

3. RADIOLARIANS FROM THE CASCADIA MARGIN, LEG 146¹

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ABSTRACT

Radiolarian abundances, recorded in core-catcher samples from all Leg 146 sites, are consistently rare. Detailed occurrences of 65 taxa were estimated in additional samples for Sites 889/890 and 892. Radiolarian events, calibrated with well-documented paleomagnetic reversals during Ocean Drilling Program Leg 145 (Morley and Nigrini, in press), offer a relatively good chronology for both locations. A continuous sequence of late Pliocene to Pleistocene age, with an average sedimentation rate of 9 cm/k.y., is recognized at Site 889 on the accretionary wedge off Vancouver Island. An intricate biostratigraphy, with several hiatuses and three stratigraphic reversals involving late Pliocene intervals underlying middle or late Miocene deposits, is proposed for Site 892 on the accretionary wedge under the Oregon continental slope. The occurrence of many species characteristic of upwelling environments records upwelling activity along the Oregon margin through most of the late Pliocene to early Pleistocene.

INTRODUCTION

Ocean Drilling Program (ODP) Leg 146 investigated tectonic dewatering in the accretionary wedge at the Cascadia continental margin. The sedimentary section resting on oceanic lithosphere is a thick accumulation (up to 3.5 km) of turbidites and hemipelagic deposits. In this sedimentological environment, radiolarian fractions isolated from the sediment bulk are usually very small (less than 1%), or often absent.

At all sites drilled during Leg 146 (Fig. 1), radiolarian debris is indeed very rare, but occurrence of some well-preserved stratigraphic

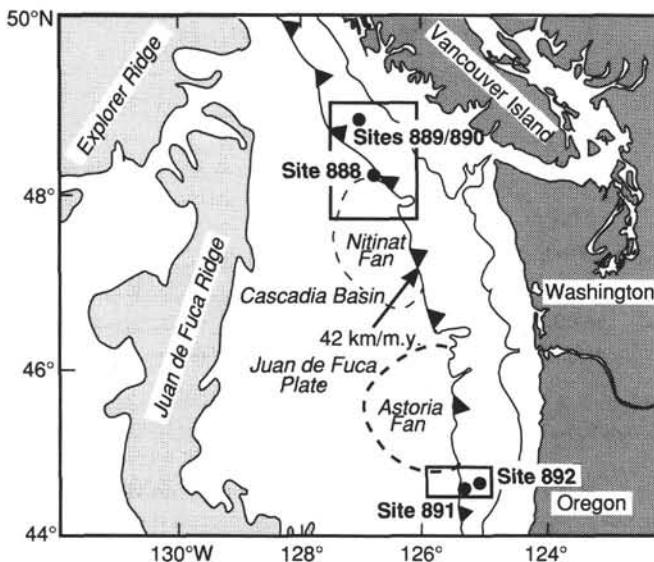


Figure 1. Location map of the Cascadia Margin, showing the convergent plate motion, major physiographic features, the position of the continental margin, and the position of drilling sites.

¹Carson, B., Westbrook, G.K., Musgrave, R.J., and Suess, E. (Eds.), 1995. *Proc. ODP, Sci. Results*, 146 (Pt. 1): College Station, TX (Ocean Drilling Program).

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markers in almost all holes allows a stratigraphic description of the different sedimentological sequences. Due to the scarcity of foraminiferal associations and a disturbed magnetostratigraphy, radiolarian biostratigraphy was very helpful during the cruise. After the cruise, shore-based studies of diatom assemblages have completed and confirmed the preliminary radiolarian biostratigraphy of the important Site 892, providing a good stratigraphic framework for the tectonic and geochemical synthesis of this site (see this volume).

The main purpose of this paper is to provide, using additional samples, a more refined biostratigraphy for Sites 889 and 892.

Radiolarian assemblages from Site 893 (Santa Barbara Basin) are not described here because the data they provide are mostly of paleoenvironmental importance and based on high-resolution counts of species assemblages that are beyond the scope of this report.

PROCEDURES

Sample preparation for microscopic examination during Leg 146 followed the standard techniques described by Sanfilippo et al. (1985). Samples were sieved first at 80 µm to eliminate clay aggregates. A second mesh of 50 µm was used for control. An additional cleaning of clay-rich sediments in a 5% ammonia solution was performed on samples prepared onshore.

For each sample examined, qualitative estimates of radiolarian abundance and preservation were made. Radiolarian assemblage abundance was assessed as follows:

- A = abundant (>500 specimens on slide)
- C = common (100–500 specimens on slide)
- F = few (50–100 specimens on slide)
- R = rare (<50 specimens on slide)
- VR = very rare (<10 specimens on slide)

Preservation of the radiolarian assemblage was based on the following:

- G (good) = radiolarians show no sign of dissolution with only minor fragmentation.
- M (moderate) = radiolarians show evidence of moderate dissolution with obvious fragmentation.
- P (poor) = radiolarians show signs of a high degree of dissolution with very few intact specimens.

Radiolarian Biostratigraphy

During the cruise itself, a radiolarian zonation with six zones was used to date the Pliocene/Pleistocene sediments (see "Explanatory Notes" chapter in Westbrook, Carson, Musgrave, et al., 1994).

Mainly constructed from the work of Hays (1970), Kling (1971), Foreman (1975), Riedel and Sanfilippo (1970, 1971, 1978), and Morley (1985), radiolarian zones were defined as follows:

- Botryostrobus aquilonaris* Zone, Hays (1970): the base is defined by the last occurrence of *Stylarctus universus*. This zone extends to the top of Recent sediments.
- Stylarctus universus* Zone, Hays (1970): the base is defined by the last occurrence (LO) of *Eucyrtidium matuyamai*.
- Eucyrtidium matuyamai* Zone, Hays (1970); emend. Foreman (1975): the base is defined by the evolutionary transition from *E. calvertense* to *E. matuyamai*.
- Lamprocrytis heteroporus* Zone, Hays (1970); emend. Foreman (1975): the base is defined by the LO of *Stichocorys peregrina*.
- Sphaeropyle langii* Zone, Foreman (1975): the base is defined by the FO of *S. langii*.
- Stichocorys peregrina* Zone, Riedel and Sanfilippo (1970); emend. Foreman (1975): the base is defined by the first evolutionary appearance of *S. peregrina*.

Using magnetostratigraphy, Morley and Nigrini (in press) have recently calculated good magnetic ages for many Neogene radiolarian events from the North Pacific. Radiolarian events occurring in Leg 146 sediments are now dated following the time chart established by Morley and Nigrini (in press) for the North Pacific radiolarian assemblages (Table 1).

As the last occurrence of *S. peregrina* (base of the *L. heteroporus* Zone) is now dated of 6.4 Ma for the northeast Pacific, the Pliocene/Pleistocene radiolarian zonation used in this paper is thus modified, and consists only of the four previous upper zones because the older *S. langii* and part of the *S. peregrina* Zones are now included in the *L. heteroporus* Zone (first occurrence of *S. langii* dated at 4.8 Ma in the northeast Pacific).

Table 1. Radiolarian events in the North Pacific (after Morley and Nigrini, in press).

Event	Species	Age (Ma)
LO	<i>Lychnocanoma nipponica sakai</i>	0.05
LO	<i>Stylacontarium acqilonium</i>	0.4
LO	<i>Stylarctus universus</i>	0.55
FO	<i>Lamprocrytis nigriniae</i>	0.8-1.2
LO	<i>Lamprocrytis neoheteroporus</i>	0.9
LO	<i>Eucyrtidium matuyamai</i>	1.0
LO	<i>Lamprocrytis heteroporus</i>	1.7
LO	<i>Sphaeropyle robusta</i>	1.5-1.7
FO	<i>Eucyrtidium matuyamai</i>	2.0
FO	<i>Lamprocrytis neoheteroporus</i>	2.6-2.8
FO	<i>Cycladophora davisianna</i>	2.9
FO	<i>Sphaeropyle langii</i>	4.8
LO	<i>Dictyophimus splendens</i>	4.8
LO	<i>Theocorys redondoensis</i>	4.9
LO	<i>Stichocorys peregrina</i>	6.4
FO	<i>Lamprocrytis heteroporus</i>	6.6
LO	<i>Stichocorys delmontensis</i>	6.8
Transition	<i>S. delmontensis-S. peregrina</i>	7.5
FO	<i>Stylacontarium acqilonium</i>	7.7
LO	<i>Cyrtocapsella japonica</i>	10.0
FO	<i>Dictyophimus splendens</i>	15.5
FO	<i>Theocorys redondoensis</i>	16.25

RADIOLARIAN BIOSTRATIGRAPHY OF LEG 146 SITES

Site 888 (Nitinat Fan)

An apparently continuous sequence of upper Pleistocene sediments was cored at Site 888 ($48^{\circ}10.00'N$ - $126^{\circ}39.79'E$). Radiolarians occurred in only a few intervals and in very low abundance.

