

31. CHEMOSTRATIGRAPHY OF MADEIRA ABYSSAL PLAIN MIOCENE–PLEISTOCENE TURBIDITES, SITE 950¹

Ian Jarvis,² Jennifer Moreton,² and Martine Gérard³

ABSTRACT

Miocene to Holocene sediments on the Madeira Abyssal Plain (MAP), northeast Atlantic, are dominated by thick-bedded distal mud turbidites. These turbidites record the history of sedimentary source areas and slope failure on the margins of the Canary Basin since ~15 Ma. Major elements and selected trace elements have been determined in 488 turbidite samples collected between 0 and 325 mbsf (Miocene–Pleistocene) at Site 950, on the western MAP. Carbonate and Ti/Al ratio data have been plotted against a detailed sedimentary log to demonstrate the distribution of turbidite chemofacies through the succession. Three major turbidite groups, recognized previously from Quaternary cores, are confirmed to continue through the older sediment record on the plain. Organic-rich, volcanic, and calcareous turbidites are clearly differentiated on chemostratigraphic logs. Organic-rich turbidites dominate both volumetrically and numerically. They have been deposited since the middle Miocene (~15 Ma) and may be subdivided into three geochemically distinct subgroups, the relative importance of which has changed through time. The oldest sediments are Al rich, reflecting more kaolinitic compositions; two K- and Mg-rich subgroups become dominant upward, implying a trend toward more chloritic and illitic clay-mineral assemblages. These changes indicate an increasing importance of northerly source areas on the northwest African continental slope, and/or climatic changes promoting mineralogical shifts in sediments on the margin. The onset of significant volcanic turbidite deposition occurred in the mid–late Miocene, ~14–16 Ma, with the deposition of low-Ti sediments derived from the vicinity of an evolved volcanic source, possibly the slopes of the Canary Islands off Lanzarote or Gomera. A major change toward more basaltic sources occurred in the late Pliocene (~3.5 Ma), possibly associated with the early development of La Palma. Wide ranges in trace-element compositions and a shift toward less Ti-rich compositions indicate the continued existence of multiple sources with increasing volcanic fractionation since that time. Calcareous turbidites have been deposited regularly since the Miocene, but underwent a major decrease in their volcaniclastic component ~3.5 Ma. This is interpreted to indicate the subsidence and draping of the seamount chains to the west of the MAP, which are believed to be the source area for these turbidites.

INTRODUCTION

The Madeira Abyssal Plain (MAP) lies in the deepest part of the Canary Basin at water depths of ~5.4 km. It consists of three linked sub-basins, which when combined occupy an area of approximately 68,000 km² (Fig. 1). The morphology of the MAP is virtually flat, interrupted only by small abyssal hills that rise a few hundred meters above the plain. Quaternary sequences on the MAP represent one of the most comprehensively studied sedimentary records in the deep ocean (Weaver and Kuijpers, 1983; Colley et al., 1984, 1989; Colley and Thomson, 1985, 1992; Kuijpers and Weaver, 1985; Wilson et al., 1985, 1986; Thomson et al., 1986, 1987, 1993; Weaver et al., 1986, 1992, 1994; Weaver, Buckley et al., 1989; Weaver, Thomson, et al., 1989; De Lange et al., 1987, 1989; Jarvis and Higgs, 1987; Kidd et al., 1987; Searle, 1987; Weaver and Rothwell, 1987; Middelburg and De Lange, 1988; Pearce, 1991; De Lange, 1992a, 1992b; Jones et al., 1992; McArthur et al., 1992; Middelburg, 1993; Pearce and Jarvis 1992a, 1992b, 1995; Rothwell et al., 1992; Weaver, 1993; Weaver and Thomson, 1993).

Sediments have been recovered primarily by piston cores to a depth of up to 35 m, representing deposition since ~750 ka. More than 160 cores have been collected from the plain, sequences typically comprising meter-thick, fine-grained distal turbidites, interbedded with thin, centimeter to decimeter pelagic ooze, marl, or clay (Weaver et al., 1986; Weaver and Rothwell, 1987). Thicker turbidites may have a coarser basal facies (Rothwell et al., 1992), comprising thin

sands and silts (Units T_{bcd} of Bouma, 1962), but these units progressively disappear distally, leaving only wispy laminated silts and muds at the base (Units T_{3.5} of Stow and Shanmugam, 1980). Throughout the area, the bulk of each turbidite is typically composed of T_e (T_{6.8}) ungraded muds.

MAP Quaternary sequences have been dated using nannofossil biostratigraphy of the pelagic marls and oozes. The turbidites them-

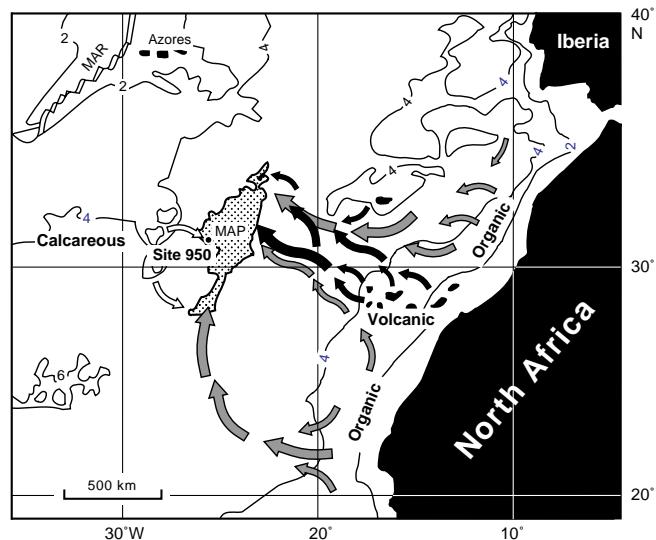


Figure 1. Location map for the Madeira Abyssal Plain (MAP) and ODP Site 950. MAR = Mid-Atlantic Ridge; bathymetry in kilometers; sediment transport pathways from Pearce and Jarvis (1995), gray arrows = organic-rich, black = volcanic, white = calcareous turbidites.

