

15. HIGH-RESOLUTION PLEISTOCENE DIATOM BIOSTRATIGRAPHY AND PALEOCEANOGRAPHY OF SITE 919 FROM THE IRMINGER BASIN¹

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ABSTRACT

Ocean Drilling Program Site 919 from the Irminger Basin provides one of the first opportunities to study high-resolution Pleistocene high-latitude North Atlantic diatom biostratigraphy and paleoceanography. Three Pleistocene diatom datum events are identified and tied directly to the stable oxygen isotope record of Site 919. These datum events are the (1) last occurrence of *Thalassiosira nidulus* at 0.310 Ma, (2) last occurrence of *Proboscia curvirostris* at 0.335 Ma, and (3) last occurrence of *Neodenticula seminae* at 0.817–0.895 Ma. Based on these datums, a Pleistocene diatom zonation for the high-latitude North Atlantic and the Nordic Seas is proposed using high-latitude cold species. The last occurrences of *Thalassiosira nidulus* and *Proboscia curvirostris* appear to be isochronous in the middle- to high-latitude North Pacific and the North Atlantic Oceans.

The diatom abundance and preservation record of Site 919 is used to infer polar front fluctuations during the last 1 m.y. Results indicate that during glacial Stages 2, 6, 10, 12, and 16 the polar front migrated eastward, covering Site 919 with sea-ice. Milder climate is indicated during glacial Stages 4, 8, 14, 18, and 20, as the polar front is interpreted to lie to the west of Site 919.

INTRODUCTION

The Cenozoic diatom biostratigraphy of the North Atlantic Ocean up to 60°N is well established (Baldauf, 1985, 1987). However, north of 60°N, there is no well established diatom biostratigraphy partly due to the poor coring techniques employed during Deep Sea Drilling Project (DSDP) Leg 38 and Ocean Drilling Program (ODP) Leg 104, which resulted in poor recovery and poor magnetostratigraphy. These handicaps and the rare occurrence of many of the marker species limited the diatom biostratigraphic studies (Schrader and Fenner, 1976; Dzinoridze et al., 1978; Bodén, 1992). Recent drilling in the Nordic Seas (ODP Leg 151) provided excellent recovery and magnetostratigraphy, which led to the initial development of a detailed and reliable Cenozoic diatom biostratigraphy for the Nordic Seas (Koç and Scherer, in press).

The North Atlantic diatom zonation of Baldauf (1985, 1987) is not applicable in the Nordic Seas due to the endemic character of the species in this region. However, sites in the high-latitude North Atlantic (between 60°–65°N) south of Iceland might be a useful link to combine the Nordic Sea biostratigraphic scheme (Koç and Scherer, in press) with the North Atlantic zonation (Baldauf, 1985, 1987).

During ODP Leg 152, a transect of six sites was drilled on the Southeast Greenland Margin (Larsen, Saunders, Cliff, et al., 1994). The main drilling objective was deep sampling of the volcanic basement. To reach this objective, four sites (914–917) were drilled on the shelf and two sites (918–919) in the adjacent deep Irminger Basin. Site 918 is located on the upper continental rise and is highly influenced by a submarine fan system that supplied coarse detritus to this area (see “Lithostratigraphy” section in Shipboard Scientific Party, 1994a).

Site 919 (62°40.20'N, 37°27.611'W) is located in 2088.2 m of water on the continental rise of Southeast Greenland, within the western part of the Irminger Basin (Fig. 1). Two holes were drilled at Site 919: Hole 919A (10 cores) was cored continuously to a depth of 93.5

meters below seafloor (mbsf), and Hole 919B (eight cores) was cored and washed to a depth of 147 mbsf in two parts. Recovery of sediment in the holes exceeds 100%. The two holes, when combined, form a composite sequence 147 m thick. Due to the pelagic character, high rates of Pleistocene sedimentation, and the presence of well-preserved diatoms at this site, Site 919 was chosen for detailed study to refine and increase the diatom biostratigraphic resolution of the Pleistocene middle- to high-latitude North Atlantic and Nordic Seas.

METHODS

No samples were processed on board for diatom biostratigraphic investigations. Later on shore, one sample from each core section (1.5 m) from Holes 919A and 919B was analyzed for its diatom content. Approximately 1–2 g of each dry sample was placed in a 250-mL beaker, and 10 drops of 37% HCl acid were added. After waiting for at least 15 min for the reaction to stop, 10 drops of 30% hydrogen peroxide were added. The sample was gently heated until the liquid became light yellow. The acidity of the sample was diluted by centrifuging it with distilled water until a pH of 7 was achieved. Quantitative slides of acid-cleaned material were examined at 400× for relative diatom abundance. Assessment of total diatom abundance was qualitative. Diatoms were recorded as abundant (A) if six or more specimens were present in one field of view at 400×, common (C) if one to five specimens occurred in one field of view, few (F) if one to four specimens occurred in five fields of view, and rare (R) if one to 10 specimens occurred per horizontal traverse.

Qualitative estimates of abundance of individual taxa are based on the number of specimens observed per field of view at 400×. Species were recorded as abundant (A) if two or more specimens were present in one field of view, common (C) if one to five specimens were observed per five fields of view, few (F) if two to ten specimens were observed per horizontal traverse, rare (R) if one specimen was observed per horizontal traverse, and trace (T) if only fragments of the specimen were observed.