Hole 888A, represented by a single core, recovered uppermost Quaternary sediments. An upper Pleistocene record (~0.7 Ma) was recovered from Hole 888B.

All core-catcher samples from Holes 888A and 888B were processed and examined for radiolarians. Rare and well-preserved radiolarians characteristic of the late Quaternary *B. aquilonaris* Zone (0.55 Ma to present) were found in sediments near the top of the sequence from Hole 888A (Sample 146-888A-1H-CC at 9.5 mbsf) and in Samples 146-888B-1H-CC through -18X-CC (5.4 to 162.5 mbsf) from Hole 888B. In this interval, four samples of coarse sand (Samples 146-888B-8H-CC, -9H-CC, -12H-CC, and -13H-CC; 72 to 118 mbsf) are barren. Samples 146-888B-19X-CC through -36H-CC (172 to 319.8 mbsf) are barren. Very rare and well-preserved Pleistocene radiolarians were found in Samples 146-888B-37H-CC through -41X-CC (329.7 to 357.9 mbsf). No stratigraphic markers were recognized. All core catchers downhole are barren, with the exception of Sample 146-888B-54X-CC (471.5 mbsf), where rare Pleistocene radiolarians were observed.

All specimens are typical of the Arctic Pleistocene radiolarian assemblage. Rare species indicative of upwelling areas (Nigrini and Caulet, 1992) are present in Samples 146-888A-1H-CC (9.5 mbsf) and 146-888B-42X-CC (367.2 mbsf).

Sites 889 and 890 (Accretionary wedge off Vancouver Island)

A sequence of upper Pliocene through upper Quaternary sediments was cored at Site 889 ($48^{\circ}41.95'N$ - $126^{\circ}52.09'W$), and an apparently continuous upper Quaternary sequence was cored at Site 890 ($48^{\circ}39.75'N$ - $126^{\circ}52.89'W$). Planktonic foraminifers, benthic foraminifers, and radiolarians are rare to abundant at these sites. Diatoms are common to abundant in radiolarian-bearing samples.

Occurrences of radiolarian species related to previously determined age data from Morley and Nigrini (in press) are recorded in Table 2 for Hole 889A.

Hole 889A

Well-preserved radiolarians are rare to abundant in Samples 146-889A-1H-CC (30.2 mbsf) to -37X-CC (275.4 mbsf). Samples 146-889A-38X-CC (284.2 mbsf) to -44X-CC (345.8 mbsf) are barren of radiolarians.

Samples 146-889A-1H-CC (30.2 mbsf) through 146-889A-8H-2, 96-101 cm (89.0 mbsf) contain a well-preserved assemblage dominated by *Cycladophora davisianna*, *Lychnocanoma nipponica sakai*, and *Pterocorys clausus*. This assemblage was previously correlated with the late Pleistocene *B. aquilonaris* Zone (0.55 Ma to present) (Westbrook, Carson, Musgrave, et al., 1994). Radiolarian species indicative of upwelling environments, such as *Lamprocrytis nigriniae*, *Pterocanum auritum*, and *Pterocanum grandiporus* (Nigrini and Caulet, 1992), were found in Samples 146-889A-2H-CC (39.9 mbsf), -3H-CC (50.0 mbsf), and -7H-CC (86.4 mbsf).

This interval was previously related to the *B. aquilonaris* Zone because the LO of *S. universus* was reported at 89.0 mbsf. Additional radiolarian events published by Morley and Nigrini (in press) help to define possible absolute ages for some parts of this interval:

Table 2. Range chart of radiolarian species at Hole 889A

Notes: see text for explanation of entries. Radiolarian ages were estimated using radiolarian events calibrated by Morley and Nigrini (in press).

Table 2 (continued).

- a. the occurrence of a specimen of *L. n. sakai* in Sample 146-889A-1H-CC (30.2 mbsf) gives an age older than 0.05 Ma to this sample;
- b. the LO of *Stylocontarium acqilonium* in Sample 146-889A-2H-6, 78–80 cm (37.78 mbsf) gives an age older than 0.4 Ma to this sample.

As the LO of *S. universus* (Table 1) is older than the LO of *S. acqilonium* (Morley and Nigrini, in press), we can consider the base of the *B. aquilonaris* Zone to be just below 37.78 mbsf. The question then becomes why no representatives of *S. universus* are present in most Pleistocene sediments from Hole 889A, particularly in well-preserved and diverse assemblages. Observations on shallow sites studied in the Antarctic (Caulet, 1991) and on the Peru Margin (De Wever et al., 1990) would indicate that debris of *S. universus* may be rare to absent at shallow depths, but no conclusive data are yet published.

The LO of *E. matuyamai* dates Sample 146-889A-9H-7, 50–52 cm (101.14 mbsf) at >1.0 Ma, in the late Pliocene/early Pleistocene *E. matuyamai* Zone (2.0 to 1.0 Ma). The interval between the LO of *S. acqilonium* and the LO of *E. matuyamai* can be considered as representing the *S. universus* Zone, and the hiatus previously deduced from the absence of *S. universus* in sediments above 89.38 mbsf can be now rejected. The FO of *L. nigriniae* (transitional forms) in Sample 146-889A-7H-4, 119–121 cm (81.86 mbsf) and the LO of *L. neoheteroporus* in Sample 146-889A-8H-2, 117–120 cm (89.18 mbsf) give an age of 0.8 to 0.9 Ma to this interval.

The first occurrence of *E. matuyamai* (at the base of the *E. matuyamai* Zone) is located between Sample 146-889A-25X-CC (200.7 mbsf) and Sample 146-889A-26X-CC (215.4 mbsf), giving an age of nearly 2.0 Ma to the sediments drilled at 206.80 mbsf. The LO of *L. heteroporus* (1.7 Ma) is recorded in Sample 146-889A-20X-CC (168.0 mbsf). Reworked specimens of early Pliocene species, such as *C. davisianna* (ancestral form), *S. peregrina*, and *Sphaeropyle robusta*, occur in many samples within this interval, suggesting erosional and transportation processes during the early Pleistocene. Radiolarian species indicative of upwelling environments, such as *P. grandiporus*, *L. nigriniae*, and *Phormostichoartus crustula*, were found in Samples 146-889A-10H-1, 140–144 cm (105.4 mbsf), 146-889A-11H-3, 88–92 cm (117.4 mbsf), and 146-889A-12H-4, 124–128 cm (123.9 mbsf).

Rare to few radiolarians assign the interval between Samples 146-889A-26X-CC (215.4 mbsf) and -37X-CC (284.10 mbsf) to the late Pliocene *L. heteroporus* Zone (6.4 to 2.0 Ma). The occurrence of rare representatives of *C. davisianna* assigns an age younger than 2.9 Ma to Sample 146-889A-37X-CC (284.10 mbsf). Rare representatives of the upwelling radiolarian marker *P. crustula* suggest an upwelling influence in sediments of Samples 146-889A-30X-CC and -31X-CC between 235 and 247 mbsf. No reworked fauna was observed.

Samples 146-889A-39X-CC through -44X-CC (294.0 to 338.7 mbsf) are barren of radiolarians.

Hole 889B

Sample 146-889B-3R-CC (218.2 mbsf) is barren of radiolarians. Rare to few representatives of the North Pacific radiolarian assemblage were found in Samples 146-889B-4R-CC and -5R-CC (228.2 and 235.2 mbsf). The occurrence of *L. heteroporus* and the absence of *E. matuyamai* suggest a late Pliocene age for this interval. Samples 146-889B-6R-CC through -8R-CC (248.2 to 269.5 mbsf) are barren of radiolarians. Rare radiolarians of late Pliocene age were observed in Samples 146-889B-9R-CC and -10R-CC (273.9 and 284.4 mbsf). Farther downhole in Hole 889B, all core-catcher samples contain rare radiolarians. No stratigraphic markers were recognized. Radiolarian species characteristic of upwelling environments, such as *Acrosphaera murrayana* and *P. crustula*, were found in Sample 146-889B-14R-CC (317.8 mbsf).

