.705	3.806	0.88	3.549	2.117	1.232	0.532	0.700	700
3H-3, 14.6	15.05	0.708	3.806	0.99	3.768	2.738	1.030	0.001	0.070	070
3H-3, 20.5	15.10	0.709	3.806	0.93	3.540	2.561	0.978	0.419	0.559	559
3H-3, 26.5	15.16	0.711	3.806	0.81	3.083	2.041	1.042	0.301	0.741	741
3H-3, 29.5	15.19	0.713	3.806	1.09	3.578	2.213	1.305	0.492	1.093	1093
3H-3, 41.6	15.32	0.715	3.806	0.91	3.463	2.102	1.362	0.010	1.075	1070
3H-3, 44.5	15.34	0.715	3.806	0.87	3.311	1.842	1.469			
3H-3, 50.6	15.41	0.717	3.806	0.88	3.349	1.856	1.494			
3H-3, 59.5	15.49	0.719	3.806	0.99	3 654	1.989	1 973			
3H-3, 65.5	15.55	0.721	3.806	1.03	3.920	2.410	1.510			
3H-3, 71.5	15.61	0.723	3.806	1.01	3.844	2.434	1.410	0.605	0.806	806
3H-3, 74.6	15.65	0.724	3.806	0.92	3.502	2.122	1.380	0.593	0.787	787
3H-3, 86.6	15.77	0.723	3.806	0.87	3.197	2.013	1.185	0.330	0.017	010
3H-3, 89.5	15.80	0.727	3.806	0.90	3.425	1.843	1.583	0.662	0.921	921
3H-3, 95.5	15.85	0.729	3.806	0.82	3.121	1.584	1.537	0.731	0.806	806
3H-3, 101.5 3H-3, 104.6	15.91	0.730	3.806	0.80	3.045	1.535	1.510	0.678	0.832	832
3H-3, 110.5	16.00	0.733	3.806	0.80	3.045	1.732	1.313	0.576	0.737	737
3H-3, 116.6	16.07	0.734	3.806	0.83	3.159	1.801	1.358	0.569	0.789	789
3H-3, 119.5	16.09	0.735	3.806	0.99	3.768	2.248	1.520	0.630	0.889	889
3H-3, 125.0 3H-3, 131.6	16.10	0.737	3.806	0.93	3.540	2.128	1.412	0.443	0.806	806
3H-3, 134.5	16.24	0.739	3.806	0.93	3.540	2.253	1.286	0.499	0.787	787
3H-3, 140.5	16.31	0.741	3.806	0.96	3.654	2.292	1.362	0.592	0.770	770
3H-3, 146.5	16.36	0.742	3.806	0.88	3.349	2.104	1.245	0.560	0.685	685
3H-4, 5.5 3H-4, 11,5	16.51	0.744	3.806	0.91	3.403	2.071	1.392	0.507	0.825	825
3H-4, 14.5	16.55	0.747	3.806	0.82	3.121	1.742	1.379	0.522	0.856	856
3H-4, 20.6	16.61	0.749	3.806	0.85	3.235	1.810	1.425	0.576	0.849	849
3H-4, 26.5	16.66	0.750	3.806	0.96	3.654	2.192	1.462	0.544	0.918	918
3H-4, 41.5	16.82	0.751	3.806	0.88	3.549	2.107	1.185	0.560	0.823	825
3H-4, 44.6	16.85	0.755	3.806	0.94	3.578	2.203	1.374	0.659	0.715	715
3H-4, 50.6	16.91	0.757	3.806	0.94	3.578	2.514	1.064			
3H-4, 56.5	16.97	0.758	3.806	0.94	3.578	2.550	1.027			
3H-4, 65.5	17.06	0.761	3.806	0.96	3.654	2.626	1.027	0.447	0.580	580
3H-4, 71.7	17.12	0.762	3.806	0.88	3.349	2.391	0.959	0.411	0.547	547
3H-4, 74.6	17.15	0.763	3.806	0.84	3.197	2.013	1.185	0.495	0.690	690
3H-4, 80.5 3H-4, 86, 5	17.20	0.764	3.806	0.93	3.540	2.206	1.334	0.505	0.829	829
3H-4, 89.5	17.30	0.767	3.806	0.97	3.692	1.758	1.933	0.777	1.157	1157
3H-4, 95.5	17.35	0.768	3.806	0.90	3.425	1.522	1.904	0.684	1.219	1219
3H-4, 101.6	17.42	0.770	3.806	1.06	4.034	1.963	2.072	0.734	1.338	1338
3H-4, 104.7	17.50	0.772	3.806	0.98	3.750	2 317	1 451	0.547	0.999	999
3H-4, 116.7	17.57	0.774	3.806	1.13	4.301	3.051	1.250	0.100	0.777	
3H-4, 119.5	17.59	0.775	3.806	1.13	4.301	3.118	1.183	0.551	0.632	632
3H-4, 125.5 3H-4, 131.5	17.05	0.778	3.806	1.03	3.920	2.762	1.158	0.565	0.593	593
3H-4, 134.6	17.75	0.779	3.806	0.93	3.540	2.572	0.968	0.480	0.488	488
3H-4, 140.5	17.81	0.781	1.429	0.88	1.258	0.857	0.401	0.181	0.220	220
3H-4, 146.5	17.86	0.784	1.429	0.99	1.415	0.921	0.494	0.191	0.303	303
3H-5, 5.5 3H-5, 11, 5	17.95	0.790	1.429	0.94	1.343	0.778	0.565	0.259	0 293	203
3H-5, 14.5	18.05	0.798	1.429	0.98	1.400	1.107	0.294	0.403	0.275	270
3H-5, 20.5	18.10	0.801	1.429	0.98	1.400	1.071	0.330	2012/02/2012	1000000	0-0222
3H-5, 26.5	18.16	0.805	1.429	0.94	1.343	1.019	0.324	0.145	0.179	179
3H-5, 35,5	18.25	0.807	1.429	0.98	1.400	0.805	0.389			
3H-5, 41.5	18.32	0.816	1.429	0.94	1.343	0.877	0.466	0.199	0.268	268
3H-5, 44.6	18.35	0.818	1.429	0.86	1.229	0.798	0.431	11		
3H-5, 50.5 3H-5, 56, 5	18.40	0.822	1.429	0.90	1.286	0.881	0.405			
3H-5, 59.5	18.49	0.828	1.429	1.01	1.443	0.920	0.449			
3H-5, 65.5	18.56	0.833	1.429	0.91	1.300	0.929	0.372			
3H-5, 71.5	18.61	0.837	1.429	1.01	1.443	1.099	0.345	0.150	0.195	195
3H-5, 74.5 3H-5, 80 5	18.70	0.840	1.429	0.90	1.372	0.957	0.345			
3H-5, 86.6	18.77	0.848	1.429	0.92	1.315	0.945	0.370			
3H-5, 89.5	18.80	0.850	1.429	0.93	1.329	0.960	0.369			
3H-5, 95.5 3H-5, 101 5	18.85	0.854	1.429	0.91	1.300	0.942	0.358	0.193	0.226	336
3H-5, 104.5	18.94	0.856	1.429	0.90	1.286	0.708	0.518	0.182	0.330	550
3H-5, 110.7	19.01	0.865	1.429	0.93	1.329	0.739	0.590			
3H-5, 116.6	19.07	0.869	1.429	0.89	1.272	0.675	0.596			
311-5, 119.5	19.09	0.870	1.429	0.94	1.343	0.778	0.565			