Diatom preservation was recorded as good (G) when both finely silicified and heavily robust forms were present and no significant alteration of the frustules other than minor fragmentation was observed, moderate (M) when the assemblage showed moderate breakage and slight dissolution, and poor (P) when finely silicified forms

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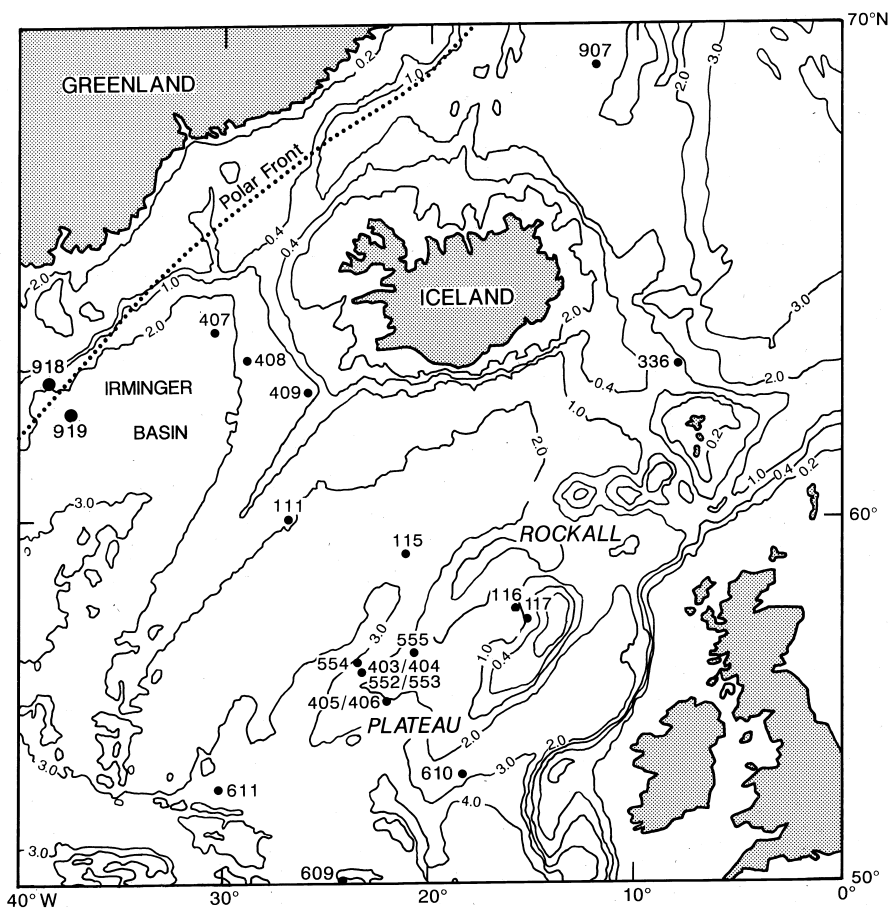


Figure 1. Location of ODP Sites 919 and 918 in the western Irminger Basin. Plotted also are previous DSDP and ODP sites in the area. Position of the polar front is shown by dotted line.

were rare or absent and the assemblage was dominated by robust forms and fragments.

PLEISTOCENE DIATOM BIOSTRATIGRAPHY

At present, the Irminger Basin is influenced by the southward surface flow of the cold and ice-laden East Greenland Current (EGC) in the west, and in the central parts by the confluence of these polar waters with the Atlantic waters of the Irminger Current creating modified North Atlantic waters (Fig. 2). Site 919 in the western Irminger Basin today is primarily under the influence of the EGC. The Pleistocene diatom species observed in sediment recovered from Site 919 primarily consist of modern polar and arctic species (Koç Karpuz and Schrader, 1990; Koç, unpubl. data) (Plate 1). These include *Actinocyclus curvatulus*, *Thalassiosira gravida* spore and vegetative cells, *Thalassiosira trifulta*, *Thalassiothrix longissima*, and the warm-water species *Thalassiosira oestrupii* (Table 1).

Three biostratigraphic events are observed in Holes 919A and 919B: the last occurrence (LO) of *Thalassiosira nidulus*, the LO of *Proboscia curvirostris*, and the LO of *Neodenticula seminae* (Fig. 3; Plate 2). The LO of *Thalassiosira nidulus* is observed in Sample 152-919A-5H-4, 40–41 cm (41.4 mbsf). The LO of *Proboscia curvirostris* is observed in Sample 152-919A-5H-7, 35–36 cm (45.85 mbsf). The LO of *Neodenticula seminae* is observed in Sample 152-919B-7H-7, 40–41 cm (137.4 mbsf) (Table 1).

Correlation to the Isotope Record

Precise dating of these biostratigraphic events is achieved by correlating the events to the oxygen isotope record of Site 919 (Flower,

this volume). The LO of *T. nidulus* and the LO of *P. curvirostris* occur within isotope Stage 9 (Fig. 3). Interpolation of age using the stage boundaries gives an age of 0.309 Ma for the LO of *T. nidulus* and an age of 0.335 Ma for the LO of *P. curvirostris*. The LO of *N. seminae* occurs within isotope Stage 22; however, because there is a barren interval above this event, the age is interpolated to be between 0.817 and 0.895 Ma. The LO of *T. nidulus* is dated between 0.280 and 0.390 Ma in the Northwest Pacific, showing a possible diachrony from high to low latitudes (Koizumi and Tanimura, 1985). The datum of 0.309 Ma for the LO of *T. nidulus* from Site 919 falls closer to the younger datum from the North Pacific. It is possible to conclude that this event is synchronous between the middle to high latitudes of both oceans.

The LO of *P. curvirostris* is dated at 0.300–0.350 Ma in the middle- to high-latitude North Pacific (Koizumi and Tanimura, 1985; Koizumi, 1986). Baldauf (1985) correlated the LO of *P. curvirostris* datum at Site 552 in the North Atlantic to isotope Stage 10 and assigned an approximate age of 0.360 Ma. However, this must be taken as a minimum age because it is based on observations from a single sample. At Site 646, from the southern Labrador Sea, the LO of *P. curvirostris* is observed in Sample 105-646A-3H-1, 128–130 cm, and estimated to be 0.27 Ma based on a linear interpolation between the core top and the Brunhes/Matuyama boundary (Monjanel and Baldauf, 1989). The overlap in age observed between the North Atlantic and the North Pacific suggests synchronicity between these two oceans for this event. Baldauf (1985, 1987) recorded the occurrence of *P. curvirostris* from Sites 552, 606, 607, 609, and 611 in the North Atlantic. *Proboscia curvirostris* is also recorded from the Nordic Seas Sites 907 (Koç and Scherer, in press) and 986 (Jansen and Raymo, in press). This suggests that *P. curvirostris* has a distribution that stretches from 40°N in the North Atlantic to the high-latitude Nordic