Hole 889D

Rare to few radiolarians were found in Samples 146-889D-1H-CC through -3X-CC (89.7 to 141.2 mbsf). No stratigraphic markers were recognized. Sample 146-889D-4N-CC (151.9 mbsf) is barren of radiolarians.

Holes 890A and 890B

All core-catcher samples from Site 890 were examined for radiolarians. Sample 146-890A-1H-CC (3.7 mbsf) contains a diverse radiolarian assemblage assigned to the Pleistocene *B. aquilonaris* Zone (0.55 Ma to present).

Few radiolarians, representative of the late Pleistocene *B. aquilonaris* Zone, were found in Sample 146-890B-1H-CC (7.3 mbsf). Samples 146-890B-2H-CC to -5H-CC (17.5 to 48.8 mbsf) are barren of radiolarians.

Site 891 (Oregon continental margin)

The sediments recovered at Site 891 (44°38.64'N–125°19.55'W) afford poor biostratigraphic control throughout most of the sequence. Factors contributing to the uncertainty in age determinations include: (1) the occurrence of several intervals barren of microfossils, (2) poor sedimentary recovery, and (3) the absence of zonal markers.

All core-catcher samples from Holes 891A and 891B were processed and examined for radiolarians. Rare to common and well-preserved radiolarians characteristic of the late Quaternary were found in sediments near the top of the sequence from Hole 891A (Samples 146-891A-1H-CC to -3H-CC, 0 to 9.5 mbsf). Radiolarian species indicative of upwelling environments (*L. nigriniae* and *P. crustula*) were found in Sample 146-891A-2H-CC, at 7.3 mbsf.

Only rare radiolarian debris was found in the upper part of Hole 891B (Samples 146-891A-1X-CC to -10X-CC, 0 to 83.6 mbsf), except in Sample 146-891B-7X-CC (56.2 mbsf) where two stratigraphic and/or upwelling markers (*L. nigriniae* and *P. auritum*) indicate a Pleistocene age. Samples 146-891B-10X-CC to -31X-CC (83.6 to 244.5 mbsf) are barren of radiolarians.

Radiolarians were very rare to rare in Samples 146-891B-32X-CC to -40X-CC (252.4 to 306.1 mbsf). No stratigraphic or paleoenvironmental markers were observed. The species composition of these sparse assemblages could be indicative of a Pleistocene age.

Samples 146-891B-42X-CC to -52X-CC (323.1 to 411.6 mbsf) are barren of radiolarians. Radiolarians occur rarely in the interval between Samples 146-891B-55X-CC and -57X-CC (440.1 to 455.3 mbsf). The occurrence of *C. davisianna* in Sample 146-891B-57X-CC (455.3 mbsf) indicates a Pliocene–Pleistocene age. Sample 146-891B-58X-CC (465.9 mbsf) is barren of radiolarians.

Site 892 (Accretionary wedge under the Oregon continental slope)

A discontinuous sequence of upper Neogene sediments was cored at Site 892 (44°40.45'N–125°7.13'W). Poorly to moderately preserved planktonic foraminifers and poorly to well-preserved benthic foraminifers and radiolarians are rare to common at this site. Diatoms are common to abundant in radiolarian-bearing samples.

Occurrences of radiolarian species are recorded in Table 3 for Hole 892A, and in Table 4 for Holes 892D and 892E. Radiolarian events and inferred radiolarian ages in Hole 892A are summarized in Table 5.

Hole 892A

Sample 146-892A-1X-1, 128–130 cm (1.29 mbsf), contains a well-preserved assemblage. The co-occurrence of *L. nigriniae* and well-preserved specimens of *L. neoheteroporus*, and the absence of

Table 3. Range chart of radiolarian species at Hole 892A.

Radiolarian ages (Ma)	Core, section, interval (cm)	Depth (mbst)	Abundance	Preservation																												
					<i>Anthocyrtidium diocenica</i>	<i>Bathyopera deflandrei</i>	<i>Bathystrobus aquilonaris</i>	<i>Bathystrobus aquilonaris</i> aff.	<i>Bathystrobus b. Bramlettei</i>	<i>Bathystrobus b. costatus</i>	<i>Bathystrobus b. seriatus</i>	<i>Ceratospyris borealis</i>	<i>Ceratocyrtis histricosa</i> gr.	<i>Cycladophora b. amphora</i>	<i>Cycladophora b. helios</i>	<i>Cycladophora davissiana</i>	<i>Cycladophora pliocenica</i>	<i>Cycladophora robusta</i>	<i>Cyriocapsella iaponica</i>	<i>Dicyophimus crisiae</i>	<i>Dicyophimus splendens</i>	<i>Encyrtidium calverense</i>	<i>Encyrtidium matuyamai</i>	<i>Gondwanaria dovieli</i>	<i>Halimeta miocenea</i>	<i>Lamprocyclas hawaii</i>	<i>Lamprocyclas maratensis</i>	<i>Lamprocyris heteroporus</i>	<i>Lamprocyris neoheteroporus</i>	<i>Lipmanella acanthica</i>	<i>Lychnocanum n. nipponica</i>	<i>Lychnocanum n. sakai</i>
1/0.9	1X-1, 128-130	1.29	F G	A																												
1.7/1.0	1X-2, 126-129	2.77	R G		F																											
	1X-3, 7-9	3.04	R G		R																											
	1X-CC	4.30	C G		C																											
2/1.7	2X-3, 52-54	13.02	R G		R																											
	2X-CC	13.50	F G																													
2.9/2	3X-1, 14-16	19.14	VR M																													
	3X-2, 28-30	20.78	R G		+ R																											
	3X-CC	23.20	F G																													
7.5/6.4	4X-1, 64-66	29.14	VR G																													
	4X-2, 26-30	29.80	VR P																													
	4X-CC	30.30	R G																													
	6X, TOP	39.00	R G																													
	6X-2, 93-95	41.43	R G	+ R																												
15.5/10	6X-CC	45.70	R G																													
	7X-1, 49-52	48.99	R G																													
>6.6	7X-2, 10-15	50.10	R M																													
2.9/2.8	7X-2, 41-43	50.43	R G																													
	7X-3, 30-33	51.82	no markers																													
	7X-4, 63-65	53.67	F G		+ R																											
	7X-6, 98-100	57.02	F G																													
	7X-CC	58.30	F G																													
	8X-2, 116-118	60.68	no markers																													
	8X-3, 24-26	61.27	C M																													
7.5/6.4	8X-4, 104-106	63.49	VR P																													
	8X-5, 38-40	64.33	no markers																													
	8X-CC	64.70	R M																													
2.9/2.8	9X-1, 42-45	67.92	R M																													
	9X-CC	68.60																														
	11X-1, 44-46	78.44	VR M																													
?	11X-3, 16-19	80.61	R M																													
	11X-CC	81.30	VR P																													
	12X-1, 15-18	87.65	no markers																													
	12X-CC	88.60	no markers																													
	13X-2, 36-38	97.65	barren																													
	13X-4, 55-57	99.33	VR P																													
	13X-7, 138-14	103.33	no markers																													
	13X-8, 91-93	104.36	no markers																													
	13X-CC	106.10	VR P																													
	14X-CC	107.00	no markers																													
?	15X-1, 42-44	116.42	no markers																													
	15X-CC	117.40	no markers																													
	16X-CC	126.30	no markers																													
	17X-CC	136.80	no markers																													
	18X-CC	146.90	VR P																													
<2.9	20X-1, 87-92	164.40	VR P																													
?	20X-1, 142-14	164.92	no markers																													
	20X-2, 44-46	165.46	no markers																													
	20X-CC	167.40	no markers																													
	21X-CC	173.10	no markers																													

Notes: see text for explanation of entries. Radiolarian ages were estimated using radiolarian events calibrated by Morley and Nigrini (in press).

Table 3 (continued).