Core, section,	Depth	Age	LSR	DBD	MAR	CAR	NCAR	OMAR	$TMAR^1$	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(c	$m^2 \times k.y$	.) <sup>-1</sup>		$mg(cm^2 \times k.y.)^{-1}$
3H-5, 125.5	19.15	0.874	1.429	0.83	1.186	0.673	0.513			
3H-5, 128.6	19.19	0.877	1.429	0.82	1.172	0.686	0.486	0.174	0.312	312
3H-5, 134.5	19.24	0.881	1.429	0.83	1.186	0.713	0.473			
3H-6, 5, 5	19.40	0.892	1.429	0.95	1.329	0.753	0.322			
3H-6, 11.5	19.51	0.900	1.429	0.96	1.372	0.902	0.470	0.440	0.030	30
3H-6, 14.5	19.55	0.902	1.429	0.98	1.400	0.853	0.547			
3H-6, 20.6 3H-6, 26, 5	19.61	0.907	1.429	0.84	1.200	0.678	0.522			
3H-6, 29.5	19.69	0.912	1.429	0.83	1.186	0.873	0.313			
3H-6, 35.5	19.75	0.916	1.429	0.98	1.400	1.056	0.344			10000
3H-6, 41.7	19.82	0.921	1.429	1.03	1.472	1.057	0.415	0.173	0.242	242
3H-6, 50 5	19.84	0.923	1.429	0.91	1.300	1.003	0.372			
3H-6, 56.5	19.97	0.932	1.429	0.96	1.372	0.879	0.492			
3H-6, 59.5	19.99	0.933	1.429	0.93	1.329	0.793	0.536			
3H-6, 65.5 3H-6, 71.5	20.06	0.938	1,429	0.87	1.243	0.748	0.496	0 179	0 202	202
3H-6, 74.5	20.15	0.945	1.429	0.86	1.229	0.766	0.463	0.177	0.272	272
3H-6, 80.5	20.20	0.948	1.429	0.93	1.329	0.963	0.366			
3H-6, 86.5	20.26	0.952	1.429	0.99	1.415	1.164	0.250			
3H-6, 95,5	20.30	0.955	1.429	1.04	1.480	1.242	0.243			
3H-6, 101.5	20.41	0.963	1.429	0.96	1.372	1.162	0.210	0.111	0.099	99
3H-6, 104.6	20.45	0.965	1.429	1.09	1.558	1.229	0.328			
3H-6, 116.6	20.50	0.969	1.429	0.95	1.343	0.900	0.429			
3H-6, 119.5	20.59	0.975	1.429	0.84	1.200	0.766	0.434	0.143	0.291	291
3H-6, 125.5	20.65	0.979	1.429	0.96	1.372	0.939	0.433			
3H-6, 131.5 3H-6, 134.7	20.72	0.984	1.429	1.07	1.529	0.958	0.571			
3H-6, 140.7	20.81	0.993	0.348	0.96	0.334	0.279	0.055	0.030	0.025	25
3H-6, 146.5	20.86	1.007	0.348	0.90	0.313	0.239	0.075	1200	E 222	1222
3H-7, 5.5	20.95	1.033	0.348	0.93	0.324	0.212	0.112	0.041	0.071	71
3H-7, 20.5	21.10	1.076	0.348	0.98	0.363	0.203	0.070			
3H-7, 26.5	21.16	1.093	0.348	0.98	0.341	0.253	0.088	0.036	0.052	52
3H-7, 29.5	21.19	1.102	0.348	0.92	0.320	0.248	0.072			
3H-7, 41.5	21.23	1.119	0.348	0.90	0.315	0.195	0.122			
3H-7, 44.6	21.35	1.148	0.348	0.91	0.317	0.215	0.101	0.034	0.068	68
3H-7, 50.5	21.40	1.162	0.348	0.88	0.306	0.201	0.105			
4H-1, 5.5 4H-1, 11.5	21.45	1.1//	0.348	0.90	0.313	0.160	0.147			
4H-1, 14.5	21.55	1.206	0.348	0.90	0.313	0.191	0.122	0.050	0.072	72
4H-1, 20.5	21.60	1.220	0.348	0.87	0.303	0.163	0.140			
4H-1, 26.5	21.66	1.237	0.348	0.91	0.317	0.189	0.128			
4H-1, 35.6	21.76	1.241	1.461	0.85	1.242	0.892	0.350		0.390	390
4H-1, 41.7	21.82	1.250	1.461	0.99	1.446	0.906	0.540			
4H-1, 44.7	21.85	1.252	1.461	0.94	1.373	1.076	0.297			
4H-1, 56.5	21.97	1.260	1.461	0.90	1.446	1.053	0.394	0.203	0.190	190
4H-1, 59.5	21.99	1.262	1.461	0.96	1.403	0.965	0.438			
4H-1, 65.5	22.06	1.267	1.461	0.99	1.446	0.899	0.547			
4H-1, 74.6	22.12	1.271	1.461	0.87	1.271	0.744	0.542	0.258	0.284	284
4H-1, 80.6	22.21	1.277	1.461	0.99	1.446	1.208	0.238			
4H-1, 86.6	22.27	1.281	1.461	0.99	1.446	1.223	0.223			
4H-1, 89.5 4H-1, 95.5	22.30	1.285	1.461	0.99	1.534	1.322	0.212			
4H-1, 101.5	22.41	1.291	1.461	0.91	1.330	1.113	0.217	0.134	0.083	83
4H-1, 104.6	22.45	1.293	1.461	0.98	1.432	1.179	0.253			
4H-1, 110.6 4H-1, 116.5	22.51	1.298	1.461	0.94	1.373	0.969	0.273			
4H-1, 119.5	22.59	1.303	1.461	0.91	1.330	1.023	0.307			
4H-1, 125.5	22.65	1.307	1.461	1.09	1.592	1.220	0.372			
4H-1, 131.5	22.72	1.312	1.461	1.03	1.505	1.137	0.368	0.162	0.206	206
4H-1, 140.5	22.81	1.313	1.461	0.98	1.432	0.998	0.337			
4H-1, 146.5	22.86	1.321	1.461	0.91	1.330	0.872	0.458	0.187	0.271	271
4H-1, 149.5	22.90	1.324	1.461	0.94	1.373	0.841	0.532			
4H-2, 11.5	23.01	1.320	1.461	1.07	1.303	1.258	0.323			
4H-2, 14.5	23.05	1.334	1.461	1.07	1.563	1.187	0.377	0.058	0.318	318
4H-2, 20.5	23.10	1.338	1.461	1.03	1.505	0.946	0.559			
4H-2, 20.0 4H-2, 29.5	23.19	1.343	1.461	1.05	1.534	1 417	0.922			
4H-2, 35.5	23.25	1.348	1.461	1.01	1.476	0.810	0.665			
4H-2, 41.6	23.32	1.353	1.461	0.94	1.373	1.077	0.296	0.162	0.134	134
4H-2, 44.5 4H-2, 50.5	23.34	1.354	1.461	1.01	1.476	1.174	0.301	0.163	0.138	138
4H-2, 56.6	23.47	1.363	1.461	0.96	1.403	1.083	0.320	0.172	0.148	148
4H-2, 59.5	23.49	1.365	1.461	0.99	1.446	1.147	0.300	0.131	0.168	168
4H-2, 65.5 4H-2, 71.7	23.56	1.369	1.461	0.94	1.373	1.021	0.353	0.193	0.160	160
4H-2, 74.7	23.65	1.375	1.461	0.95	1.403	1.060	0.343	0.167	0.176	176
4H-2, 80.5	23.70	1.379	1.461	0.98	1.432	0.935	0.497	0.213	0.284	284
4H-2, 86.5	23.76	1.383	1.461	0.94	1.373	0.741	0.632			