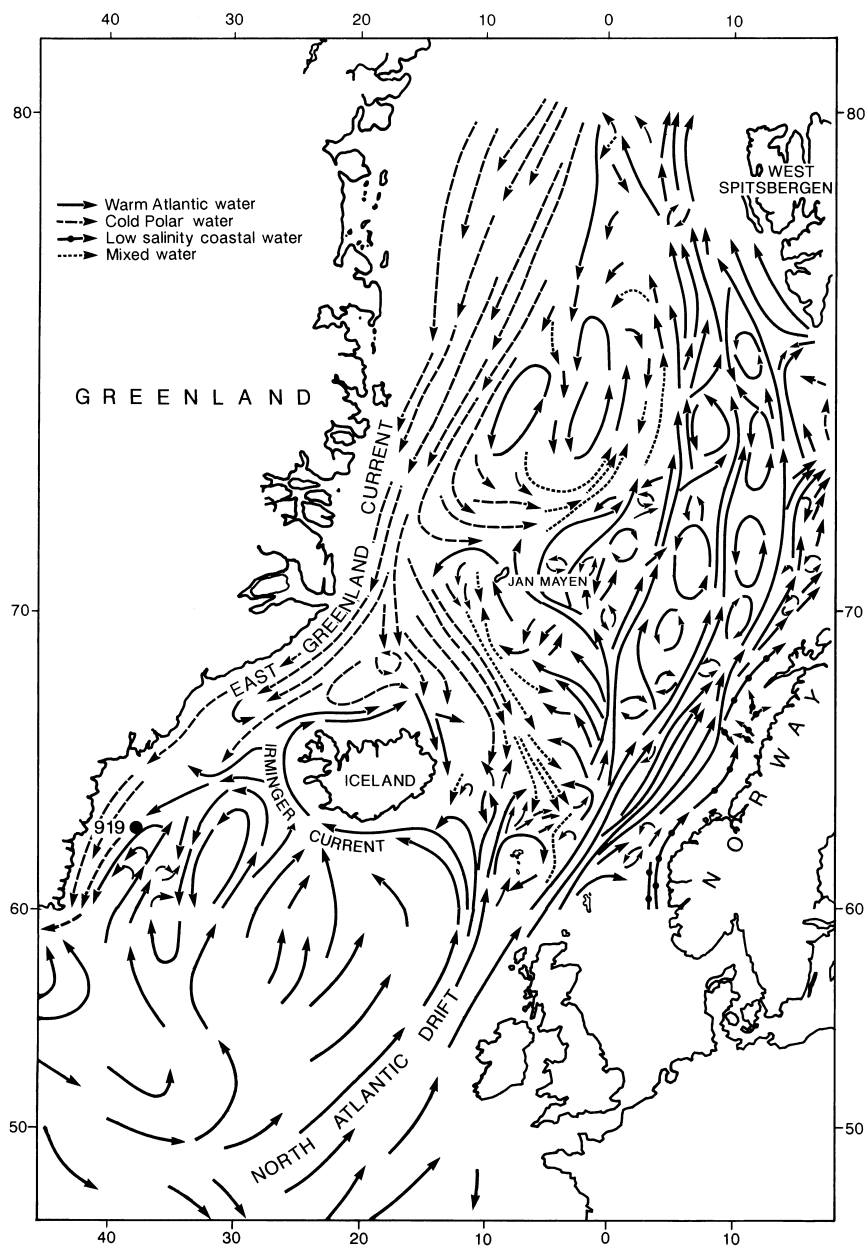


Figure 2. Modern surface currents in the Nordic Seas and the North Atlantic from Alekseev and Istochin (1956). Also plotted is the position of ODP Site 919 in the western Irminger Basin.

Seas, and therefore may be a useful biostratigraphic datum. It also has a middle- to high-latitude distribution in the North Pacific, which makes correlations between the high-latitude North Atlantic and the North Pacific possible.

Denticulopsis seminae belongs to the modern assemblage of the middle- to high-latitude North Pacific (Barron, 1981). However, it went extinct in the North Atlantic during the late Quaternary (Baldauf, 1986). Baldauf (1986) found the LO of *Denticulopsis seminae* to be diachronous between Site 607 (about 0.8 Ma) and Sites 609 and 611 (about 0.7 Ma). Since the age from Site 919 (0.817–0.895 Ma) is closer to the one from Site 607, which is the southernmost site, it seems that the calculated ages from Sites 609 and 611 are too young for this event. Monjanel and Baldauf (1989) recorded the LO of *D. seminae* at Site 646 from the southern Labrador Sea in Sample 105-646A-7H-5, 120–122 cm. We have correlated this sample to the oxygen isotope record of Hole 646A (Aksu et al., 1989). It plots near the midpoint of isotope Stage 21, thereby giving an age of about 0.818 Ma. This age correlates well with the age of 0.817–0.895 Ma from Site 919.

Zonation

The diatom zones defined for the North Atlantic (Baldauf, 1985, 1987) are based partially on eastern equatorial North Pacific diatom zonation schemes (Barron, 1981; Baldauf, 1986). Thus, most of the zones are defined by warm-temperate species. The youngest North Atlantic diatom zone, the *Pseudoeunotia doliolus* Zone (Baldauf, 1985, 1987), was defined by Burckle (1977) from the equatorial Pacific as the interval containing *Pseudoeunotia doliolus* that ranges stratigraphically above the LO of *Nitzschia reinholdii* (0.440 Ma) to the present. This zone also was recognized in DSDP Leg 81 and 94 sites from the middle- and high-latitude North Atlantic (Baldauf, 1985, 1987). However, it was impossible to recognize this zone both at Sites 919 and 907 due to the absence of *Nitzschia reinholdii* and the very sparse occurrence of *Pseudoeunotia doliolus*.

A Pleistocene diatom zonation scheme is proposed below using high-latitude cold-water species for the high-latitude North Atlantic and the Nordic Seas (modified from Koç and Scherer, in press) (Fig. 4).