	<i>Phormosiochaetus schneideli</i>	<i>Prunopyle titan</i> aff.	<i>Pseudodictyophimus gracilipes</i>	<i>Pierocanium auritum</i>	<i>Pierocanium grandiporus</i>	<i>Pierocanium korotevi</i>	<i>Pierocanium trilobatum</i>	<i>Pterocorys clausus</i>	<i>Siphocampe arachnea</i>	<i>Siphocampe modebensis</i>	<i>Sphaeropyle langii</i>	<i>Sphaeropyle robusta</i>	<i>Spiraea</i> spp.	<i>Spongodiscus osculatus</i>	<i>Spongotorchus glacialis</i>	<i>Stauropiphos communis</i>	<i>Stichocorys peregrina</i>	? <i>Stichocorys wolffi</i>	<i>Stylecontarium acquinum</i>	<i>Stylaccont. aff. S. hispiculum</i>	<i>Stylaractus universus</i>	<i>Theocorys redondensis</i>	<i>Theocorythium trachelium</i> old	
R		A + R	F R A A		R	R	R		R R									F	R	R	R	R		
	R	+ R	+ +		R		R		R R									R	R	R	R	R		
	R	+ +		+ R	R		R		R		R						R	R	R	R	R			
+	+ R	R	+							R	C						R	R	R	R	R			
				R					+ + + +		R						A	R	A	R	C +			
									+ + + +		R +						R	R	R	R	R			
+				R	+	R	C			R +	A						R	R	R	R	R			
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Table 4. Range chart of radiolarian species at Hole 892D and 892E.

Radiolarian ages (Ma)	Core, section, interval (cm)	Depth (mbf)	Abundance	Preservation	<i>Actinomma popofskii</i>	<i>Botryopera deflandrei</i>	<i>Botryostrobus aquilonaris</i>	<i>Botryostrobus aquilonaris aff.</i>	<i>Botryostrobus b. bramlettei</i>	<i>Botryostrobus b. costatus</i>	<i>Botryostrobus b. seriatus</i>	<i>Ceratospysis borealis</i>	<i>Ceratocystis histricosa</i> gr.	<i>Cycladophora b. amphora</i>	<i>Cycladophora b. helios</i>	<i>Cycladophora pilionenica</i>	<i>Cycladophora davisiiana</i>	<i>Cyrtocapsella japonica</i>	<i>Dicyrophilus crisiae</i>	<i>Dicyrophilus splendens</i>	<i>Eucyrtidium calvertense</i>	<i>Gondwanaria dozieri</i>	<i>Lamprocyclas hawaii</i>	<i>Lamprocyclas marquesensis</i>	<i>Lamprocyclis heteroporos</i>	<i>Lipmanella acanthica</i>	<i>Lychnocanium n. nipponica</i>	<i>Lychnocanium n. sakai</i>	<i>Phormostichoartus crustula</i>	<i>Phormostichoartus fistula</i>	<i>Phormostichoartus intermedius</i>	<i>Phromopyle titan</i>		
1.2 - 0.9	892D-1X, CC	8.50	A	G																														
	892D-2X-1, 6-9	8.56	C	G			R																											
	892D-2X-2, 6-8	10.06	R	G		R +																												
2 - 1	892D-2X, CC	18.00	R	G																														
	892D-3X-1, 5-7	18.05	R	G		C																												
	892D-3X, CC	27.50	R	G																													R	
2.9 - 2	892D-4X-1, 29-31	27.79	R	G																														
	892D-4X-3, 28-30	30.30	R	G		R																												
	892D-4X, CC	37.00	R	G																														R
	892D-5X-1, 52-54	37.52	R	G		R R																												
15.5 - 10	892D-5X-2, 95-97	39.45	R	G	C																												R	
	892D-5X-3, 125-127	41.25	R	G	C																													
	892D-5X, CC	46.50	R	G	R																													
	892D-6X-1, 83-86	47.33	R	G	R																													
	892D-6X-3, 91-94	50.16	R	G																														
	892D-6X-4, 74-77	51.49	VR	G																														
	892D-6X-5, 67-70	52.92	R	G																														
	892D-6X, CC	54.00	R	G																														
	892D-7X-1, 73-76	54.73	R	G	R	R																												
	892D-7X-2, 25-28	55.75	VR	G																														
	892D-7X-3, 74-77	57.50	R	G																														
2.9 - 2.8	892D-7X-4, 113-116	59.39	VR	G																														
	892D-7X-5, 32-34	60.08	VR	G																														
	reworked	892D-7X-6, 21-23	61.24	C	G																													
	892D-7X, CC	61.70	R	M																														
	892D-8X-1, 43-45	62.13	R	G																														
	892D-8X-2, 57-60	63.57	VR	G																														
	892D-8X-2, 117-119	64.17	VR	G																														
	892D-8X-3, 65-67	65.15	VR	G																														
?	892D-8X, CC	69.30	barren																															
	892D-9X, CC	77.60	VR	M																														
< 2.9 - 2.8	892D-10X-3, 45-48	102.70	barren																															
	892D-10X-6, 26-27	107.31	barren																															
	892D-10X, CC	109.50	barren																															
	892D-11X-1, 20-22	109.70	barren																															
< 2.9 - 2.8	892D-11X-2, 46-48	111.46	VR	M																														
	892D-11X, CC	119.00	VR	M																														
?	892D-12X-2, 91-93	121.41	no markers																															
	892D-12X-3, 81-83	122.81	VR	P																														
	892D-12X, CC	128.50	no markers																															
	892D-13X, CC	138.00	no markers																															
	892D-14X, CC	147.50	no markers																															
	892D-15X-1, 80-83	148.30	no markers																															
	892D-15X-2, 37-39	149.37	VR	P																														
	892D-15X-3, 77-79	121.27	no markers																															
	892D-16X-4, 106-108	162.09	no markers																															
	892D-16X-5, 53-55	163.06	no markers																															

?	892E-1X-2, 38-40	0.64	C G +																													
?	892E-3H-1, 20-22	33.22	no markers																													
?	892E-3H-3, 54-56	35.70	VR M +																													
?	892E-4H-4, 47-49	37.11	VR M																													
?	892E-4H-6, 04-106	39.83	no markers																													

Table 4 (continued).

Table 5. Radiolarian event chart for Hole 892A.

Core, section, interval (cm)	Depth (mbsf)	Age (Ma)	Events
892A-1X-1, 128–130	1.29	1/0.9	Co-occurrence of <i>L. nigriniae</i> and <i>L. neoheteroporos</i>
1X-2, 126–129, to	2.77		Occurrences of <i>E. matuyamai</i> , <i>P. pitomorphus</i> , and <i>L. neoheteroporos</i> ;
1X-CC	4.30	1.7/1	absence of <i>L. heteroporos</i>
2X-3, 52–54, to	13.02	2/1.7	Occurrences of <i>L. heteroporos</i> , <i>L. neoheteroporos</i> , and <i>E. matuyamai</i>
2X-CC	13.50		
3X-1, 14–16, to	19.14	2.9/2	Occurrences of <i>C. davisiana</i> , <i>L. neoheteroporos</i> ; absence of <i>E. matuyamai</i>
3X-CC	23.20		Occurrences of <i>L. neoheteroporos</i> and <i>L. heteroporos</i> ; no <i>C. davisiana</i>
4X-1, 64–66, to	29.14	7.5/6.4	Occurrences of <i>S. peregrina</i> and <i>D. splendens</i>
6X-2, 93–95	41.43		Occurrences of <i>T. redondoensis</i> and <i>A. pliocenica</i>
6X-CC, to	45.70	15.5/10	Occurrences of <i>C. japonica</i> , <i>D. splendens</i> , <i>P. vema</i> , <i>T. redondoensis</i> , and <i>B. bramlettei</i>
7X-1, 49–52	48.99		
7X-2, 10–15	50.10	< 6.6	Occurrences of <i>P. fistula</i> and <i>L. heteroporos</i>
7X-2, 41–43, to	50.43	2.9/2.8	Occurrence of <i>L. neoheteroporos</i> ; absence of <i>C. davisiana</i> ;
8X-3, 24–26	61.27		<i>S. peregrina</i> present
8X-4, 104–106, to	63.49	7.5/6.4	Occurrences of <i>S. peregrina</i> , <i>D. splendens</i> , and <i>T. redondoensis</i>
8X-CC	64.70		
9X-1, 42–45, to	67.92	2.9/2.8	Occurrence of common <i>L. neoheteroporos</i> ;
11X-1, 44–46	78.44		absence of <i>C. davisiana</i>
11X-3, 16–19, to	80.61	?	Occurrences of <i>L. heteroporos</i> and <i>S. peregrina</i> ;
11X-CC	81.30		absence of <i>L. neoheteroporos</i>
12X-1, 15–18, to	87.65	Barren	
892A-13X-2, 36–38	97.65		
13X-4, 55–57	99.33	?	Occurrence of <i>L. heteroporos</i>
13X-7, 138–140, to	103.33	Barren	
13X-8, 91–93	104.36		
13X-CC	106.10	?	Occurrence of <i>S. peregrina</i>
14X-CC, to	107.00	Barren	
17X-CC	136.80		
18X, CC	146.90	?	Occurrences of <i>H. miocenica</i> and <i>B. bramlettei</i>
20X-1, 87–92	164.40	< 2.9	Occurrence of <i>C. davisiana</i>
89A-20X-1, 142–144, to	164.92	Barren	
21X-CC	173.10		

E. matuyamai, place this interval in the base of the *S. universus* Zone between 0.9 and 1.0 Ma.

The interval between Sample 146-892A-1X-2, 126–129 cm, and Sample 146-892A-1X-CC (2.77 to 4.30 mbsf), contains representatives of *E. matuyamai* and *L. neoheteroporos*, but no forms such as *L. nigriniae* or *L. heteroporos*. Its age is 1.7 to 1.0 Ma between the LO of *L. heteroporos* and the LO of *E. matuyamai*. Radiolarian species indicative of upwelling environments, such as *P. auritum*, *P. grandiporus*, *P. crustula*, and *Phormostichoartus schneideri* (Nigrini and Caulet, 1992), were found in many samples from this interval.