Core. section	Denth	Age	LSR	DBD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(c	$cm^2 \times k.y$	)^-1		$mg(cm^2 \times k.y.)^{-1}$
4H-2, 89.5	23.80	1.386	1.461	0.98	1.432	1.015	0.417	0.259	0.158	158
4H-2, 95.6 4H-2, 101.5	23.80	1.390	1.461	0.95	1.388	1.153	0.235	0.140	0.095	95 70
4H-2, 104.5	23.94	1.395	1.461	0.99	1.446	1.232	0.214	0.143	0.071	71
4H-2, 110.6	24.01	1.400	1.461	1.07	1.563	1.325	0.238	0.149	0.089	89
4H-2, 110.5 4H-2, 119.5	24.07	1.404	1.461	1.11	1.622	1.349	0.273	0.176	0.097	107
4H-2, 125.5	24.15	1.410	1.461	0.90	1.315	1.113	0.202	0.110	0.092	92
4H-2, 131.5	24.22	1.415	1.461	0.99	1.446	1.210	0.236	0.147	0.090	90
4H-2, 134.5 4H-2, 140.5	24.24	1.416	1.461	1.01	1.476	1.235	0.240	0.145	0.095	95
4H-2, 146.6	24.31	1.421	1.461	1.15	1.476	1.375	0.276	0.182	0.133	133
4H-2, 149.5	24.40	1.427	1.461	1.03	1.505	1.156	0.349	0.148	0.200	200
4H-3, 5.6	24.46	1.431	1.461	0.91	1.330	0.956	0.373	0.141	0.232	232
4H-3, 14.6	24.51	1.434	1.461	0.94	1.375	0.815	0.505	0.200	0.305	305
4H-3, 20.5	24.60	1.441	1.461	0.98	1.432	0.907	0.525	0.122	0.403	403
4H-3, 26.5	24.66	1.445	1.461	0.90	1.315	0.737	0.578	0.199	0.379	379
4H-3, 35.7	24.09	1.447	1.461	0.90	1.315	0.766	0.549	0.202	0.340	323
4H-3, 41.6	24.82	1.456	1.461	0.96	1.403	0.811	0.592	0.203	0.389	389
4H-3, 44.5	24.84	1.457	2.457	0.90	2.211	1.575	0.636	0.358	0.278	278
4H-3, 50.5 4H-3, 56.5	24.90	1.459	2.457	0.94	2.310	1.811	0.499	0.244	0.255	308
4H-3, 59.5	24.99	1.463	2.457	1.03	2.531	2.070	0.461	0.263	0.198	198
4H-3, 65.5	25.06	1.466	2.457	0.98	2.408	1.951	0.457	0.292	0.165	165
4H-3, 71.5 4H-3, 74.6	25.11	1.468	2.457	0.94	2.310	1.815	0.495	0.327	0.167	167
4H-3, 80.6	25.21	1.472	2.457	0.99	2.432	2.071	0.361	0.198	0.163	163
4H-3, 86.5	25.26	1.474	2.457	0.94	2.310	1.937	0.373	0.248	0.125	125
4H-3, 89.6 4H-3, 95.6	25.30	1.476	2.457	0.87	2.138	1.778	0.360	0.215	0.145	145
4H-3, 101.5	25.41	1.480	2.457	0.90	2.162	1.724	0.302	0.223	0.198	198
4H-3, 104.5	25.44	1.481	2.457	0.91	2.236	1.622	0.614	0.269	0.345	345
4H-3, 110.5	25.50	1.484	2.457	0.87	2.138	1.469	0.669	0.280	0.389	389
4H-3, 119.5	25.59	1.487	2.457	0.88	2.102	1.288	0.874	0.379	0.495	617
4H-3, 125.5	25.65	1.490	2.457	0.98	2.408	1.467	0.941	0.263	0.678	678
4H-3, 134.5	25.74	1.494	2.457	1.25	3.071	1.731	1.340	0.339	1.001	1001
4H-3, 140.7 4H-3, 146.5	25.81	1.497	2.457	0.91	2.236	1.290	0.946	0.322	0.623	023 464
4H-3, 149.7	25.90	1.500	2.457	1.01	2.482	1.857	0.624	0.393	0.232	232
4H-4, 5.5	25.95	1.502	2.457	0.94	2.310	1.842	0.468	0.384	0.083	83