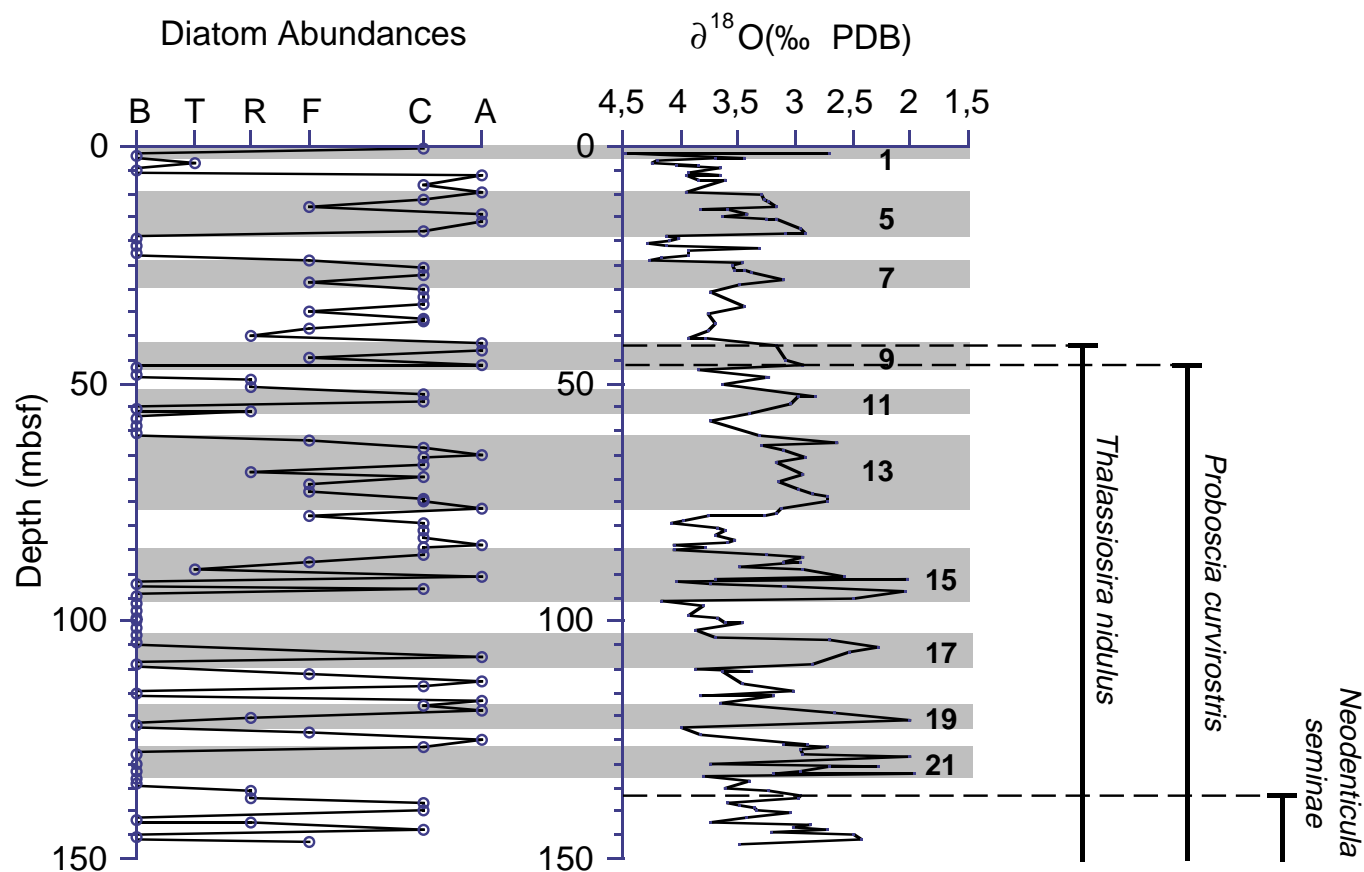


Figure 3. Site 919 relative diatom abundances (B = barren, T = trace, R = rare, F = few, C = common, and A = abundant) plotted with the $\delta^{18}\text{O}$ record against composite depth. Interglacial stages are shaded and numbered. Diatom biostratigraphic events and their ranges are indicated and correlated to the $\delta^{18}\text{O}$ record.

Thalassiosira oestrupii Zone

Authors. Koç and Scherer (in press).

Definition. The interval from the LO of *Proboscia curvirostris* to the present.

Remarks. Koç and Scherer (in press) could not date the base of this zone precisely because the LO of *P. curvirostris* occurred within an interval that was mostly barren.

Correlation. This zone correlates to the *Neodenticula seminae* (NPD12) Zone (Donahue, 1970) of the North Pacific, which also uses the LO of *P. curvirostris* to define its base (Fig. 5).

Age. Quaternary, 0.335–0 Ma.

Proboscia curvirostris Zone

Author. Koç (this study).

Definition. The top of the zone is defined by the LO of *P. curvirostris* and the base by the LO of *Neodenticula seminae*.

Remarks. The base of this zone is modified from its original description by Koç and Scherer (in press) as the interval from the LO of *Thalassiothrix miocenica* to the LO of *P. curvirostris*. This interval included an extended, mostly barren section at its original site of description (Site 907). The LO of *Thalassiosira nidulus* occurs right above the top of this zone (Fig. 4).

Correlation. This zone correlates to the *Rhizosolenia curvirostris* (synonym for *Proboscia curvirostris*) Zone of the North Pacific (Donahue, 1970) (Fig. 5).

Age. Quaternary, 0.817–0.894 to 0.335 Ma.

Neodenticula seminae Zone

Author. Koç (this study).

Definition. The top of the zone is defined by the LO of *Neodenticula seminae* and the base by the LO of *Thalassiothrix miocenica*.

Remarks. The base of this zone is not observed at Site 919.

Correlation. This zone encompasses the North Pacific zones from the *Actinocyclus oculatus* Zone to the *Neodenticula koizumii*-*N. kamschatica* Zones (Koizumi, 1973).

Age. Late Pliocene through late Pleistocene, 3.328 to 0.894–0.817 Ma.

DIATOM ABUNDANCES: PALEOCEANOGRAPHIC IMPLICATIONS

Site 919 lies in a position to monitor fluctuations in the polar front (Fig. 1). Diatom abundance and preservation at this site fluctuate in response to changing oceanographic conditions affecting diatom productivity. Correlating variation in diatom abundance to the oxygen isotope record of Site 919 shows that generally the interglacials are related to high diatom productivity (common to abundant diatoms) (Fig. 3), whereas glacial Stages 2, 6, 10, 12, and 16 are barren of diatoms. However, during glacial Stages 4, 8, 14, 18, and 20, there was significant diatom production.