¹Weaver, P.P.E., Schmincke, H.-U., Firth, J.V., and Duffield, W. (Eds.), 1998. *Proc. ODP, Sci. Results*, 157: College Station, TX (Ocean Drilling Program).

²School of Geological Sciences, Kingston University, Penrhyn Road, Kingston upon Thames, Surrey KT1 2EE, United Kingdom. I.Jarvis@kingston.ac.uk

³Laboratoire Petrologie et Mineralogie, Orstom, 72 Route d'Aulnay, 93143 Bondy Cedex, France.

selves exhibit distinctive color and thickness variation, and yield unique mixed coccolith assemblages. By combining these data, individual turbidites have been successfully correlated between cores (Weaver and Kuijpers, 1983; Weaver et al., 1986; Weaver, Buckley et al., 1989; Weaver, Thomson, et al., 1989; Weaver, 1993). In addition, De Lange et al. (1987, 1989) studied the bulk-geochemical analyses of turbidites from two representative MAP cores, and recognized three broad compositional groups: green organic-rich turbidites with 0.3%–2.5% organic carbon (C_{org}); gray “volcanic” turbidites containing volcaniclastic debris and high concentrations of elements such as Ti; and white calcareous turbidites with high $CaCO_3$ (>75%) contents. This work was extended by Pearce (1991) and Pearce and Jarvis (1992a, 1995), who used geochemical data for >500 samples from 23 cores, to demonstrate that individual MAP turbidites have distinctive and consistent geochemical compositions over distances of >500 km, which enable the chemostratigraphic correlation of sequences across the whole plain.

A combination of geochemical and sedimentological evidence (De Lange et al., 1987, 1989; Jones, 1988; McCave and Jones, 1988; Pearce, 1991; Pearce and Jarvis 1992a, 1995; Rothwell et al., 1992), mineralogical data (Weaver and Rothwell, 1987; Pearce and Jarvis, 1992b), and geological arguments (Weaver, Buckley et al., 1989; Weaver, Thomson, et al., 1989; 1992, 1994), were employed by Pearce and Jarvis (1995) to demonstrate that the turbidite compositional groups represented deposition from flows derived from five distinct source areas. Organic-rich turbidites are sourced from the northwest African continental slope (Fig. 1), one subgroup being derived from north of the Canaries, probably off the Moroccan coast; the second subgroup was derived from south of the Canaries, off Western Sahara. Volcanic turbidites originate from the slopes of the Canary Islands, one subgroup coming largely from the western islands, the other being derived from the northern flanks of the central and eastern islands. Calcareous turbidites are derived from the margins of the seamount chains lying to the west of the MAP.

Despite their young age, minimal burial, and deep-water depositional setting, MAP Quaternary turbidites display evidence of significant early diagenetic modification. Many turbidites exhibit a distinctive two-tone coloration with decimeter-thick pale gray to brown-colored tops, and thicker, darker, green lower portions. These ‘bleached’ tops have been produced by the post-depositional oxidation of organic matter in the upper parts of the turbidites by oxygen diffusing into the sediment from bottom water (Colley et al., 1984; Colley and Thomson, 1985; Wilson et al., 1985, 1986; Thomson et al., 1986, 1987, Chap. 32, this volume; Jarvis and Higgs, 1987; Weaver, Thomson, et al., 1989). Oxidation proceeds by the downward migration of a sharply defined redox boundary that remains active until the turbidite is cut off from its oxygen supply by the deposition of a subsequent flow, generally within a few 10 k.y. The color change, which is caused by the destruction of organic matter and oxidation of reduced transition-metal species, is particularly apparent in organic-rich (>0.3% organic carbon) turbidites. Many elements, (As, Cd, Co, Cu, Fe, Mn, Ni, S, Sb, Se, Tl, U, V, and Zn; Thomson et al., 1993, Chap. 32, this volume) are mobilized and relocated in response to the changing redox conditions, but it is important to note that others, particularly Si, Ti, Al, Mg, K, and Zr (Jarvis and Higgs, 1987; Pearce and Jarvis 1995; Thomson et al., Chap. 32, this volume), appear to be unaffected by these early diagenetic processes.

Cores obtained during ODP Leg 157 provide the first opportunity to study the pre-late Quaternary record on the MAP. In this paper, we have applied the methods of De Lange et al. (1987, 1989) and Pearce and Jarvis (1992a, 1995) to sediments obtained at Site 950 to develop a chemostratigraphic framework that will be used to document the history of turbidite sedimentation on the plain since the early middle Miocene (~15 Ma).