4H-4, 11.5 4H-4, 14.5	26.01	1.505	2.457	0.98	2.408	1.964	0.444	0.275	0.168	168
4H-4, 20.5	26.10	1.508	2.457	0.94	2.310	1.915	0.394	0.247	0.147	147
4H-4, 26.5	26.16	1.511	2.457	0.96	2.359	1.930	0.429	0.070	0.170	1.52
4H-4, 29.6 4H-4, 35.5	26.20	1.512	2.457	0.87	2.138	1.722	0.415	0.262	0.153	153
4H-4, 41.5	26.32	1.517	2.457	0.88	2.162	1.759	0.403	0.243	0.161	161
4H-4, 44.5	26.34	1.518	2.457	0.96	2.359	1.846	0.513	0.299	0.214	214
4H-4, 50.5 4H-4, 56.6	26.40	1.520	2.457	0.88	2.162	1.666	0.496	0.308	0.188	188
4H-4, 59.6	26.50	1.525	2.457	0.88	2.162	1.418	0.744	0.394	0.350	350
4H-4, 65.5	26.56	1.527	2.457	0.88	2.162	1.380	0.782	0.414	0.368	368
4H-4, 71.0 4H-4 74 5	26.62	1.529	2.457	0.79	1.941	1.189	0.752	0.340	0.412	412 524
4H-4, 80.5	26.70	1.533	2.457	0.94	2.310	1.282	1.027	0.046	0.024	021
4H-4, 86.5	26.76	1.535	2.457	0.93	2.285	1.496	0.789	0.203	0.586	586
4H-4, 89.5 4H-4, 95.7	26.80	1.537	2.457	0.98	2.408	1.464	0.944	0.587	0.357	357
4H-4, 101.6	26.92	1.542	2.457	0.87	2.138	1.208	0.930	0.395	0.535	535
4H-4, 104.5	26.94	1.543	2.457	0.87	2.138	1.272	0.866	0.320	0.546	546
4H-4, 110.5 4H-4, 116.6	27.00	1.545	2.457	0.91	2.236	1.213	1.023	0.500	0.523	523
4H-4, 119.5	27.09	1.549	2.457	0.99	2.432	1.867	0.565	0.349	0.216	216
4H-4, 125.7	27.16	1.551	2.457	0.90	2.211	1.769	0.442	0.237	0.204	204
4H-4, 131.6 4H-4, 134.5	27.22	1.554	2.457	0.99	2.432	1.819	0.613	0.290	0.323	323
4H-4, 140.6	27.31	1.558	2.457	0.90	2.236	1.677	0.559	0.342	0.217	217
4H-4, 146.5	27.36	1.560	2.457	0.91	2.236	1.756	0.480	0.250	0.230	230
4H-4, 149.5	27.40	1.561	2.457	0.93	2.285	1.756	0.529			
4H-5, 11.5	27.51	1.566	2.457	0.87	2.138	1.021	0.516			
4H-5, 14.6	27.55	1.567	2.457	0.94	2.310	1.802	0.508	0.238	0.269	269
4H-5, 20.6	27.61	1.570	2.457	1.01	2.482	1.962	0.519			
4H-5, 20.5	27.69	1.572	2.457	1.03	2.531	1.903	0.628	0.272	0.502	503
4H-5, 35.5	27.75	1.575	2.457	0.88	2.162	1.383	0.779			202924 
4H-5, 41.5	27.82	1.578	2.457	0.88	2.162	1.115	1.048	0.400	0.647	647
4H-5, 44.5 4H-5, 50.6	27.84	1.579	2.457	0.87	2.138	1.151	0.987			
4H-5, 56.6	27.97	1.584	2.457	0.93	2.285	1.611	0.674	0.337	0.336	336
4H-5, 59.5	27.99	1.585	2.457	0.91	2.236	1.321	0.915			
4H-5, 65.5 4H-5, 71.6	28.06	1.588	2.457	0.96	2.359	1.693	0.666			
4H-5, 74.6	28.15	1.592	2.457	0.95	2.334	1.808	0.527	0.216	0.311	311
4H-5, 80.7	28.21	1.594	2.457	0.88	2.162	1.676	0.486	0.000 A.M.O	1	THE REAL PROPERTY AND A DESCRIPTION OF A DE
4H-5, 86.6	28.27	1.597	2.457	1.04	2.555	1.912	0.643			