Lack of surface-water productivity during Stages 2, 6, 10, 12, and 16 indicates that these were the intervals during which the polar front migrated farthest east, thereby covering Site 919 with sea-ice. Considering the proximity of the sea-ice margin to the area at present, it is very probable that these diatom-barren glacial intervals reflect periods when the site was under heavy sea-ice cover. We can, thus, conclude that these were the severest glacial periods of the last 1 m.y.

The presence of significant diatom production during glacial Stages 4, 8, 14, 18, and 20 indicates open marine conditions over Site 919 during these times. Benthic oxygen isotope records show light isotopic excursions, indicating some warming during these glacials (Shackleton et al., 1990). Planktonic oxygen isotope records from the Nordic Seas also record these light isotopic excursions (Fronval and Jansen, in press). Results indicate that during the glacial Stages 4, 8,

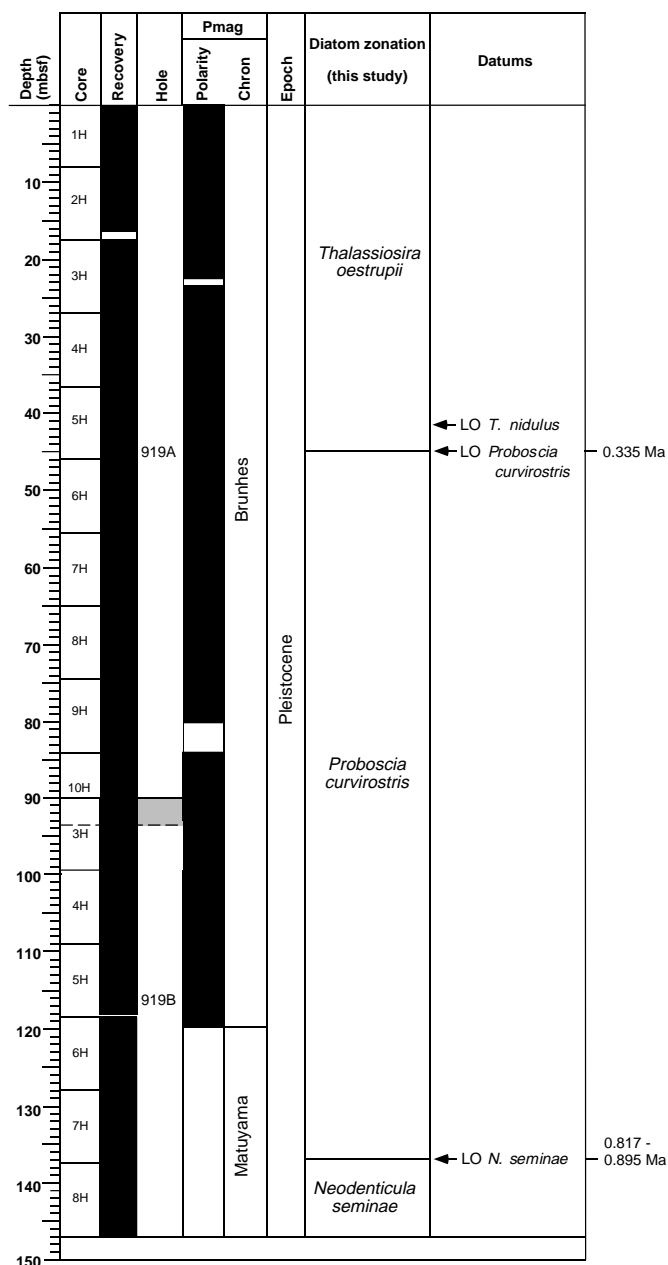


Figure 4. Stratigraphic position of cores, recovery, paleomagnetic stratigraphy, epoch, diatom zonation, and datums. The magnetostratigraphy shown is that of Shipboard Scientific Party (1994b). LO = last occurrence.

14, 18, and 20, the polar front was positioned west of Site 919, and that these glacial periods were milder than Stages 2, 6, 10, 12, and 16.

CONCLUSIONS

Diatom investigations of Site 919 have refined and increased the Pleistocene diatom biostratigraphic resolution of the high-latitude North Atlantic considerably.

Three Pleistocene diatom datum events are identified and tied directly to the stable oxygen isotope record of Site 919. These datum events are (1) the LO of *Thalassiosira nidulus* at 0.310 Ma, (2) the LO of *Proboscia curvirostris* at 0.335 Ma, and (3) the LO of *Neodenticula seminae* around 0.817–0.895 Ma.

Three diatom zones, based on these datums, are proposed for the Pleistocene in the high-latitude North Atlantic and the Nordic Seas.

The record of diatom abundance and preservation at Site 919 indicates that during glacial Stages 2, 6, 10, 12, and 16, the polar front migrated eastward covering Site 919 with sea-ice. During glacial Stages 4, 8, 14, 18, and 20, the polar front was positioned west of Site 919, indicating that these glacial periods were milder than Stages 2, 6, 10, 12, and 16.

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APPENDIX

List of Flora

- Actinocyclus curvatulus* Janisch, 1878, in Schmidt et al. (1874–1959), pl. 57, fig. 31; Hustedt, 1958, pp. 129–130, pl. 8, fig. 81.
- Bacterosira fragilis* Gran, 1900, p. 114; Jousé, 1962, fig. 2, no. 15; Sancetta, 1982, p. 227, pl. 2, figs. 1–4.
- Coscinodiscus asteromphalus* Ehrenberg, 1844, in Hustedt, 1930, pp. 452–454, fig. 250.
- Coscinodiscus marginatus* Ehrenberg, 1841, p. 142; Hustedt, 1930, p. 416, fig. 223.
- Coscinodiscus oculus iridis* Ehrenberg, 1839, p. 147; Hustedt, 1930, pp. 454–459, fig. 252.
- Coscinodiscus radiatus* Ehrenberg, 1839, p. 148, pl. 3, figs. 1a–c; Hustedt, 1930, p. 420, fig. 255.
- Detonula confervacea* (Cleve) Gran. Hasle, 1973, pp. 1–49.
- Neodenticula seminae* (Simonsen and Kanaya) Akiba and Yanagisawa, 1986, p. 491, 492, pl. 24, figs. 1–11, pl. 26, figs. 1–10. Synonym: *Denticula seminae* Simonsen and Kanaya, 1961, p. 503, pl. 1, figs. 26–32; *Denticulopsis seminae* (Simonsen and Kanaya) Simonsen, 1979, p. 65.
- Nitzschia cylindrica* (Grunow) Hasle, 1965, pl. 12, figs. 6–12, pl. 14, figs. 1–10, pl. 17, figs. 2–4; Sancetta, 1982, pl. 3, figs. 6, 7.
- Nitzschia grunowii* Hasle, 1965, pl. 1, figs. 15–19, pl. 2, figs. 6–9, pl. 3, figs. 1, 2, pl. 16, figs. 1, 2; Sancetta, 1982, pl. 3, figs. 8–10.