MATERIALS

Site 950 is located in the southwestern part of the central MAP (Fig. 1) at 31°9.01'N, 25°36.00'W, at a water depth of 5438 m. This location is only 25 km west-southwest of the site of a 34-m-long giant piston core, MD10, which contains a complete turbidite sequence beginning ~690 ka (isotope Stages 1 through 17). Core MD10 was geochemically characterized by De Lange et al. (1987, 1989) and, along with Core D10688, was the first to be used to demonstrate the application of chemostratigraphic methods to sediment provenance studies. The location of the site on the western margin of the plain ensures that it contains a relatively complete sediment record dominated by very distal facies, with little evidence of basal sand or silt development except in calcareous turbidites, which were derived from the west. These characteristics made the site the most appropriate choice for a chemostratigraphic type section. Shipboard (Schmincke, Weaver, Firth, et al., 1995) and subsequent biostratigraphic data (Howe and Sblendorio-Levy, Chap. 29, this volume) confirm that Site 950 exhibits a thick turbidite record down to the lowest middle Miocene.

An informal lithostratigraphy for Site 950 was proposed by Schmincke, Weaver, Firth, et al. (1995): Unit I (0–306 mbsf; meters below seafloor), comprises thick (decimeter to meter) turbidite muds separated by thin (centimeter) pelagic ooze, marl, or clay interbeds; Unit II (306–333 mbsf) consists of carbonate debris flows; Unit III (333–370 mbsf) is predominantly pelagic red clay, with thin carbonate-rich turbidites and zeolitic ash bands; Unit IV (370–381 mbsf) consists of volcaniclastic sandstones and siltstones. Unit I was divided (Schmincke, Weaver, Firth, et al., 1995) into Subunits Ia (0–150 mbsf) and Ib (150–306 mbsf), on the basis that pelagic intervals below 150 mbsf are clays rather than mixed lithologies (clays, clayey nannofossil mixed sediments, and nannofossil oozes). Following the work of De Lange et al. (1987, 1989) and Pearce and Jarvis (1992a, 1995), shipboard sedimentologists employed a genetic classification for MAP turbidites that included green (organic-rich), gray (volcanic-rich), and white (calcareous) turbidites. Other lithologies recognized (Schmincke, Weaver, Firth, et al., 1995) included gray-green (intermediate-type), light brown or white nannofossil clay, and calcareous turbidites with volcanic clasts.

In total, 488 10-mL sediment plug samples were collected from the mud portions of each turbidite >20 cm thick from 0–325 mbsf (Cores 157-950A-1H to 36X); samples were taken from below redox fronts and above any silty basal facies. Three equally spaced samples were collected from turbidites >1 m thick, and one or two samples from thinner beds. The aim was to adequately characterize the primary geochemical composition of each major turbidite, and to use these data to develop a chemostratigraphic framework, which could be used to document changes in sediment composition and provenance through the Miocene–Pleistocene.

ANALYTICAL METHODS

Sample preparation methods are described elsewhere (Jarvis, 1992). Briefly, unwashed samples were freeze-dried, and ground by hand to a fine powder in an agate mortar and pestle. Homogenized samples were redried overnight at 65°C; 0.250-g subsamples were fused with 1.250 g of lithium metaborate ($LiBO_2$) at 1050°C, and the melts dissolved in dilute HNO_3 . Final solutions were prepared in 250 mL 0.5 M HNO_3 .

Geochemical data were obtained using a Jobin Yvon JY70 Plus ICP-AES at Kingston University. Analytical procedures and operating conditions are listed in Jarvis and Jarvis (1992) and Totland et al. (1992). In this study, nine major elements (Si, Ti, Al, Fe, Mn, Mg, Ca,

Na, and K) and four trace elements (Ba, Cr, Sr, and Zr) were determined. Calibration of the ICP-AES was achieved using nine well-characterized rock reference materials (RRMs) and a procedural blank, selected to cover the range of elemental concentrations expected in samples. Data are reported as weight percent oxides for Si, Ti, Al, Fe, Mg, Na, K, and P, and as $\mu\text{g/g}$ (parts per million) for Ba, Cr, Sr, and Zr. Calcium data are also presented as CaCO_3 , since the bulk of the Ca present occurs in the carbonate fraction.

Analytical precision, reproducibility, and accuracy were determined by replicate analyses of multiple digestions of four different RRMs (Table 1) analyzed on a routine basis with each batch of unknowns. Assessments were based on preparations and determinations made on a number of different days, over a period of 1 yr. Long-term reproducibility was generally better than 2% (with a short-term precision of 0.5% to 1%) for all elements present at or above shale-like concentrations, deteriorating to ~10% for elements at low concentrations in limestones. With reference to published data (Table 1), accuracy is considered generally to lie within the range of the reproducibility.

Data have been presented as absolute concentrations and as values normalized to Al, the latter being used to exclude the masking effects of high but variable CaCO_3 contents in samples. The rationale for this procedure is that biogenic CaCO_3 is relatively pure and free from other elements (Mg and Sr excepted), so differences in the bulk chemistry of the noncarbonate fraction are emphasized by relating values to the Al contents of samples. In the absence of any evidence for pore-water advection, significant import or export of Al ions is considered unlikely for sediments of 50%–60% porosity buried to <400 m.

RESULTS AND DISCUSSION

Chemostratigraphy

De Lange et al. (1987, 1989), Pearce (1991), and Pearce and Jarvis (1995) demonstrated that, although their three major turbidite

groups were defined using a range of sedimentological and geochemical criteria, members of each group could be distinguished using only CaCO_3 and Ti/Al data. Organic-rich turbidite muds (i.e., excluding basal sands and silts) have the lowest CaCO_3 contents, typically 45%–55% (a notable exception is turbidite a_1 , with 18%–32% CaCO_3), and low but constant Ti/Al ratios of ~0.05. Volcanic turbidite muds have intermediate carbonate contents, generally 55%–65%, and high but variable Ti/Al ratios of 0.08–1.5, while calcareous turbidite muds were defined as having >75% CaCO_3 , and yielded marginally higher Ti/Al ratios than the organic-rich group, with values ~0.06.