APPENDIX C (continued).

#### TMAR<sup>2</sup> MAR CAR NCAR OMAR TMAR<sup>1</sup> Core, section, Depth Age LSR DBD $mg(cm^2 \times k.y.)^{-1}$ $g(cm^2 \times k.y.)^{-1}$ interval (cm) (mbsf) (Ma) (cm/k.v.) $(g/cm^3)$ 4H-5, 89.5 28.30 0.063 0.054 0.038 1.629 0.85 0.016 4H-5, 95.7 28.36 1.723 0.063 0.99 0.062 0.043 0.019 4H-5, 101.6 0.042 0.022 28.42 1.818 0.063 1.02 0.064 0.011 0.011 11 4H-5, 104.6 4H-5, 110.5 4H-5, 116.5 28.45 1.865 0.063 0.94 0.059 0.048 0.011 0.007 0.005 5 28.50 1.944 0.99 0.063 0.062 0.050 0.012 28.57 1.964 1.468 0.314 0.94 1.380 1.066 4H-5, 119.5 4H-5, 125.5 0.425 28 59 1 965 1.468 0.90 1.321 0.896 28.65 1.969 0.948 0.212 212 1.468 0.91 0.176 1.336 4H-5, 131.5 28.72 1.974 1.468 0.90 1.321 0.825 0.496 4H-5, 134.5 28 74 1 976 1.468 0.93 1.365 0.710 0.655 0.280 0.375 375 4H-5, 140.5 28.81 1.980 1.468 1.067 0.416 1.01 1.483 0.96 4H-5, 146.5 28.86 1.984 1.468 1.409 0.978 0.431 4H-5, 149.5 28.90 1.987 1.468 0.99 1.453 1.238 0.215 0.097 0.118 118 4H-6, 5.5 28.95 1.990 1.468 1.01 1.483 1.275 0.207 4H-6, 11.5 4H-6, 14.5 20.01 1 994 1.468 0.99 1.453 1.266 0.187 29.05 1.997 1.468 0.93 0.179 1.365 1.186 4H-6, 20.5 29.10 2.000 1.468 0.93 1.365 1.188 0.177 4H-6, 26.6 4H-6, 29.5 29.17 2.005 1.468 0.92 1.351 1.178 0.172 0.179 29.19 2.006 1.468 1.157 0.110 0.068 68 0.91 1.336 4H-6, 35.5 29.25 2.010 1.468 0.93 0.246 1.365 1.119 4H-6, 41.5 29.32 2015 1.468 0.93 1.365 1 030 0.335 4H-6, 44.5 0.351 29.34 2.016 1.468 0.94 1.380 1.029 0.181 0.171 171 4H-6, 50.6 4H-6, 56.5 29.41 2.021 1.468 0.96 1.409 0.391 1.018 29.47 2.025 1.468 0.96 1.409 1.097 0.312 4H-6, 59.6 29.50 1.571 0.303 2.027 1.468 1.07 1.267 0.139 0.164 164 4H-6, 65.6 4H-6, 71.5 20 56 2.032 1.468 0.96 1.409 1.074 0.335 29.61 0.291 0.212 212 2.035 1.468 0.94 1.380 0.877 0.503 4H-6, 74.7 29.65 2.038 1.468 1.318 0.311 1.11 1.629 4H-6, 80.5 4H-6, 86.5 29.70 2.041 1.468 1.01 1.483 1.220 0.262 0.239 29.76 2.045 1.391 1.468 1.11 1.629 4H-6, 89.5 29.80 2.048 1.468 1.571 1.373 0.197 0.124 0.074 74 1.07 4H-6.95.5 29.85 2 051 1 468 1.09 1 600 1.384 0.216 4H-6, 101.5 0.250 29.91 2.055 1.468 1.380 1.11 1.629 4H-6, 104.5 29.94 2.057 1.468 1.07 1.571 1.367 0.204 4H-6, 110.5 30.00 2.062 1 468 0.98 1 439 1 206 0 233 4H-6, 116.7 1.468 0.96 0.251 30.07 2.066 1.409 1.158 4H-6, 119.5 30.09 2.068 1.468 1.13 1.659 1.305 0.353 4H-6, 125.5 30.15 0.188 0.126 2.0721.468 1.066 0.314 126 0.94 1.3804H-6, 131.5 30.22 2.076 1.468 0.979 0.298 0.87 1.277 4H-6, 134.5 4H-6, 140.6 30.24 2.078 1.468 1.01 1.483 1.205 0.278 0.245 30.31 1.076 2.083 1.468 0.90 1.321 4H-6, 146.5 30.36 2.086 