The FO of *E. matuyamai* in Sample 146-892A-2X-CC (13.50 mbsf) assigns an age of 1.7 to 2 Ma to the upper interval between 13.02 and 13.50 mbsf.

Below this depth, the occurrence of *C. davisiana*, *L. neoheteroporos*, and *L. heteroporos* in Samples 146-892A-3-1, 14–16 cm, to -3X-CC (19.14 to 23.20 mbsf), places this interval in the upper *L. heteroporos* Zone (2.9 to 2 Ma). Radiolarian species, common in advected waters, testify to a continuous activity of the local upwelling systems.

Immediately below, at 29.14 mbsf (Sample 146-892A-4X-1, 64–66 cm), the composition of the radiolarian assemblage is quite different, with the occurrence of late Neogene forms such as *Cycladophora pliocenica* and *S. peregrina* and the absence of all other species characteristic of Arctic Pleistocene assemblages. The interval (29.14 to 41.3 mbsf) between this sample and Sample 146-892A-6X-2, 93–95 cm (41.43 mbsf), can be dated at 7.5 (FO of *S. peregrina*) to 6.4 Ma (LO of *S. peregrina*) because late Miocene forms such as *Botryopera deflandrei*, *Theocorys redondoensis*, and *Prunopyle titan* co-occur with *S. peregrina*. No typical representatives of *S. delmontensis* were found. Some close morphotypes reported as ?*Stichocorys wolffii* may be taken as local forms of *S. peregrina*, considering the great variability of this species in upwelling environments (see taxonomic remarks in the species list).

A hiatus spanning nearly 3 m.y., located between 23.20 and 29.14 mbsf, is inferred from these data.

The radiolarian assemblage observed in Samples 146-892A-6X-CC (45.70 mbsf) and 146-892A-7X-1, 49–52 cm (48.99 mbsf), is older because of common middle Miocene forms such as *Cyrt-*

capsella japonica. The co-occurrence of these species with representatives of *Dictyophimus splendens* and *T. redondoensis*, and the absence of typical specimens of *S. peregrina*, suggest an age of 15.3 to 10 Ma for this interval.

A second hiatus spanning at least 2 to 3 m.y. is located between 41.43 and 45.70 mbsf, above the middle Miocene sequence.

Rare and moderately preserved debris of *L. heteroporos* in Sample 146-892A-7X-2, 10–15 cm at 50.10 mbsf, assigns an age younger than 6.6 Ma to this level, suggesting a stratigraphic reversal above this sample.

The stratigraphic reversal below the middle to upper Miocene interval is confirmed by the occurrence of abundant, well-preserved representatives of *L. neoheteroporos* in the interval between Sample 146-892A-7X-2, 41–43 cm (50.43 mbsf), and Sample 146-892A-8X-3, 24–26 cm (61.27 mbsf). The absence of *C. davisianna* and *S. peregrina* gives an age of 2.8–2.9 Ma to this sequence. Many Arctic forms of late Pliocene to Pleistocene age and upwelling indicators are common. Assemblages are generally well-preserved, and species diversity is high. Rare reworked forms of late Miocene age can be observed.

Two meters down, an older assemblage can be observed in Sample 146-892A-8X-4, 104–106 cm (63.49 mbsf). The occurrence of *S. peregrina*, *D. splendens*, and *T. redondoensis* gives an age of 7.5 to 6.4 Ma to the interval between 63.49 and 64.70 mbsf (Sample 146-892A-8X-CC), suggesting a hiatus of probably 3 m.y. located between 61.27 and 63.49 mbsf.

Further below this upper Miocene sequence, a second reversal is inferred, as assemblages identical in composition to those observed between 50.43 and 61.27 mbsf are reported for the interval between Samples 146-892A-9X-1, 42–45 cm (67.92 mbsf) and -11X-1, 44–46 cm (78.44 mbsf). Common abundance of *L. neoheteroporos* and *S. acutum*, and the absence of *C. davisianna*, also suggest an age of 2.9–2.8 Ma for this interval. The reversal is located between 64.70 and 67.92 mbsf.

Downhole, radiolarians are very rare and poorly preserved when present. In the interval between Samples 146-892A-11X-3, 16–19 cm (80.61 mbsf) and -18X-CC (146.90 mbsf), rare representatives of *L. heteroporos*, *Stauroxiphos communis*, and *S. peregrina* suggest a late Miocene to early Pliocene age, but preservation is so poor and abundance so rare that no age assignment can be realistically attempted.

Sample 146-892A-20X-1, 87–92 cm, at 164.4 mbsf, contains rare and poorly preserved radiolarians. But the occurrence of many easily recognizable representatives of *C. davisianna* suggests a late Pliocene to Pleistocene age (<2.9 Ma) for this sample. Co-occurrence of many representatives of the foraminifer *Neogloboquadrina asanoi* in Samples 146-892A-20X-1, 87–92 cm, and 146-892A-20X-CC assigns a late Pliocene age (1.9 to 2.8 Ma) to the interval between 164.4 and 167.4 mbsf.

A possible stratigraphic reversal is thus inferred between 146.90 and 164.4 mbsf. Although the depth of this inversion cannot be located precisely because the interval lacks well-preserved fauna, its occurrence within this interval is consistent with the position of the major fault zone inferred at this site. Rare recrystallized radiolarians not identifiable as to the species level were found in most samples below 87.65 mbsf.

Hole 892D

The co-occurrence of rare but well-preserved representatives of *L. neoheteroporos* and *L. nigriniae* in Samples 146-892D-1X-CC (8.50 mbsf) to -2X-2, 6–8 cm (10.06 mbsf) assigns an age of 1.2–0.9 Ma to the uppermost part of Hole 892D. Radiolarian assemblages of Arctic and upwelling forms are common to abundant in this interval.

The occurrence of rare specimens of *E. matuyamai* in Samples 146-892D-2X-CC (18.00 mbsf) and -3X-1, 5–7 cm (18.05 mbsf)

gives an age of 2–1 Ma to the interval between 18.00 and 27.50 mbsf (Sample 146-892D-3X-CC). This last sample is tentatively included in the interval, although no representative of *E. matuyamai* was found in it, because no other changes occur in the radiolarian fauna.

The lower interval between Samples 146-892D-4X-1, 29–31 cm, to -5X-1, 52–54 cm (27.79 to 37.52 mbsf) is dated 2.9 to 2 Ma because no specimens of *E. matuyamai* were found in assemblages composed of forms such as *L. neoheteroporos*, *C. davisianna* (old forms), *C. pliocenica*, and *E. calvertense*.

Two meters below this sequence, a completely different assemblage of middle Miocene age was found in Sample 146-892D-5X-2, 95–97 cm (39.45 mbsf). Species such as *C. japonica*, *D. splendens*, *B. deflandrei*, *P. titan*, *L. n. nipponica*, and *T. redondoensis* are present in this sample and in the sediments down to Sample 146-892D-7X-5, 32–34 cm (60.08 mbsf). An age of 15.5 to 10 Ma is proposed for this interval, suggesting a major hiatus of nearly 7 m.y. between 37.52 and 39.45 mbsf.