The above criteria have been applied to turbidite geochemical data obtained from Site 950 (Table 2). Carbonate contents and Ti/Al ratios have been plotted against a detailed lithologic log for the interval 0–325 mbsf (Fig. 2), to illustrate relationships between sediment type, bedding characteristics and geochemical composition. A number of simplifications have been made in the construction of Figure 2. First, only turbidites are shown; pelagic sediments have been omitted for clarity, so beds are defined by the base of each turbidite. Secondly, the geochemical profiles assume constant or little compositional variation within beds. This has been demonstrated to be the case in the Quaternary of the MAP by Jarvis and Higgs, (1987), De Lange et al. (1987, 1989), and Pearce and Jarvis (1992a, 1995). The validity of the approach is confirmed by the close geochemical similarity of multiple samples from most beds (Fig. 2; Table 2); data are commonly within analytical error. Exceptions are generally confined to the lower parts of what must be subtly graded beds. Where possible, only unoxidized sediments from below fossil redox fronts were sampled from Site 950, so the geochemical profiles in Figure 2 reflect primary compositional variation within the sequence. Postdepositional oxidation of organic matter in the bleached tops of organic-rich turbidites leads to the dissolution of carbonate, producing characteristically stepped profiles for CaCO_3 (Jarvis and Higgs, 1987; Thomson et al., Chap. 32, this volume); these tops are not represented in our data.

Table 1. Results obtained for rock reference materials used to assess analytical data quality.

Material	Element	Mean	SD	Reference	Material	Mean	SD	Reference
MAG-1 marine mud*								
Major element (wt%)								
SiO_2	50.6	0.3	50.36		1.14	0.13	1.13	
TiO_2	0.756	0.005	0.751		0.0180	0.0023	0.016	
Al_2O_3	16.3	0.1	16.37		0.369	0.031	0.36	
Fe_2O_3	6.87	0.05	6.80		0.293	0.018	0.28	
MnO	0.0981	0.0009	0.098		0.0153	0.0004	0.016	
MgO	3.06	0.03	3.00		21.3	0.2	21.03	
CaO	1.54	0.07	1.37		30.1	0.2	30.12	
Na_2O	3.84	0.04	3.83		0.04	0.03	0.03	
K_2O	3.55	0.06	3.55		0.11	0.02	0.10	
Trace elements ($\mu\text{g/g}$)								
Ba	479	3	479		<10	4	ND	
Cr	95	2	97		22	5	ND	
Sr	147	1	146		63	1	64	
Zr	128	2	126		17	4	ND	
SCo-1 Cody Shale*								
Major element (wt%)								
SiO_2	62.7	0.8	62.78		0.92	0.17	0.97	
TiO_2	0.621	0.009	0.628		0.0161	0.0034	0.017	
Al_2O_3	13.7	0.2	13.67		0.314	0.055	0.3	
Fe_2O_3	5.06	0.07	5.14		0.180	0.021	0.17	
MnO	0.0512	0.0005	0.053		0.0058	0.0005	0.007	
MgO	2.72	0.04	2.72		2.81	0.03	2.91	
CaO	2.66	0.05	2.62		51.4	0.5	52.12	
Na_2O	0.92	0.03	0.90		0.05	0.04	0.048	
K_2O	2.77	0.02	2.77		0.07	0.01	0.082	
Trace elements ($\mu\text{g/g}$)								
Ba	566	5	570		<10	4	6.6	
Cr	72	1	68		20	5	7.4	
Sr	172	3	174		282	3	284	
Zr	169	3	160		16	4	8	
CCH-1 limestone**								

Notes: Originators: * = U.S. Geological Survey; † = National Institute of Standards and Technology, U.S.A.; ** = University of Liège, Belgium. Reference values from Govindaraju (1994); number of determinations = 14; SD = standard deviation (σ). Total iron expressed as Fe_2O_3 . ND = no data available.

Table 2. Geochemical composition of turbidites, Site 950.