1.468 0.96 1.409 0.966 0.444 1.439 1.483 5H-1.5.5 30.95 2.126 1.468 0.98 1.116 0.322 5H-1, 11.5 1.468 0.328 31.01 2.130 1.154 1.01 5H-1, 17.6 31.08 2.135 1.468 1.483 1.214 0.269 1.01 0.138 63 5H-1, 20.5 31.10 2.136 1 468 1 05 1 541 1 340 0.201 0.063 5H-1, 26.5 31.16 2.141 1.468 1.483 1.285 0.197 1.01 5H-1, 32.5 5H-1, 38.5 31.23 31.28 2.145 1.468 1.629 1.277 0.353 1.11 2.149 1.468 1.09 1.600 1.358 0.242 5H-1, 41.5 31.32 2.151 1.511 1.238 0.213 0.96 1.451 5H-1, 44.5 5H-1, 50.6 2.153 2.157 31.34 1.511 0.94 1.420 1.140 0.281 31.41 0.284 1.511 1.03 1.556 1.272 5H-1, 56.6 31.47 2.161 1.511 1.541 1.225 0.316 1.02 5H-1, 59.5 31.49 2.163 1.511 1.03 1.556 1.162 0.394 0.191 0.204 204 5H-1, 65.5 31.56 2.167 1.511 1.738 1.239 0.499 1.15 5H-1, 71.5 5H-1, 74.5 31.61 2.171 1.511 1.03 1.556 1.032 0.524 31.65 2.173 1.511 0.94 1.420 0.893 0.527 5H-1, 80.5 31.70 2.177 0.756 1.511 0.91 1.375 0.619 0.245 0.374 374 0.553 0.332 5H-1, 86.5 31.76 2.180 1.511 0.87 1.315 0.762 5H-1, 95.6 31.86 2.187 1.511 1.09 1.647 1.315 5H-1, 101.5 31.91 2.190 1.511 1.07 1.617 1.308 0.308 5H-1, 104.5 5H-1, 110.5 31.94 2 192 1.511 1.03 1.556 1.302 0.254 32.01 0.256 2.197 1.391 0.157 0.099 99 1.511 1.647 1.09 5H-1, 116.5 32.06 2.200 1.511 0.98 .481 1.244 0.237 5H-1, 119.5 32.10 2.203 1.511 1.04 1.571 1.238 0.334 5H-1, 125.5 32.15 2.206 1.511 1.556 1.177 0.379 1.03 2.211 2.212 5H-1, 131.5 32.22 1.511 0.96 1.451 1.031 0.419 32.24 0.195 0.225 225 5H-1, 134.5 1.511 1.01 1.526 1.106 0.420 5H-1, 140.5 32.31 2.217 1.511 1.526 0.315 1.212 1.01 5H-1, 146.5 32 36 2.220 2.223 1.511 1.07 1.617 1.298 0.319 5H-1, 149.5 5H-2, 5.5 32.40 0.956 1.511 0.88 1.330 0.374 1.556 32.46 2.227 1.511 0.385 1.03 1.172 5H-2, 11.5 5H-2, 14.5 0.410 0.380 32.52 2.231 1.511 1.05 1.587 1.177 32.55 2.233 0.206 174 0.174 1.511 0.99 1.496 1.116 5H-2, 20.5 5H-2, 29.5 32.61 2.237 1.511 1.03 .556 1.079 0.477 32.69 2 242 1 511 0.99 1 4 9 6 0.940 0 556 5H-2, 35.5 32.76 2.247 1.511 1.03 1.556 1.095 0.461 5H-2, 41.5 5H-2, 44.5 32.82 2.251 2.253 1.511 0.98 .481 1.021 0.460 32.85 0.511 0.178 0.333 333 1.511 0.88 1.330 0.819 5H-2, 50.5 32.91 2.257 0.259 1.511 1.01 1.526 1.267 5H-2, 56.5 5H-2, 59.5 32.97 2 260 1.511 0.96 1.451 1.178 0.273 2.263 33.00 1.511 1.209 0.994 0.215 0.80 5H-2, 65.5 2.266 0.294 33.06 1.511 1.05 1.587 1.293 5H-2, 71.5 5H-2, 74.5 33.12 33.15 2.270 2.272 1.511 1.01 1.526 1.254 0.272 1.349 0.298 0.169 129 1.647 0.129 1.09 5H-2, 80.5 33.21 2.276 1.511 0.96 1.451 1.172 0.279 5H-2, 86.7 33.27 2 280 1.511 1.03 1.556 1.244 0.312 5H-2, 89.6 33.30 2.282 1.511 1.061 0.360 0.94 1.420