The occurrence of common well-preserved representatives of *L. neoheteroporos*, *L. heteroporos*, and *C. pliocenica* in Sample 146-892D-7X-6, 21–23 cm (61.24 mbsf), dated at 2.9–2.8 Ma, suggests a big stratigraphic reversal between 60.09 and 61.24 mbsf. No representatives of *C. davisianna* are present in this sample, but transitional forms of *L. heteroporos*/*L. neoheteroporos* suggest an age close to the FO of both *C. davisianna* and *L. neoheteroporos*. An identical radiolarian assemblage of the same age is present in all samples down to 77.60 mbsf (Sample 146-892D-9X-CC). All samples between 102.70 and 109.70 mbsf are barren of radiolarians.

Occurrence of old forms of *C. davisianna*, with representatives of *L. neoheteroporos* and *L. heteroporos* give an age slightly younger than 2.9–2.8 Ma to the interval between Samples 146-892D-11X-2, 46–48 cm, and -11X-CC (111.46 to 119.00 mbsf). The scarcity and the poor state of preservation of radiolarians does not permit placement of any hiatus or stratigraphic reversal between this sequence and the upper interval.

From Sample 146-892D-12X-2, 91–93 cm (121.41 mbsf) to 146-892D-16X-5, 53–55 cm (163.06 mbsf), radiolarian assemblages and radiolarian stratigraphic markers are rare and poorly preserved.

Hole 892E

Five samples from 0.64 to 39.83 mbsf were studied. Radiolarian species are usually rare so no good stratigraphy can be proposed.

CONCLUSIONS

Radiolarian assemblages are generally rare and moderately to well-preserved in detrital sediments from the Cascadia Margin.

Detailed occurrence of 65 taxa were estimated in 92 samples at Hole 889A on the accretionary wedge off Vancouver Island. Age estimates obtained from the radiolarian events range from 0.4 Ma at 29.50 mbsf to younger than 2.9 Ma at 284.10 mbsf (Fig. 2). An average sedimentation rate of 9 cm/k.y. can be estimated for the early Pleistocene. No hiatuses are recorded.

LO and FO of stratigraphic markers were recognized in 50 samples from Hole 892A, and in 47 samples from Hole 892D, on the accretionary wedge of the Oregon continental slope. An intricate stratigraphy for Hole 892A is proposed (Fig. 3), from the top to the bottom:

1. approximately 1.5 m of Pleistocene sediments (1–0.9 Ma);
2. a small interval (1.6 m) of lower Pleistocene material (1.7–1 Ma);
3. a hiatus of about 3 m.y., located between 23.2 and 29.1 mbsf;
4. a short sequence (12 m) of late Miocene age (7.5–6.4 Ma);

5. second hiatus spanning 2 to 3 m.y., located between 41.43 and 45.70 mbsf;
6. the first stratigraphic reversal between 49 and 50.50 mbsf;
7. an upper Pliocene sequence (2.9–2.8 Ma) from 50.43 to 61.27 mbsf;
8. a third hiatus of approximately 3 m.y., located between 61.27 and 63.49 mbsf;
9. a very short interval (1.3 m thick) of late Miocene age (7.5 to 6.4 Ma);
10. a second stratigraphic reversal located between 65 and 67 mbsf;
11. a Pliocene interval between 67.92 and 78.44 mbsf, identical in age (2.9–2.8 Ma) and fauna to the upper Pliocene interval above the third hiatus;

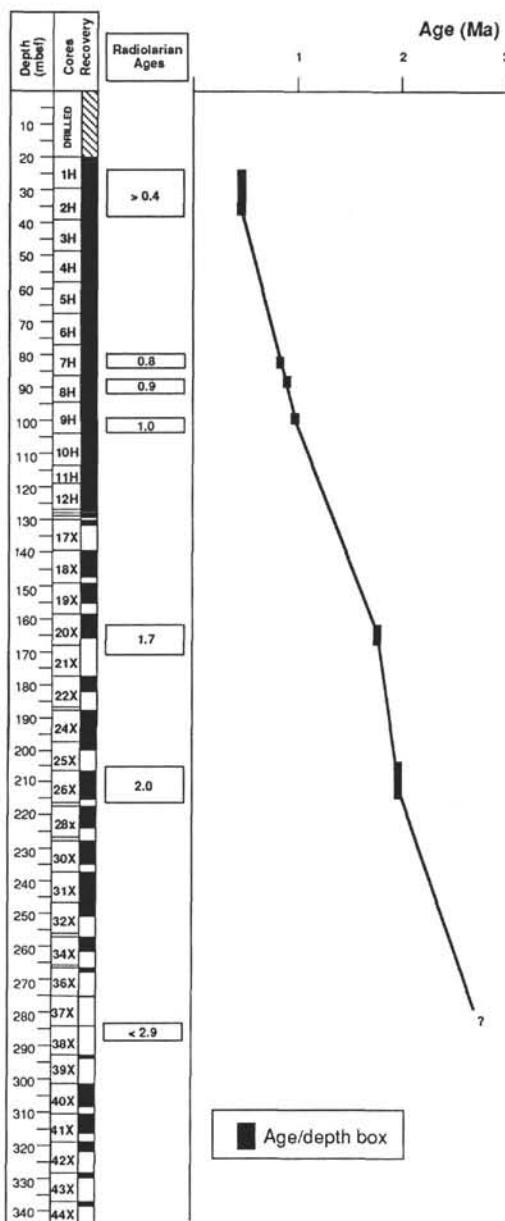


Figure 2. Age-depth plot for Hole 889A.

12. a thick sequence of sediments with no characteristic radiolarians and unknown age from 80.61 to 146.9 mbsf;
13. a possible third stratigraphic reversal between 146 and 164.54 mbsf;
14. a sample younger than 2.9 Ma at 164.4 mbsf;
15. a downhole sequence barren of radiolarians.

Occurrence of many radiolarian species characteristic of upwelling environment suggests that upwelling activity occurred along the Oregon margin during most of the late Pliocene to early Pleistocene.

SPECIES LIST

This list includes bibliographic references for all the taxa that were counted in this study, and some additional remarks about new taxonomic combinations, or morphological peculiarities. The only literature references given are to the original description and to my present concept of the species if different from the original one.

Some groups (mostly actinomids, trissocyklids, and botryoids) were purposely not included in this study because taxonomic concepts of these groups are not yet clear.

Some apparently new morphotypes were not described because their occurrences were scarce and no stratigraphic range could be determined for them.

- Acrosphaera murrayana* (Haeckel). *Choenicospaera murrayana* Haeckel, 1887, p. 102, pl. 8, fig. 4. *Acrosphaera murrayana* (Haeckel) in Hilmers, 1906, p. 63.
- Actinomma leptodermum* (Jørgensen), n. comb. *Echinomma leptodermum* Jørgensen, 1905, p. 116, pl. 8, fig. 33. Included in the genus *Actinomma* without a specific validation by M.G.Petrushevskaya, in Petrushevskaya and Kozlova, 1979, p. 98.
- Actinomma popofskii* (Petrushevskaya). *Echinomma popofskii* Petrushevskaya, 1967, p. 23, fig. 12, I–III. *Actinomma popofskii* (Petrushevskaya) in Caulet, 1986, p. 851.
- Anthocyrtidium pliocenica* (Seguenza). *Anthocyrtis ehrenbergi* Stöhr var. *pliocenica* Seguenza, 1880, p. 232. *Anthocyrtidium pliocenica* (Seguenza) in Nigrini and Caulet, 1988, p. 355, pl. 2, figs. 5–6.
- Botryopera deflandrei* Petrushevskaya, 1975, p. 592, pl. I 1, figs. 30–32.
- Botryostrobus aquilonaris* (Bailey). *Eucyrtidium aquilonaris* Bailey, 1856, p. 4, pl. 1, fig. 9. *Botryostrobus aquilonaris* (Bailey) in Nigrini, 1977, p. 246, pl. 1, fig. 1.
- Botryostrobus aquilonaris* aff. This morphotype differs from the *B. aquilonaris* type by its larger size and additional post-thoracic segments. Stratigraphic range unknown.
- Botryostrobus bramlettei* (Campbell and Clark). *Lithomitra bramlettei* Campbell and Clark, 1944, p. 53, pl. 7, figs. 10–14. *Botryostrobus b. bramlettei* (Campbell and Clark) in Caulet, 1979, p. 129, pl. 1, fig. 8.
- Botryostrobus bramlettei costatus* (Stöhr). n. comb. *Dictyomitra costata* Stöhr, 1880, p. 101, pl. 3, fig. 23. This form has been recently rediscovered in the Messinian of Sicily (Caulet, in prep.).
- Botryostrobus bramlettei pretumidulus* Caulet, 1979, p. 129, pl. 1, fig. 5.
- Botryostrobus bramlettei seriatus* (Jørgensen). *Eucyrtidium seriatum* Jørgensen, 1905, p. 150. *Botryostrobus bramlettei seriatus* (Jørgensen) in Caulet, 1979, p. 130, pl. 1, fig. 6, text-fig. 4.
- Botryostrobus bramlettei tumidulus* (Bailey). *Eucyrtidium tumidulus* Bailey, 1856, p. 5, pl. 1, fig. II. *Botryostrobus bramlettei tumidulus* (Bailey) in Caulet, 1979, p. 131, pl. 1, fig. 9.
- Ceratocyrtis histrionica* (Jørgensen). *Helotholus histrionosa* Jørgensen, 1905, p. 137, pl. 16, figs. 86–88. *Ceratocyrtis histrionica* (Jørgensen) in Petrushevskaya, 1971b, p. 98, fig. 52, 2–4.
- Ceratospyris borealis* Bailey, 1856, p. 31, pl. 1, fig. 3; Nigrini and Moore, 1979, N, Pl. 19, figs. 1a–d.
- Cycladophora bicornis amphora* Lombardi and Lazarus, 1988, p. 110, pl. 4, figs. 6–12. Under this name are grouped those *Cycladophora* having a thorax clearly divided into an elongated upper thorax, and an elongated to equant, narrowly flared, and flat in outline, lower thorax. One of the two forms described by Bailey (1856, fig. 13) under the name of *Halicalyptia*?