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr				
157-950A-																		
1H-2, 63-65	2.03	20.7	0.908	6.75	3.90	0.110	2.20	30.8	2.78	1.15	332	57	1150	124	55.0	0.152	v	
1H-2, 101-103	2.41	20.5	0.873	6.71	3.82	0.115	2.17	31.1	2.93	1.19	343	40	1160	120	55.4	0.147	v	
1H-2, 136-138	2.76	19.9	0.654	6.60	3.40	0.105	1.91	32.2	2.53	1.17	289	44	1150	104	57.5	0.112	v	
1H-3, 18-20	3.08	16.6	0.308	5.45	2.15	0.0737	1.28	36.1	2.17	1.00	240	47	1160	81	64.3	0.0641	o	Gray - b
1H-3, 113-115	4.03	29.4	0.540	10.3	3.71	0.0683	2.81	22.5	2.17	2.08	408	86	780	98	40.1	0.0595	o	Pale green - c
1H-4, 130-132	5.70	21.8	0.433	7.56	2.76	0.0621	2.06	30.1	2.42	1.58	476	49	1120	83	53.7	0.0648	o	Green - d
1H-5, 64-66	6.54	22.1	0.441	7.38	2.77	0.0618	2.02	29.5	2.09	1.44	474	48	1100	87	52.7	0.0676	o	
1H-6, 63-65	8.03	22.6	0.453	7.31	2.77	0.0653	2.04	29.6	2.35	1.49	442	45	1090	90	52.8	0.0703	o	Green - e
2H-1, 79-81	9.60	24.4	0.366	7.32	2.66	0.0471	1.72	27.1	2.04	1.17	582	60	993	79	48.4	0.0567	o	
2H-2, 79-81	11.01	26.2	0.384	7.51	2.78	0.0501	1.78	27.4	2.12	1.24	563	72	1020	91	49.0	0.0579	o	
2H-3, 79-81	12.49	25.7	0.374	7.33	2.80	0.0508	1.76	27.3	2.20	1.22	559	72	990	90	48.6	0.0578	o	Green - f
2H-4, 98-100	14.19	18.1	0.617	6.03	2.90	0.0949	1.63	34.0	2.84	1.12	306	36	1240	149	60.7	0.116	v	
2H-5, 104-106	15.75	18.2	0.619	6.15	2.82	0.112	1.58	34.1	2.37	1.17	308	43	1240	150	60.9	0.114	v	Gray - g
2H-6, 78-80	16.98	25.2	0.429	8.26	3.02	0.0630	1.92	27.0	1.74	1.38	442	69	990	94	48.2	0.0588	o	
2H-6, 89-91	17.09	25.3	0.434	8.31	3.05	0.0633	1.94	27.2	1.79	1.34	430	74	993	98	48.6	0.0592	o	
2H-6, 96-98	17.16	24.9	0.432	8.40	3.05	0.0637	1.90	27.0	1.79	1.43	421	94	978	98	48.2	0.0583	o	Green - h
2H-7, 4-6	17.73	6.08	0.110	2.05	0.887	0.125	0.628	47.7	1.39	0.34	172	15	1490	31	85.1	0.0609	c	
2H-7, 11-13	17.80	5.41	0.0975	1.83	0.794	0.119	0.651	48.1	1.88	0.41	199	24	1520	24	85.9	0.0604	c	White - j
2H-7, 35-37	18.04	29.2	0.515	11.6	3.95	0.0482	1.92	21.7	1.75	1.64	662	111	842	94	38.8	0.0503	o	
2H-CC, 10-12	18.17	25.1	0.455	10.1	3.54	0.0534	1.70	24.6	1.89	1.35	550	90	933	86	43.9	0.0509	o	Green - k
3H-1, 40-42	18.81	23.8	0.410	8.67	3.35	0.0797	2.25	29.1	1.49	1.90	273	75	1030	66	52.0	0.0536	o	Green - l1
3H-1, 98-100	19.39	25.7	0.466	9.59	3.59	0.0783	2.39	26.8	1.43	2.18	273	60	948	78	47.9	0.0551	o	Green - l3
3H-3, 23-25	20.43	16.2	0.629	5.54	2.99	0.113	1.65	34.9	2.28	1.05	348	47	1260	90	62.3	0.129	v	
3H-3, 50-52	20.70	16.2	0.575	5.60	2.87	0.108	1.62	35.2	2.20	1.08	359	25	1260	84	62.8	0.116	v	Gray - n
3H-4, 25-27	21.79	17.8	0.639	5.90	3.10	0.104	1.56	34.8	2.18	1.09	351	32	1270	123	62.1	0.123	v	
3H-4, 74-76	22.28	18.7	0.654	5.89	3.64	0.110	1.71	33.2	2.56	1.13	348	52	1220	129	59.2	0.126	v	
3H-5, 12-14	23.15	17.3	0.629	5.85	2.70	0.121	1.58	34.9	2.34	1.09	323	33	1270	118	62.3	0.122	v	Gray - o
3H-6, 32-34	24.85	16.9	0.755	5.67	3.05	0.126	1.81	35.6	2.65	0.99	335	49	1310	99	63.6	0.151	v	
3H-6, 70-72	25.23	16.3	0.698	5.47	2.91	0.110	1.71	35.6	2.40	0.90	321	36	1320	95	63.6	0.145	v	Gray - p