### APPENDIX C (continued).

5H-2, 95.5

33.36

2.286

1.511

0.94

1.420

0.998

0.422

25 27			5	Conserved in	MAR	CAR	NCAR	OMAR	TMAR	TMAR <sup>2</sup>
Core, section,	Depth (mbsf)	Age	LSR (cm/ky)	DBD	Mirtix	ola	$m^2 \times k v$	)-1	1111/111	$mg(cm^2 \times k.v.)^{-1}$
intervar (cm)	(most)	(Ma)	(CHUK.y.)	(g/cm <sup>-</sup> )		gu	лп ~ к.у			ing(cin (ki.j.)
5H-2, 101.7	33.42	2.290	1.511	0.91	1.375	0.858	0.517	0.050	0.205	205
5H-2, 104.6 5H-2, 110.6	33.45	2.292	1.511	0.91	1.375	0.822	0.555	0.258	0.295	295
5H-2, 119.5	33.60	2.302	1.511	1.07	1.617	1.318	0.299	0.198	0.102	102
5H-2, 125.5	33.66	2.306	1.511	1.13	1.707	1.431	0.276			
5H-2, 131.5 5H-2, 134.5	33.75	2.310	1.511	1.07	1.617	1.274	0.308			
5H-2, 143.5	33.84	2.318	1.511	0.93	1.405	1.101	0.305	0.172	0.133	133
5H-2, 149.6	33.90	2.322	1.511	0.98	1.481	1.085	0.395			
5H-3, 5.5	33.96	2.324	1.511	0.90	1.451	1.057	0.303			
5H-3, 11.5	34.02	2.330	1.511	0.88	1.330	0.957	0.373	0.004	0.100	102
5H-3, 14.5	34.05	2.332	1.511	0.80	1.209	0.811	0.398	0.206	0.192	192
5H-3, 26.5	34.17	2.330	1.511	0.91	1.375	1.061	0.314			
5H-3, 29.6	34.20	2.342	1.511	1.03	1.556	1.151	0.405	0.226	0.179	179
5H-3, 35.5	34.26	2.346	1.511	0.91	1.375	0.863	0.512			
5H-3, 44.5	34.32	2.350	1.511	1.07	1.617	0.592	1.025	0.597	0.428	428
5H-3, 50.5	34.41	2.356	1.511	0.99	1.496	1.203	0.293			
5H-3, 56.5	34.47	2.360	1.511	0.94	1.420	1.140	0.280			
5H-3, 65.5	34.56	2.362	1.511	1.03	1.451	1.135	0.266			
5H-3, 71.5	34.62	2.370	1.511	1.01	1.526	1.227	0.299		02100724/1	17127
5H-3, 74.6	34.65	2.372	1.511	0.98	1.481	1.151	0.330	0.179	0.151	151
5H-3, 80.5	34.77	2.370	1.511	0.99	1.496	1.130	0.350			
5H-3, 89.5	34.80	2.382	1.511	1.01	1.526	1.207	0.319			
5H-3, 95.5	34.86	2.386	1.511	0.94	1.420	1.144	0.276			
5H-3, 101.6 5H-3, 104.5	34.92	2.390	1.511	0.98	1,481	1.226	0.255	0.148	0.071	71
5H-3, 110.5	35.01	2.395	1.511	0.94	1.420	1.203	0.217	01110	01071	
5H-3, 116.7	35.07	2.399	1.511	1.05	1.587	1.346	0.241			
5H-3, 119.6 5H-3, 125.5	35.10	2.401	1.511	1.07	1.617	1.310	0.300			
5H-3, 131.5	35.22	2.409	1.511	0.96	1.451	0.940	0.511	0.235	0.276	276
5H-3, 134.5	35.25	2.411	1.511	1.03	1.556	1.170	0.386			
5H-3, 140.7 5H-3, 146, 5	35.31	2.415	1.511	0.98	1.481	1.072	0.408	0.237	0.289	289
5H-3, 149.5	35.40	2.421	1.511	1.09	1.647	1.014	0.633	01207	0.205	107
5H-4, 5.5	35.45	2.425	1.511	1.01	1.526	1.181	0.345	0.057	0.152	152
5H-4, 11.5 5H-4, 14, 5	35.51	2.429	1.511	0.96	1.451	1.041	0.410	0.257	0.155	155
5H-4, 20.5	35.60	2.431	1.511	1.01	1.526	1.265	0.261			
5H-4, 26.5	35.66	2.438	1.511	0.96	1.451	1.162	0.288	0.169	0.120	120
5H-4, 29.5 5H-4, 35.5	35.69	2.441	1.511	0.99	1.496	1.119	0.377			
5H-4, 41.5	35.81	2.448	1.511	0.90	1.420	0.989	0.433			
5H-4, 44.5	35.84	2.450	1.511	0.84	1.269	0.777	0.493	0.040	0.040	2/2
5H-4, 50.5 5H-4, 56,5	35.90	2.454	1.511	0.85	1.284	0.753	0.531	0.268	0.263	263
5H-4, 59.5	35.99	2.460	1.511	0.88	1.330	0.940	0.389			
5H-4, 65.5	36.05	2.464	1.511	0.84	1.269	1.058	0.211			
5H-4, 71.6	36.11	2.468	1.511	0.93	1.405	1.177	0.228	0 163	0.086	86
5H-4, 80.6	36.20	2.474	1.511	0.93	1.405	1.145	0.260	0.105	0.000	50
5H-4, 86.5	36.26	2.478	1.511	0.81	1.224	0.999	0.224			
5H-4, 89.5	36.29	2.480	1.511	0.91	1.375	1.103	0.272			
5H-4, 98.6	36.38	2.484	1.511	0.88	1.420	1.121	0.300			
5H-4, 107.5	36.47	2.492	1.511	0.87	1.315	0.868	0.446	0.223	0.223	223
5H-4, 110.5	36.50	2.494	1.511	0.91	1.375	0.980	0.395			
5H-4, 119.5	36.59	2.500	1.511	0.93	1.405	1.059	0.346			
5H-4, 125.6	36.65	2.504	1.511	0.90	1.360	0.962	0.398			
5H-4, 131.5	36.71	2.508	1.511	1.03	1.556	1.061	0.496			
5H-4, 140.5	36.80	2.510	1.511	0.87	1.375	0.595	0.780	0.399	0.381	381
5H-4, 146.5	36.86	2.518	1.511	0.88	1.330	0.655	0.675			
5H-4, 149.5	36.89	2.520	1.511	0.94	1.420	1.175	0.245			
5H-5, 11.5	37.02	2.529	1.511	1.05	1.587	1.366	0.220			
5H-5, 14.5	37.05	2.530	1.511	1.03	1.556	1.324	0.232	0.142	0.090	90
5H-5, 20.5	37.11	2.534	1.511	1.03	1.556	1.304	0.252			
5H-5, 20.5	37.19	2.539	1.511	1.01	1.526	1.220	0.307			
5H-5, 35.5	37.26	2.544	1.511	0.99	1.496	1.159	0.337			
5H-5, 41.5	37.32	2.548	1.511	0.99	1.496	1.197	0.299			
5H-5, 50.5	37.41	2.550	1.511	0.99	1.496	1,103	0.393	0.230	0.163	163
5H-5, 56.5	37.47	2.558	1.511	0.96	1.451	0.976	0.475	ಜನಗಳು	14941206220	
5H-5, 59.5	37.50	2.560	1.511	0.88	1.330	0.783	0.547	0.605	0.426	426
5H-5, 68.5	37.59	2.562	1.511	0.96	1.420	0.905	0.515	0.065	0.420	420
5H-5, 74.5	37.65	2.570	1.511	0.98	1.481	1.160	0.320	0.128	0.193	193
5H-5, 80.6	37.71	2.574	1.511	0.99	1.496	1.202	0.294			
5H-5, 89.5	37.80	2.578	1.511	1.03	1.550	1.439	0.263			
5H-5, 95.5	37.86	2.584	1.511	1.05	1.587	1.346	0.240			

APPENDIX C (continued).