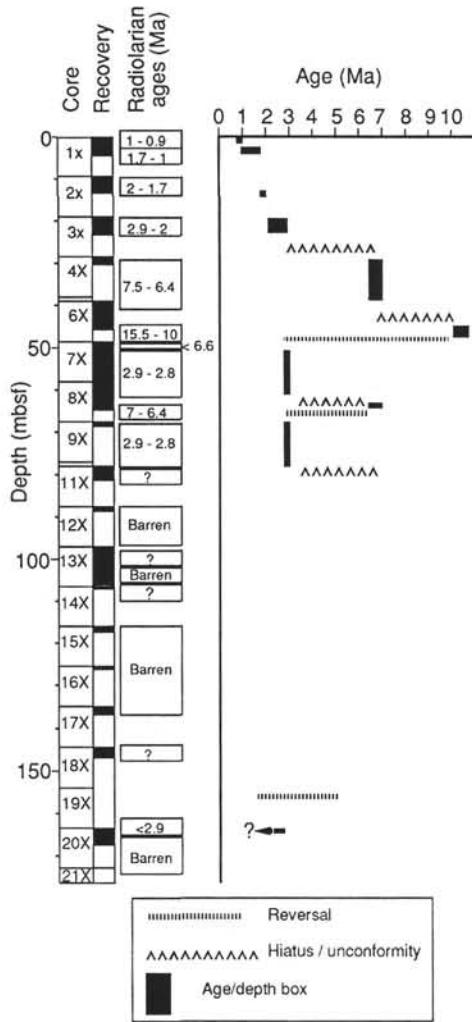


Figure 3. Age-depth plot for Hole 892A.

cornuta (1856, p. 5, pl. 1) looks very similar to the morphotypes counted under this name, but Bailey's name has mostly been used to described those *Cycladophora* having a cylindrical thorax not divided in two parts, similar to the second form described by Bailey (1856, fig. 15).

Cycladophora bicornis helios Lombardi and Lazarus, 1988, p. 114, pl. 5, figs. 1–8.

Cycladophora davisihana davisihana (Ehrenberg). *Cycladophora? davisihana*, Ehrenberg, 1861, p. 297; 1873, pl. 2, fig. 11. *Cycladophora davisihana* Ehrenberg in Petrushevskaya, 1967, p. 122, fig. 69 (I–VII). Early representatives of *C. davisihana* are more robust and shorter than their common descendants. They are referred here as "C. davisihana old".

Cycladophora pliocenica (Hays). *Clathrocyclas bicornis* Hays, 1965, p. 179, pl. 3, fig. 3. *Cycladophora pliocenica* (Hays) Lombardi and Lazarus, 1988, p. 104.

Cycladophora robusta Lombardi and Lazarus, 1988, p. 105, pl. 2, figs. 1–14.

Cyrtocapsella japonica (Nakaseko). *Eusyringium japonicum* Nakaseko, 1963, p. 193, pl. 4, figs. 1–3. *Cyrtocapsella japonica* (Nakaseko) in Sanfilippo and Riedel, 1970, p. 452, pl. 1, figs. 13–15.

Dicyophimus crisiae Ehrenberg, 1854, p. 241. Nigrini, 1967, p. 66, pl. 6, figs. 7a, b.

Dicyophimus splendens (Campbell and Clark). *Pterocorys (Pterocyrtidium) splendens* Campbell and Clark, 1944, p. 46, pl. 6, figs. 19, 20. *Dicyophimus splendens* (Campbell and Clark) in Caulet, 1986, p. 852. Emend. Morley and Nigrini, in press.

Eucyrtidium calvertense Martin, 1904, p. 450, pl. 130, fig. 5.

Eucyrtidium matuyamai Hays, 1970, p. 213, pl. 1, figs. 7, 9.

Eucyrtidium teuscheri teuscheri Caulet, 1986, p. 851, pl. 5, figs. 5–8.

Gondwanaria dogieli (Petrushevskaya). *Sethoconus? dogieli* Petrushevskaya, 1967, p. 95, pl. 53, figs. 1, 2. *Gondwanaria dogieli* (Petrushevskaya), in Petrushevskaya, 1975, p. 585.

Haliommetta miocenica (Campbell and Clark). *Heliosphaera miocenica* Campbell and Clark, 1944a, p. 16, pl. 2, figs. 10–14. *Haliommetta miocenica* (Campbell and Clark), in Petrushevskaya and Kozlova, 1972, p. 517, pl. 9, figs. 8, 9.

Lamprocyclas hannai (Campbell and Clark). *Calocyclus hannai* Campbell and Clark, 1944a, p. 48, pl. 69, figs. 21, 22. *Lamprocyclas hannai* (Campbell and Clark) in Caulet 1986, p. 852.

Lamprocyclas junonis (Haeckel). *Theoconus junonis* Haeckel, 1887, p. 1401, pl. 69, fig. 7. *Lamprocyclas junonis* (Haeckel) in Caulet, 1986, p. 852, pl. 4, fig. 10.

Lamprocyclas margatensis (Campbell and Clark). *Calocyclus margatensis* Campbell and Clark, 1944a, p. 47, pl. 6, figs. 17–18. *Lamprocyclas margatensis* (Campbell and Clark) in Caulet, 1986, p. 852, pl. 4, fig. 3.

Lamprocyclis heteroporus (Hays). *Lamprocyclas heteroporus* Hays, 1965, p. 179, pl. 3, fig. 1. *Lamprocyclis heteroporus* (Hays), in Kling, 1973, p. 639, pl. 5, figs. 19–21, pl. 15, fig. 6.

Lamprocyclis neoheteroporus Kling. *Lamprocyclis neoheteroporus* Kling, 1973, p. 639, pl. 5, figs. 17, 18, pl. 15, figs. 4, 5.

Lamprocyclis nigriniae (Caulet). *Conarachnium nigriniae* Caulet, 1971 p. 3, pl. 3, figs. 1–4, pl. 4, figs. 1–4. *Lamprocyclis nigriniae* (Caulet) in Kling, 1977, p. 217, pl. 1, fig. 17.

Lipmanella acanthica (Jørgensen). n. comb. *Dictyoceras acanthicum* Jørgensen, 1905, p. 140, pl. 17, fig. 101a, pl. 18, fig. 101b. Included in the genus *Lipmanella* without a specific validation by M.G. Petrushevskaya, in Petrushevskaya and Kozlova, 1979, p. 137.

Lychnocanoma nipponica nipponica (Nakaseko). *Lychnocanium nipponicum* Nakaseko, 1963, p. 168, text-fig. 2, pl. 1, figs. 10–14. *Lychnocanoma nipponica nipponica* (Nakaseko) in Morley and Nigrini, in press.