3H-7, 42-44	26.45	9.08	0.168	2.90	1.30	0.0633	0.767	45.0	1.40	0.64	197	3	1430	47	80.4	0.0659	c	White - q
3H-7, 140-142	27.43	5.08	0.100	1.62	0.70	0.0792	0.566	48.1	1.20	0.33	132	2	1430	34	85.9	0.0700	c	White - r
3H-CC, 20-22	27.72	25.6	0.466	8.44	3.18	0.0841	2.42	28.1	1.73	1.91	629	54	973	92	50.2	0.0625	o	
4H-1, 18-20	28.09	22.9	0.420	7.56	3.05	0.0582	2.25	28.7	1.93	1.50	615	48	1100	85	51.3	0.0629	o	Green - s
4H-1, 77-79	28.66	28.5	0.512	10.6	3.44	0.0548	2.72	23.2	1.52	2.48	465	69	792	87	41.3	0.0547	o	Green - s1
4H-1, 97-99	28.85	7.55	0.131	2.30	0.799	0.0959	0.656	46.8	1.30	0.44	216	19	1480	35	83.6	0.0643	c	White - s2
4H-1, 142-144	29.29	26.4	0.370	6.90	2.63	0.0481	1.92	27.6	1.86	1.31	630	51	949	92	49.3	0.0607	o	
4H-2, 21-23	29.57	26.1	0.359	6.74	2.49	0.0476	1.85	27.3	1.90	1.19	601	60	943	89	48.7	0.0603	o	
4H-2, 80-82	30.15	26.3	0.365	6.70	2.56	0.0525	1.85	29.0	1.91	1.22	552	62	991	96	51.8	0.0618	o	Green - t
4H-3, 26-28	31.08	28.5	0.412	8.26	3.10	0.0491	1.95	24.5	1.86	1.32	653	74	879	82	43.8	0.0565	o	
4H-3, 66-68	31.47	29.3	0.420	8.22	3.03	0.0515	1.94	24.6	1.92	1.31	615	76	882	93	43.8	0.0579	o	Green - u
4H-3, 147-149	32.26	24.1	0.323	6.60	2.49	0.0611	1.69	29.9	1.97	1.22	739	56	1090	66	53.4	0.0555	o	Green
4H-4, 66-68	32.94	9.10	0.157	2.80	1.03	0.0879	0.759	45.0	1.39	0.52	195	18	1480	40	80.3	0.0635	c	White
4H-4, 112-114	33.39	24.8	0.346	6.94	2.59	0.0567	1.72	28.1	1.90	1.20	713	54	1030	75	50.2	0.0565	o	
4H-5, 50-52	34.25	25.3	0.353	7.04	2.65	0.0572	1.72	28.0	1.88	1.25	719	65	1010	69	50.0	0.0568	o	
4H-5, 122-124	34.95	25.0	0.356	7.10	3.44	0.0587	1.73	28.1	1.77	1.23	685	66	1010	80	50.2	0.0569	o	Green
4H-6, 81-83	36.01	26.8	0.511	9.03	3.80	0.0729	2.41	25.2	1.79	1.97	792	66	937	97	45.0	0.0642	o	Gray
4H-6, 129-131	36.48	32.7	0.535	11.4	4.84	0.0607	2.78	19.7	1.55	2.41	537	92	666	86	35.1	0.0532	o	Gray
4H-7, 18-20	36.86	4.59	0.088	1.58	0.603	0.0803	0.474	49.9	1.05	0.32	159	29	1620	30	89.0	0.0632	c	
4H-7, 39-41	37.06	4.75	0.0908	1.62	0.648	0.0795	0.519	48.3	1.31	0.36	149	28	1580	29	86.2	0.0636	c	White
5H-1, 21-23	37.62	18.9	0.680	6.21	3.16	0.105	1.77	33.5	2.56	1.20	335	39	1290	151	59.7	0.124	v	
5H-1, 55-57	37.95	20.2	0.713	6.44	3.50	0.115	1.89	31.6	2.70	1.17	330	43	1250	152	56.3	0.126	v	
5H-1, 76-78	38.15	18.7	0.626	5.85	3.48	0.107	1.85	31.7	2.81	1.10	307	40	1240	132	56.6	0.121	v	Gray
5H-1, 142-144	38.80	25.9	0.489	9.20	3.30	0.0606	2.34	26.1	1.61	1.88	425	71	951	93	46.6	0.0602	o	
5H-2, 8-10	38.96	25.2	0.480	8.99	3.27	0.0600	2.36	25.6	2.03	1.81	421	66	943	97	45.8	0.0605	o	
5H-2, 47-49	39.34	25.2	0.481	8.91	3.23	0.0613	2.32	25.5	1.82	1.78	405	69	937	97	45.6	0.0611	o	Pale green
5H-3, 39-41	40.72	17.5	0.844	5.57	3.55	0.106	1.94	34.7	2.06	0.89	248	54	1290	104	61.9	0.172	v	
5H-3, 117-119	41.49	17.5	0.834	5.49	3.65	0.101	2.03	33.4	2.72	0.91	242	58	1240	103	59.6	0.172	v	
5H-4, 41-43	42.20	18.6	0.830	5.70	3.81	0.0995	2.04	32.9	2.33	0.94	238	57	1230	106	58.6	0.165	v	Gray
5H-4, 142-144	43.19	9.50	0.172</															