Nitzschia panduriformis Gregory, 1857, p. 529, fig. 57; Fenner, 1978, p. 525, pl. 32, figs. 27–29; Baldauf, 1985, pl. 5, fig. 8.
Nitzschia riedelia Schrader, 1974, pl. 1, figs. 14a, b, 19.
Paralia sulcata (Ehr.) Cleve. Hustedt, 1930, fig. 119; Sancetta, 1982, pl. 3, figs. 13–15.
Porosira glacialis (Grunow) Jørgensen, 1905; Schrader and Fenner, 1976, pl. 16, figs. 1–4; pl. 197, fig. 1.
Proboscia alata Sundström, 1986, pp. 99–102, pl. 36, figs. 258–266; Synonym: *Rhizosolenia alata* Brightwell, 1858, p. 95, pl. 5, figs. 8, 8a.
Proboscia curvirostris Jordan and Priddle, 1991, p. 57; Synonym: *Rhizosolenia curvirostris* Jousé, 1959, p. 48, pl. 2, fig. 17; Akiba and Yanagisawa, 1986, p. 497, pl. 42, figs. 1, 2; pl. 45, figs. 1–6.
Pseudoeunotia doliolus (Wallich) Grunow in Van Heurck, 1881, pl. 35, fig. 22; Baldauf, 1985, pl. 5, figs. 5–7.
Rhizosolenia hebetata forma *hebetata* Sundström, 1986, p. 47, 48, pl. 4, figs. 18 a, b; pl. 17, figs. 112, 113.
Rhizosolenia hebetata forma *semispina* (Hensen) Gran, 1904, p. 524, pl. 17; Sundström, 1986, p. 48–52, pl. 4, figs. 19, 20; pl. 17, figs. 117, 118.
Rhizosolenia styliformis Brightwell, 1858, p. 95, pl. 5, figs. 5a, b, d; Baldauf, 1987, pl. 6, fig. 7.
Stephanopyxis turris (Greville and Arnott) Ralfs in Pritchard, 1861, p. 826, pl. 5, fig. 74; Schrader and Fenner, 1976, pl. 30, figs. 1–10, 14; pl. 37, figs. 17–19.

Thalassionema nitzschioides (Grunow) Van Heurck, 1896, p. 319, fig. 75; Schrader and Fenner, 1976, pl. 5, fig. 8.
Thalassiosira anguste-lineata (A. Schmidt) Fryxell and Hasle, 1977, p. 73, figs. 22–34.
Thalassiosira decipiens (Grunow) Jørgensen, 1905; Hustedt, 1930, pp. 322–323, fig. 158; Schrader and Fenner, 1976, pl. 17, fig. 9.
Thalassiosira gravida Cleve, 1896, p. 12, pl. 2, figs. 14–16; Hustedt, 1928, p. 325, fig. 161; Schrader and Fenner, 1976, pl. 16, figs. 5, 6; pl. 17, fig. 2.
Thalassiosira hyalina (Grunow) Gran. Hustedt, 1930, fig. 159; Hasle and Heimdal, 1968, fig. 22; Koizumi, 1986, pl. 4, fig. 11.
Thalassiosira nidulus (Tempère and Brun) Jousé, 1961, p. 63; Schrader and Fenner, 1976, pl. 17, figs. 13, 16.
Thalassiosira nordenskiöldii Cleve, 1873, p. 7, pl. 1, fig. 11; Hustedt, 1928, p. 321, fig. 157; Sancetta, 1982, p. 242, pl. 5, figs. 8–9.
Thalassiosira oestrupii (Ostenfeld) Proshkina-Lavrenko, 1949; Hasle, 1960, p. 8, pl. 1, figs. 5–7; Akiba, 1986, pl. 14, figs. 1–6.
Thalassiosira plicata Schrader, 1974, p. 917, pl. 3, figs. 1, 2, 4–9.
Thalassiosira trifulta Fryxell in Fryxell and Hasle, 1979, p. 16, pls. 1–5, figs. 1–24; Sancetta, 1982, p. 244, pl. 5, figs. 10–12.
Thalassiosira zabelinae Jousé, 1961, p. 66, pl. 2, figs. 1–7; Akiba, 1986, pl. 8, fig. 11.
Thalassiothrix longissima Cleve and Grunow, 1880, Hustedt, 1932, p. 247, fig. 726; Hasle and Semina, 1987, figs. 1–25.