Lychnocanoma nipponica sakai Morley and Nigrini, in press, pl. 7, figs. 1, 3. *Phormostichoartus caryoforma* (Caulet). *Lithamphora caryoforma* Caulet, 1979, p. 131, pl. 2, fig. 2. *Phormostichoartus caryoforma* (Caulet) in Nigrini and Caulet, 1992, p. 161, pl. 6, figs. 7, 9.

Phormostichoartus crustula (Caulet). *Lithamphora crustula* Caulet, 1979, p. 131, pl. 2, fig. 1. *Phormostichoartus crustula* (Caulet) in Nigrini and Caulet, 1992, p. 161, pl. 6, figs. 10–14.

Phormostichoartus fistula Nigrini, 1977, p. 253, pl. 1, figs. 11–13.

Phormostichoartus multiserialis (Ehrenberg). *Eucyrtidium multiseriatum* Ehrenberg, 1860, p. 768. *Lithocampe? multiseriata* (Ehrenberg) in Petrushevskaya, 1967, p. 135, fig. 16, 1–111. *Phormostichoartus multiserialis* (Ehrenberg) in Caulet, 1986, p. 853.

Phormostichoartus pitomorphus Caulet, 1986a, p. 850, pl. 3, figs. 3, 4, 9, 10, 12.

Phormostichoartus schneideri Nigrini and Caulet, 1992, p. 160, pl. 6, figs. 3–6.

Prunopyle titan gr. Campbell and Clark, 1944a, p. 20, pl. 3, figs. 1–3. Arctic Pacific forms are rare and smaller than Antarctic forms.

Pseudocubus vema (Hays). *Helotholus vema* Hays, 1965, p. 176, pl. 2, fig. 3, text fig. A. *Pseudocubus vema* (Hays) in Petrushevskaya, 1971a, p. 46, fig. 24, I–IV. Arctic forms are smaller than Antarctic ones. They are so scarce that a serious morphological study cannot be undertaken.

Pseudodictyophimus gracilipes (Bailey) gr. *Dictyophimus gracilipes* Bailey, 1856, p. 4, pl. I, fig. 8. *Pseudodictyophimus gracilipes* (Bailey) in Petrushevskaya, 1971b, p. 93, figs. 47–49.

Pterocanium auritum Nigrini and Caulet, 1992, p. 152, pl. 4, figs. 6–8.

Pterocanium grandiporus Nigrini, 1968, p. 57, pl. 1, fig. 7.

Pterocanium korotnevi (Dogel). *Pterocorys korotnevi* Dogel in Dogel and Reshetnyak, 1952, p. 17, fig. 11. *Pterocanium korotnevi* (Dogel) in Nigrini, 1970, p. 170, pl. 3, figs. 10, 11.

Pterocanium trilobum (Haeckel). *Dictyopodium trilobum* Haeckel, 1860, p. 839; 1862, p. 340, pl. 8, figs. 6–10. *Pterocanium trilobum* (Haeckel) in Haeckel, 1887, p. 1333.

Pterocorys clausus (Popofsky). *Lithornithium clausum* Popofsky, 1913, p. 393, text fig. I 1. *Pterocorys clausus* (Popofsky) in Petrushevskaya and Kozlova, 1972, p. 545, pl. 36, figs. 16–18.

Rhizosphaera antarctica (Haeckel). *Spongoplegma antarctica* Haeckel, 1887, p. 90. *Rhizosphaera antarcticum* (Haeckel) in Caulet, 1986a, p. 853.

Siphocampe arachnea (Ehrenberg) group. *Eucyrtidium lineatum arachneum* Ehrenberg, 1861, p. 299. *Lithomitra arachnea* (Ehrenberg) in Riedel, 1958, p. 242, pl. 4, figs. 7, 8. *Siphocampe arachnea* (Ehrenberg) gr. in Nigrini, 1977, p. 255, pl. 3, figs. 7, 8.

- Siphocampe modeloensis* (Campbell and Clark). *Lithocampe modeloensis* Campbell and Clark, 1944a, p. 59 pl. 7, figs. 28–30. *Siphocampe modeloensis* (Campbell and Clark) in Caulet, 1986, p. 853.
- Sphaeropyle langii* Dreyer, 1889, p. 13, pl. 4, fig. 54. *Emend* Foreman, 1975, p. 618, pl. 9, figs. 30–31. *Emend* Morley and Nigrini, in press.
- Sphaeropyle robusta* Kling, 1973, p. 634, pl. 1, figs. 11–12, pl. 6, figs. 9–13, pl. 13, figs. 1–5. *Emend* Foreman, 1975, p. 618, pl. 9, figs 24–26. *Emend* Morley and Nigrini, in press.
- Spirema* sp. *Spirema* sp. in Kling, 1973, p. 635, pl. 7, figs. 23–25.
- Spongodiscus osculosus* (Dreyer). *Spongopyle osculosa* Dreyer, 1889, p. 42, figs. 99, 100. *Spongodiscus osculosus* (Dreyer) in Petrushevskaya, 1967, p. 42, figs. 20–22.
- Spongotrochus glacialis* Popofsky, 1908, p. 228, pl. 27, fig. 1, pl. 28, fig. 2.
- Spongotrochus?* *venustum* (Bailey). *Perichlamyidium venustum* Bailey, 1856, p. 5, pl. 1, figs. 16, 17. *Spongotrochus?* *venustum* (Bailey) in Nigrini and Moore, 1979, S119, pl. 15, figs. 3a, b.
- Stauroxiphus communis* Carnevale, 1908, p. 15, pl. 2, fig. 9.
- Stichocorys peregrina* (Riedel). *Eucyrtidium elongatum peregrinum* Riedel, 1953, p. 812, pl. 85, fig. 2. *Stichocorys peregrina* (Riedel) in Sanfilippo and Riedel, 1970, p. 451, pl. 1, fig. 10.
- ?*Stichocorys wolffii* Haeckel, 1887, p. 1479, pl. 80, fig. 10. Arctic forms seem to be more spinose than tropical ones. Relationships between tropical and high latitude forms are not entirely understood. There is still an uncertainty about the taxonomic determination of these forms because their last occurrence seems to be much younger (7 to 6.4 Ma) in Leg 146 material than in tropical material (LO dated around 11.6 Ma in Leg 85 material, see Nigrini, 1985).
- Stylacontarium acquilonium* (Hays). *Druppatractus acquilonius* Hays, 1970, p. 214, pl. 1, figs 4–5. *Stylacontarium acquilonium* (Hays) in Kling, 1973, p. 634, pl. 1, figs. 17–20, pl. 14, figs. 1–4.
- Stylacontarium* sp. aff. *S. bispiculum* Kling, 1973, p. 634, pl. 6, fig. 6. 19–23, pl. 14, figs. 5–8. These forms are too rare in Leg 146 material to attempt a comparison with Antarctic forms.
- Stylaractus universus* Hays, 1965, p. 167, pl. 1, fig. 6.
- Theocorys redondoensis* (Campbell and Clark). *Theocyrtsis redondoensis* Campbell and Clark, 1944a, p. 49, pl. 7, fig. 4. *Theocorys redondoensis* (Campbell and Clark) in Kling, 1973, p. 638, pl. 11, figs. 26–28.
- Theocorythium trachelium trachelium* (Ehrenberg). *Eucyrtidium trachelium* Ehrenberg, 1872a, p. 312. *Theocorythium trachelium trachelium* (Ehrenberg) in Nigrini, 1967, p. 79, pl. 8, fig. 2, pl. 9, fig. 2. Early representatives of an older form similar to *T. trachelium* are referred here under the name of "T. trachelium old". The description and a picture of this form (pl. 1, fig. 3) are given in a paper from Caulet et al. (1993).
- Tricolocapsa papillosa* (Ehrenberg) gr. *Eucyrtidium papillosum* Ehrenberg, 1872a, p. 310; 1872b, pl. 7, fig. 10. *Tricolocapsa papillosa* (Ehrenberg) gr. in Petrushevskaya and Kozlova, 1972, p. 537, pl. 22, fig. 31.
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ACKNOWLEDGMENTS

Financial support for the cruise and this study was provided by INSU grant "Soutien à ODP-France" No. 93 39 10. The author is much indebted to Joe Morley for communication of his new "radiolarian ages" for the North Pacific (Leg 145), and to David Lazarus for taxonomic discussions about Arctic forms and biostratigraphy. Thanks are due to C. Nigrini and A. Sanfilippo for the critical review of the manuscript. M. Tamby competently assisted with the preparation of samples.

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Date of initial receipt: 29 August 1994

Date of acceptance: 3 February 1995

Ms 146SR-205