Core section	Depth	4.00	1 60	DPD	MAR	CAR	NCAR	OMAR	TMAR <sup>1</sup>	TMAR <sup>2</sup>
interval (cm)	(mbsf)	(Ma)	(cm/k.y.)	(g/cm <sup>3</sup> )		g(0	$cm^2 \times k.y$	.) <sup>-1</sup>		$mg(cm^2 \times k.y.)^{-1}$
5H-5, 101.5	37.92	2.588	1.511	1.07	1.617	1.336	0.280			
5H-5, 104.5	37.94	2.589	1.511	1.20	1.813	1.458	0.355	0.215	0 182	192
5H-5, 116.5	38.07	2.594	1.511	0.98	1.481	0.974	0.398	0.215	0.182	162
5H-5, 119.5	38.10	2.600	1.531	0.94	1.439	0.645	0.794	0.275	0.519	519
5H-5, 125.6	38.16	2.604	1.531	1.01	1.546	1.153	0.394			
5H-5, 131.5 5H-5, 134.5	38.22	2.608	1.531	0.93	1.424	0.999	0.303			
5H-5, 140.5	38.31	2.614	1.531	1.01	1.546	1.104	0.442	0.308	0.134	134
5H-5, 146.5	38.37	2.618	1.531	1.17	1.791	1.514	0.277			
5H-5, 149.5	38.40	2.620	1.531	1.05	1.608	1.369	0.239	0.099	0.120	120
5H-6, 11.5	38.52	2.627	1.531	1.15	1.761	1.435	0.326	0.077	0.120	120
5H-6, 17.5	38.58	2.631	1.531	0.98	1.500	1.149	0.352			
5H-6, 20.5	38.61	2.633	1.531	1.11	1.699	1.302	0.397			
5H-6, 29.5	38.69	2.638	1.531	0.96	1.340	0.964	0.480	0.185	0.321	321
5H-6, 35.5	38.76	2.643	1.531	0.85	1.301	0.828	0.473	12201220	0.000.0	
5H-6, 41.5	38.82	2.647	1.531	1.09	1.669	1.227	0.442			
5H-6, 44.5 5H-6, 50,5	38.85	2.649	1.531	1.18	1.807	1.541	0.265			
5H-6, 56.5	38.97	2.657	1.531	1.01	1.546	1.360	0.186	0.118	0.069	69
5H-6, 59.5	39.00	2.659	1.531	1.07	1.638	1.400	0.238			
5H-6, 65.5	39.06	2.663	1.531	1.03	1.577	1.320	0.257			
5H-6, 74.5	39.12	2.669	1.531	1.03	1.577	1.298	0.309			
5H-6, 80.6	39.21	2.672	1.531	1.01	1.546	1.310	0.236			
5H-6, 86.5	39.27	2.676	1.531	1.13	1.730	1.512	0.218			
5H-6, 89.6	39.30	2.678	1.531	1.07	1.638	1.408	0.230	0.164	0.075	75
5H-6, 101.5	39.30	2.686	1.531	1.07	1.638	1.395	0.239	0.104	0.075	15
5H-6, 107.5	39.48	2.690	1.531	0.96	1.470	1.212	0.258			
5H-6, 110.6	39.51	2.692	1.531	0.93	1.424	1.128	0.296			
5H-6, 110.5	39.57	2.696	1.531	1.15	1.761	1.338	0.423			
5H-6, 125.5	39.66	2.702	1.531	0.95	1.454	1.032	0.423	0.217	0.206	206
5H-6, 131.6	39.72	2.706	1.531	0.93	1.424	1.049	0.375			
5H-6, 134.5	39.75	2.708	1.531	0.88	1.347	0.985	0.363			
5H-6, 146.6	39.81	2.712	1.531	0.96	1.393	1.100	0.304	0.149	0.106	106
5H-7, 2.5	39.93	2.720	1.531	0.94	1.439	1.203	0.236			22.2.2
5H-7, 5.5	39.96	2.721	1.531	1.01	1.546	1.305	0.242			
5H-7, 11.5 5H-7, 14.5	40.02	2.725	1.531	0.94	1.540	1.300	0.246			
5H-7, 20.5	40.11	2.731	1.531	0.87	1.332	1.026	0.306	0.175	0.131	131
5H-7, 26.5	40.17	2.735	1.531	1.04	1.592	1.177	0.416			
5H-7, 29.5	40.19	2.737	1.531	0.94	1.439	0.989	0.450			
5H-7, 41.5	40.20	2.745	1.531	0.84	1.286	0.642	0.644	0.257	0.387	387
5H-7, 44.5	40.35	2.747	1.531	0.83	1.271	0.670	0.601			
5H-7, 50.5	40.41	2.751	1.531	0.94	1.439	0.965	0.475			
6H-1, 2.5	40.45	2.754	1.531	0.96	1.470	1.410	0.344			
6H-1, 11.5	40.52	2.758	1.531	0.98	1.500	1.221	0.280			
6H-1, 14.5	40.55	2.760	1.983	0.91	1.805	1.468	0.337			
6H-1, 20.5 6H-1, 26.5	40.61	2.765	1.983	0.93	1.844	1.292	0.552			
6H-1, 29.5	40.69	2.767	1.983	1.05	2.082	1.054	1.029			
6H-1, 35.5	40.76	2.771	1.983	1.01	2.003	1.347	0.656			
6H-1, 41.5	40.82	2.774	1.983	1.13	2.241	1.454	0.787			
6H-1, 50.5	40.85	2.775	1.983	1.11	2.201	1.800	0.355			
6H-1, 53.5	40.94	2.780	1.983	1.05	2.082	1.615	0.467			
6H-1, 59.5	41.00	2.783	1.983	1.09	2.161	1.852	0.309			
6H-1, 65.5	41.06	2.786	1.983	1.11	2.201	1.863	0.338			
6H-1, 74.5	41.12	2.790	1.983	1.09	2.101	1.797	0.325			
6H-1,80.6	41.21	2.793	1.983	1.01	2.003	1.679	0.324			
6H-1, 86.5	41.27	2.796	1.983	1.11	2.201	1.836	0.365			
6H-1, 89.6	41.30	2.798	1.983	0.96	1.904	1.055	0.467			
6H-1, 101.5	41.42	2.804	1.983	1.03	2.042	1.552	0.491			
6H-1, 104.5	41.44	2.805	1.983	1.09	2.161	1.555	0.607			
6H-1, 110.6	41.51	2.808	1.983	1.07	2.122	1.573	0.549			
6H-1, 119.5	41.60	2.813	1.983	1.15	2.161	1.580	0.581			
6H-1, 125.5	41.66	2.816	1.983	1.07	2.122	1.659	0.463			
6H-1, 131.5	41.72	2.819	1.983	1.20	2.380	1.861	0.519			
6H-1, 134.5	41.75	2.820	1.983	1.05	2.082	1.793	0.289			
6H-1, 146.5	41.87	2.827	1.983	1.05	2.082	1.768	0.314			
6H-1, 149.6	41.90	2.828	1.983	0.94	1.864	1.567	0.297			

Notes: Depth (mbsf) = sub-bottom depth in meters below seafloor; age (Ma) = millions of years ago; LSR = linear sedimentation rate (cm per thousand years); DBD = dry bulk density (g/cm<sup>2</sup>); MAR = mass (bulk) accumulation rate (grams of sediment per square cm per thousand years; same units for other rates); CAR = carbonate accumulation rate; NCAR = noncarbonate accumulation rate; OMAR = noncarbonate minus colian accumulation rate; TMAR<sup>1</sup> = terrigenous accumulation rate (lithogenic material <63 µm) in units of g(cm<sup>2</sup> × k.y.)<sup>-1</sup>; TMAR<sup>2</sup> = terrigenous accumulation rate in units of mg(cm<sup>2</sup> × k.y.)<sup>-1</sup>. The carbonate and noncarbonate weight percentages that are used to calculate accumulation rates are given in Rack et al. (1993).