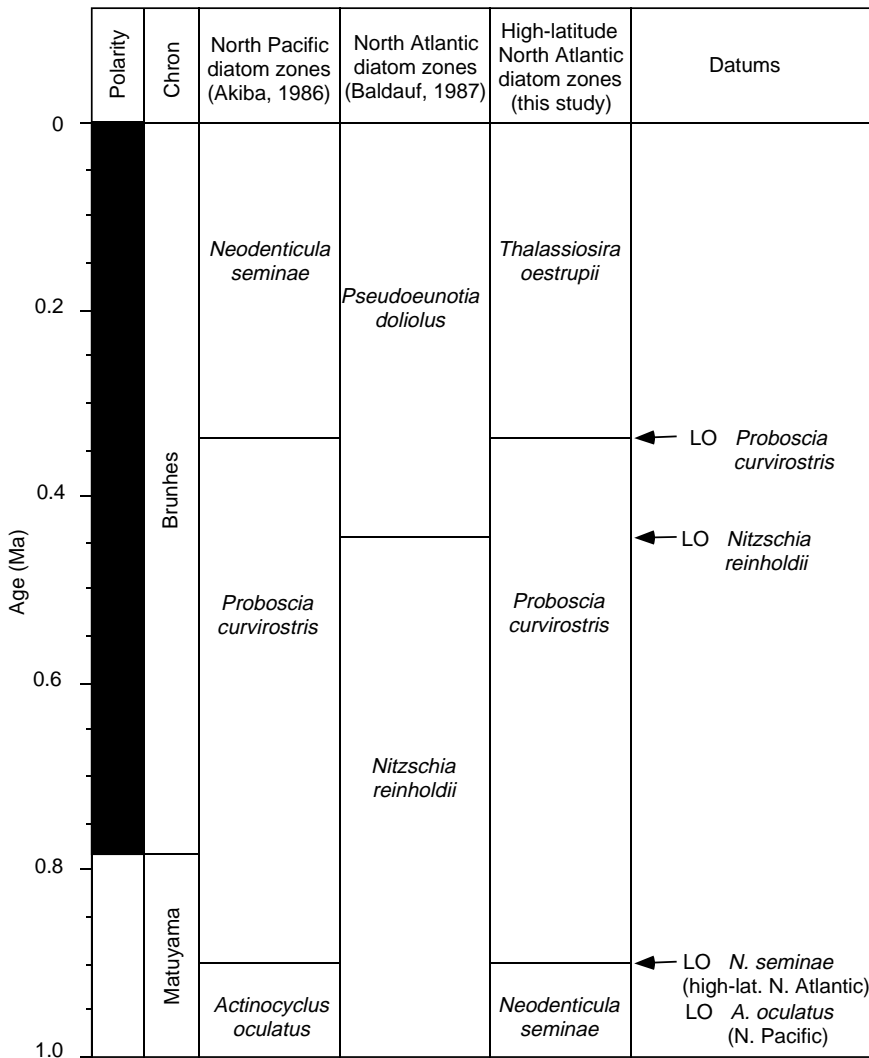


Figure 5. Correlation of various Pleistocene diatom zones and the first-order datums. LO = last occurrence.