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)									Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr					
6H-3, 8-10	49.60	29.2	0.484	9.45	3.68	0.0442	2.06	23.5	1.59	1.59	566	84	892	104	41.9	0.0580	o		
6H-3, 53-55	50.04	29.2	0.483	9.39	3.70	0.0431	2.08	23.3	1.57	1.70	571	77	881	105	41.5	0.0583	o	Green	
6H-4, 29-31	51.27	22.2	0.388	7.89	2.73	0.0569	2.05	28.3	1.44	1.73	394	66	1030	73	50.4	0.0558	o		
6H-4, 97-99	51.94	23.1	0.404	8.04	2.86	0.0562	2.17	29.1	1.63	1.77	406	60	1040	76	51.9	0.0569	o	Gray	
6H-5, 91-93	53.35	19.7	0.481	7.08	2.70	0.0611	2.09	33.2	1.46	1.47	393	74	1240	88	59.2	0.0769	o	Pale green	
6H-6, 19-21	54.12	4.43	0.0741	1.54	0.460	0.0684	0.461	48.7	1.51	0.36	115	14	1620	27	86.9	0.0546	c	White	
6H-6, 59-61	54.51	31.6	0.554	11.7	4.51	0.0509	2.89	20.7	1.74	2.99	403	84	752	93	36.9	0.0537	o	Gray	
6H-6, 76-78	54.67	11.5	0.203	3.79	1.42	0.0487	0.901	42.9	1.31	0.74	88	37	1420	57	76.6	0.0608	c	White	
6H-CC, 13-15	55.52	4.56	0.0842	1.48	0.583	0.0630	0.471	49.0	0.91	0.31	124	35	1560	31	87.5	0.0643	c		
6H-CC, 42-44	55.80	3.61	0.0725	1.24	0.538	0.0645	0.395	50.5	0.66	0.23	121	19	1640	24	90.0	0.0662	c	White	
6H-CC, 95-96	56.28	28.65	0.497	9.98	3.61	0.0380	2.01	23.8	1.88	1.51	583	103	921	103	42.5	0.0565	o		
7H-1, 22-24	56.63	28.00	0.471	9.51	3.88	0.0408	1.88	23.1	1.66	1.50	581	94	887	101	41.3	0.0561	o		
7H-1, 37-39	56.78	28.00	0.477	9.60	3.52	0.0434	1.89	24.6	1.50	1.45	580	100	937	102	44.0	0.0563	o	Green	
7H-1, 98-100	57.39	29.1	0.497	10.3	3.91	0.0674	2.71	24.3	1.37	2.45	387	82	871	83	43.4	0.0548	o		
7H-1, 108-110	57.49	27.8	0.477	9.84	3.76	0.0564	2.56	23.8	1.33	2.25	370	71	843	84	42.4	0.0549	o	Gray	
7H-1, 128-130	57.69	5.85	0.110	2.00	0.809	0.0631	0.577	48.5	1.05	0.40	81	18	1560	29	86.5	0.0623	c	White	
7H-2, 67-69	58.58	26.5	0.477	9.78	3.38	0.0464	2.56	25.7	1.38	2.42	365	77	919	83	45.9	0.0553	o	Green	
7H-3, 63-65	60.04	14.1	0.229	4.60	1.70	0.0688	1.10	39.9	1.26	0.82	354	46	1470	58	71.2	0.0563	o		
7H-4, 62-64	61.53	15.4	0.240	4.66	1.72	0.0620	1.14	36.3	1.53	0.89	389	41	1300	64	64.8	0.0583	o		
7H-5, 90-92	63.12	15.3	0.241	4.76	1.74	0.0657	1.15	38.5	1.33	0.89	371	45	1380	62	68.8	0.0575	o		
7H-6, 52-54	64.23	16.2	0.247	4.96	1.85	0.0615	1.13	36.6	1.24	0.96	432	46	1310	67	65.4	0.0566	o	Gray	
7H-7, 42-44	65.45	23.1	0.438	8.38	3.01	0.0475	2.11	30.1	1.34	1.77	413	77	1150	90	53.8	0.0593	o		
7H-7, 60-62	65.57	22.4	0.429	8.16	2.95	0.0461	2.04	29.8	1.34	1.71	392	72	1120	88	53.2	0.0596	o	Green	
8H-1, 107-109	66.58	21.8	0.436	7.67	3.22	0.0673	1.95	31.3	1.32	1.68	460	65	1180	106	55.8	0.0644	o	Gray	
8H-2, 81-83	67.80	16.7	0.718	5.36	3.09	0.0740	1.67	35.6	2.14	1.05	215	65	1320	93	63.6	0.152	v		
8H-2, 140-142	68.39	15.7	0.615	5.17	2.77	0.0754	1.53	36.2	1.74	1.08	198	54	1320	81	64.6	0.135	v	Gray	
8H-3, 122-124	69.71	16.3	0.722	5.35	2.89	0.0724	1.69	34.8	1.88	1.04	233	66	1300	78	62.1	0.153	v	Gray	
8H-4, 114-116	71.13	25.0	0.461	8.63	2.93	0.0550	2.41	28.4	1.39	1.86	401	75	1070	96	50.6	0.0605	o		
8H-5, 88-90	72.36	25.9	0.469	8.35	2.91	0.0545	2.41	27.6	1.32	1.76	392	78	1020	108	49.3	0.0636	o		
8H-6, 58-60	73.56	25.3	0.454	7.97	2.75	0.0463	2.30	27.3	1.35	1.69	386	68	993	109	48.6	0.0646	o	Green	
8H-7, 36-38	74.87	17.5	0.276	5.86	2.25	0.0422	1.50	35.4	1.37	1.02	536	61	1340	61	63.2	0.0534	o		
8H-CC, 3-5	75.15	17.8	0.276	5.83	2.26	0.0458	1.47	36.7	1.29	1.04	494	66	1390	62	65.5	0.0536	o		
9H-1, 6-8	75.47	16.5	0.258	5.48	2.11	0.0458	1.40	35.9	1.36	1.00	456	57	1320	59	64.0	0.0534	o	Green	