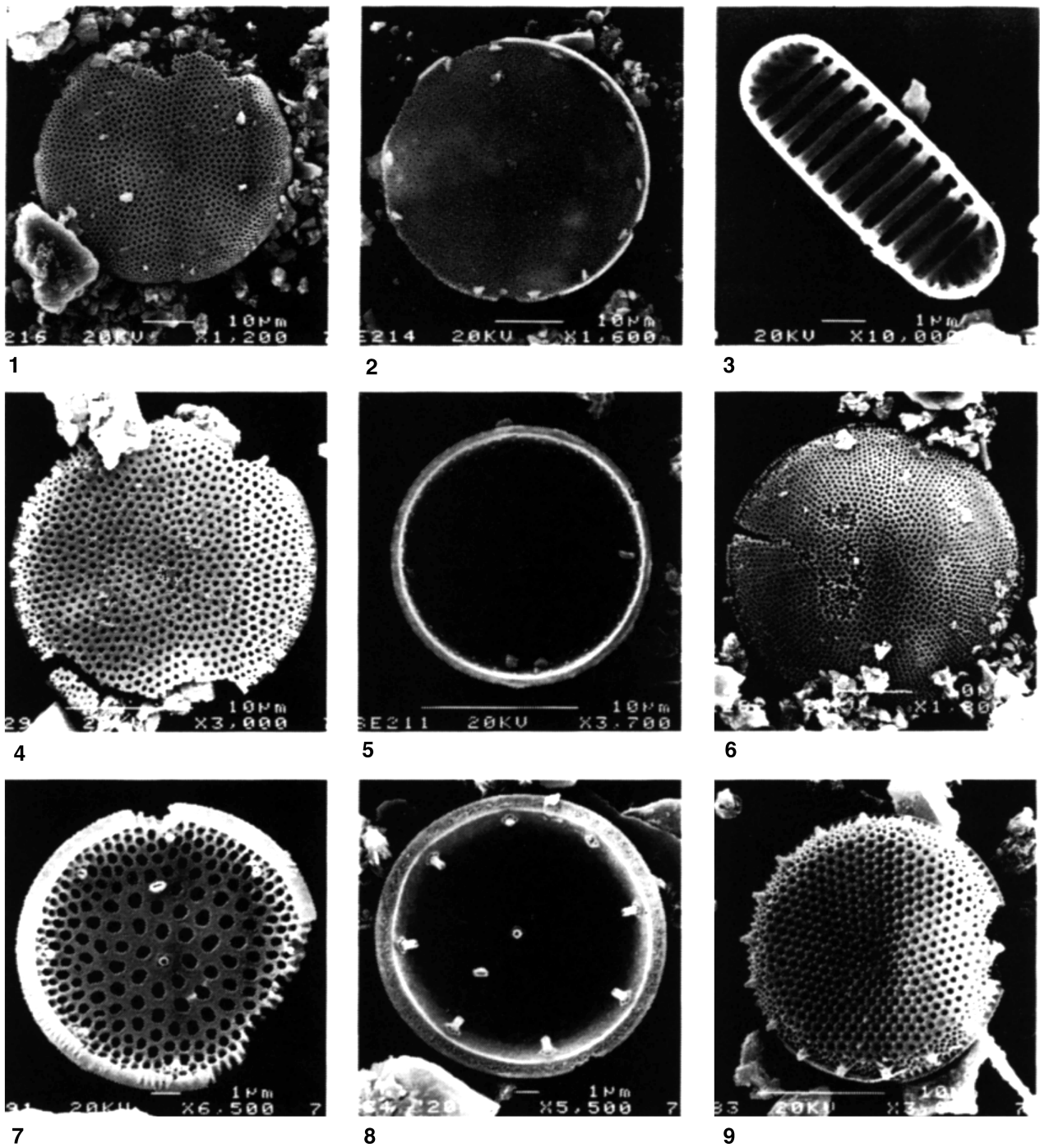


Plate 1. Scale bar = 10 μm, unless otherwise indicated. **1.** *Actinocyclus curvatulus* Janisch, SEM, external view of valve face with arrow pointing at the pseudonodulus, Sample 152-919A-5H-7, 35–36 cm. **2.** *Actinocyclus curvatulus* Janisch, SEM, internal view of valve face with labiate processes along the margin, Sample 152-919A-5H-7, 35–36 cm. **3.** *Nitzschia cylindrica* (Grunow) Hasle, SEM, internal view of valve face, Sample 152-919A-7H-7, 40–41 cm. Scale bar = 1 μm. **4.** *Thalassiosira gravida* Cleve vegetative cell, SEM, external view of valve face with more than two rows of marginal spines, cluster of central pores and spines scattered all over the valve face, Sample 152-919B-8H-5, 20–21 cm. **5.** *Thalassiosira gravida* Cleve vegetative cell, SEM, internal view of valve face with, in addition to the above features, a marginal labiate process at three o’clock, Sample 152-919A-5H-7, 35–36 cm. **6.** *Thalassiosira gravida* Cleve resting spore, SEM, external view of valve face, Sample 152-919A-7H-7, 40–41 cm. **7.** *Thalassiosira oestrupii* (Ostenfeld) Proshkina-Lavrenko, internal view of valve face with a row of marginal spines, a central process and a labiate process halfway between the center and the margin, Sample 152-919B-8H-5, 20–21 cm. Scale bar = 1 μm. **8.** *Thalassiosira oestrupii* (Ostenfeld) Proshkina-Lavrenko, internal view of valve face, Sample 152-919A-7H-7, 40–41 cm. Scale bar = 1 μm. **9.** *Thalassiosira angustilineata* (A. Schmidt) Fryxell and Hasle, SEM, external view of valve face, Sample 152-919A-7H-7, 40–41 cm.

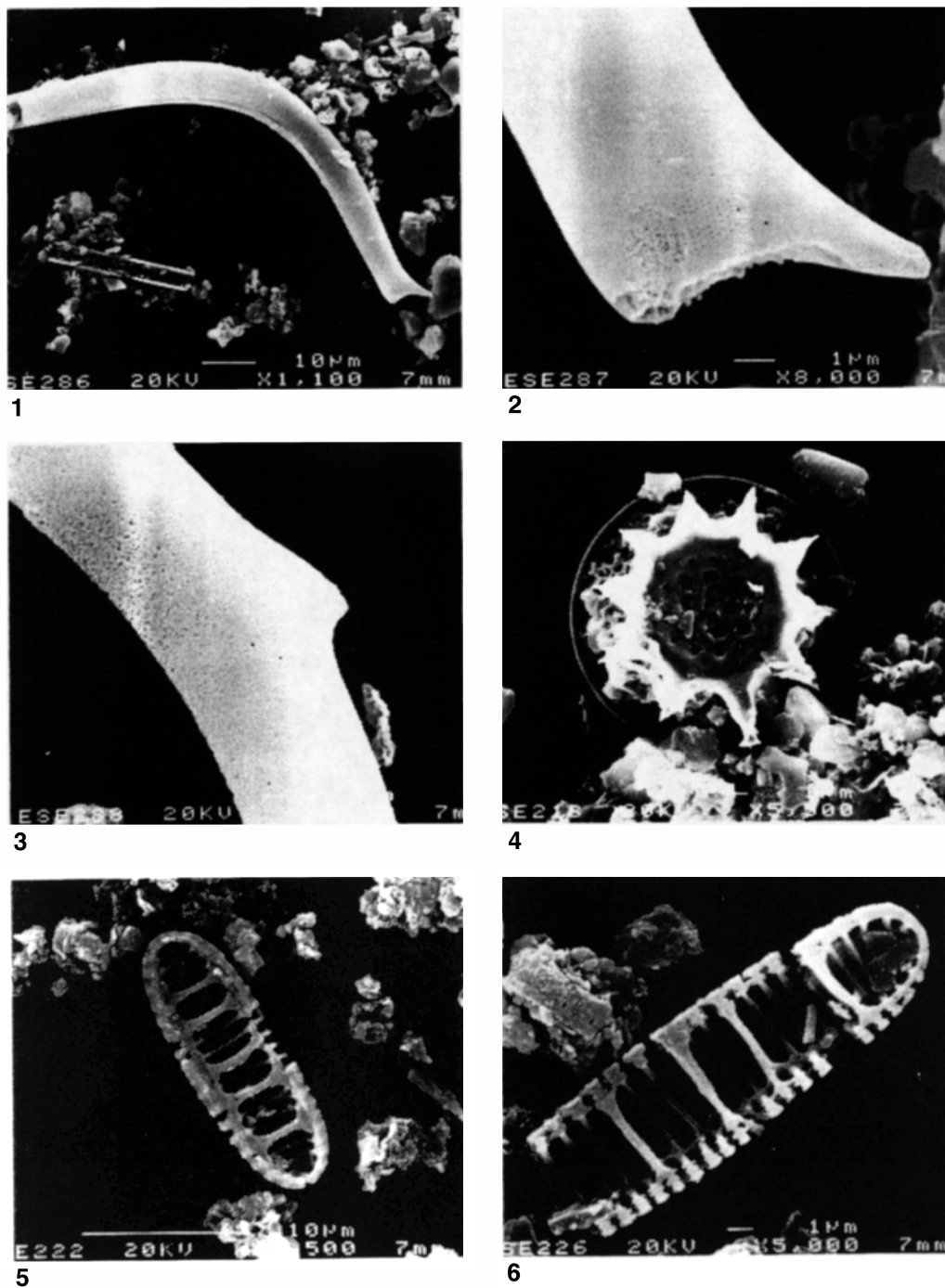


Plate 2. Scale bar = 10 μm , unless otherwise indicated. **1.** *Proboscia curvirostris* Jordan and Priddle, Sample 152-919A-7H-7, 40–41 cm. **2.** Enlarged view of Figure 1, showing dentate end. Scale bar = 1 μm . **3.** *Proboscia curvirostris* Jordan and Priddle. Detail showing a triangular spine, Sample 152-919A-7H-7, 40–41 cm. Scale bar = 1 μm ; magnification is 5,500 \times . **4.** *Thalassiosira nidulus* (Tempère and Brun) Jousé, Sample 152-919A-5H-7, 35–36 cm. Scale bar = 1 μm . **5.** *Neodenticula seminae* (Simonsen and Kanaya) Akiba and Yanagisawa, internal view of valve face, Sample 152-919B-8H-1, 95–96 cm. **6.** *Neodenticula seminae* (Simonsen and Kanaya) Akiba and Yanagisawa, internal view of valve face, Sample 152-919B-8H-1, 95–96 cm. Scale bar = 1 μm .