9H-1, 72-74	76.12	27.3	0.493	10.0	3.75	0.0607	2.75	24.8	1.29	2.45	320	90	899	87	44.3	0.0556	o	Gray	
9H-1, 117-119	76.56	3.43	0.0673	1.12	0.458	0.0592	0.457	50.7	1.06	0.21	86	22	1680	21	90.5	0.0679	c	White	
9H-2, 39-41	77.26	14.4	0.538	4.73	2.36	0.0579	1.36	37.8	1.78	0.96	194	54	1370	86	67.5	0.129	v		
9H-2, 62-64	77.49	12.7	0.487	4.29	2.11	0.0625	1.24	39.7	1.75	0.85	204	34	1470	73	70.8	0.129	v	Gray	
9H-2, 126-128	78.12	29.1	0.497	10.3	3.82	0.0561	2.84	23.6	1.23	2.52	342	75	830	85	42.1	0.0546	o		
9H-3, 26-28	78.61	4.98	0.0898	1.68	0.719	0.0503	0.507	49.4	0.98	0.33	62	12	1570	27	88.1	0.0606	c	White	
9H-3, 112-114	79.45	4.13	0.0819	1.38	0.618	0.0827	0.464	50.9	1.01	0.29	112	15	1720	28	90.8	0.0675	c	White	
9H-4, 18-20	80.00	18.4	0.890	5.52	3.93	0.0890	1.94	34.0	2.04	0.99	188	90	1300	101	60.8	0.183	v		
9H-4, 95-97	80.75	17.9	0.911	5.54	3.75	0.109	1.93	31.0	2.05	1.00	186	91	1200	96	55.3	0.186	v		
9H-5, 57-59	81.85	17.8	0.887	5.87	3.23	0.141	1.82	32.5	1.97	1.07	186	87	1270	96	58.1	0.171	v	Gray	
9H-5, 121-123	82.48	28.0	0.504	10.4	3.69	0.0563	2.80	24.1	1.24	2.30	323	75	864	86	43.1	0.0550	o	Pale green	
9H-5, 139-141	82.65	28.4	0.508	10.4	3.76	0.0561	2.83	23.8	1.26	2.50	335	76	846	79	42.4	0.0555	o		
9H-6, 66-68	83.41	4.77	0.0828	1.56	0.654	0.0841	0.505	49.2	1.04	0.33	136	15	1720	29	87.8	0.0600	c	White	
9H-6, 119-121	83.93	29.7	0.485	9.79	3.87	0.0369	2.13	23.0	1.41	1.69	747	94	872	103	41.1	0.0561	o		
9H-6, 133-135	84.06	29.3	0.481	9.76	4.09	0.0385	2.12	22.8	1.47	1.76	759	80	848	96	40.5	0.0558	o	Dark green	
9H-7, 14-16	84.37	25.3	0.479	9.66	3.31	0.0489	2.55	23.0	1.08	2.30	278	71	812	79	41.0	0.0562	o	Green	
9H-7, 40-42	84.62	26.2	0.423	8.73	3.49	0.0419	1.81	24.8	1.26	1.44	561	81	917	90	44.3	0.0549	o		
9H-CC, 8-10	84.82	26.3	0.430	8.79	3.62	0.0461	1.82	25.7	1.33	1.50	582	82	951	95	45.8	0.0554	o		
10H-1, 89-91	85.80	26.3	0.430	8.61	3.54	0.0541	1.79	26.8	1.42	1.46	553	72	979	100	47.9	0.0566	o	Dark green	
10H-2, 42-44	86.83	27.3	0.504	10.1	3.45	0.0555	2.64	25.1	1.45	2.17	379	77	970	94	44.8	0.0565	o		
10H-2, 78-80	87.19	27.7	0.512	9.71	3.54	0.0551	2.57	24.8	1.23	2.04	359	71	912	100	44.3	0.0597	o	Green	
10H-4, 21-23	88.27	8.28	0.183	3.00	1.24	0.115	0.785	44.4	0.98	0.64	172	25	1510	38	79.3	0.0693	c	White	
10H-4, 79-81	88.85	16.0	0.709	5.37	2.87	0.0911	1.47	36.5	1.79	0.99	208	70	1360	85	65.1	0.150	v		
10H-4, 105-107	89.11	16.0	0.686	5.37	2.89	0.0955	1.47	37.2	1.93	1.01	202	64	1350	85	66.4	0.145	v		
10H-4, 133-135	89.36	15.9	0.657	5.28	2.78	0.0959	1.46	36.8	1.98	1.02	204	69	1370	83	65.6	0.141	v	Pale gray	
10H-5, 138-140	90.91	24.9	0.468	9.26	3.12	0.0477	2.34	28.9	1.10	2.08	370	69	1080	85	51.6				

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr				
10H-7, 98-100	93.51	24.9	0.477	9.16	3.25	0.0605	2.55	28.0	1.45	2.08	330	65	1020	82	50.0	0.0590	o	
10H-7, 124-126	93.77	24.5	0.470	8.93	3.18	0.0602	2.50	27.5	1.20	1.91	323	68	1010	82	49.1	0.0596	o	Green
11H-1, 43-45	94.83	25.2	0.479	8.56	3.13	0.0623	2.51	28.1	1.31	1.86	304	70	1020	90	50.1	0.0634	o	
11H-1, 112-114	95.51	25.4	0.481	8.22	3.03	0.0643	2.44	28.6	1.25	1.72	307	66	1020	94	51.0	0.0663	o	
11H-2, 55-57	96.42	25.7	0.464	7.82	2.96	0.0613	2.29	28.0	1.22	1.82	287	56	970	102	49.9	0.0672	o	Green
11H-3, 30-32	97.65	2.71	0.0534	0.986	0.430	0.0647	0.397	51.5	0.97	0.17	68	15	1700	19	92.0	0.0614	c	
11H-3, 60-62	97.95	2.77	0.0519	0.958	0.420	0.0577	0.376	50.8	0.89	0.17	62	13	1710	18	90.6	0.0614	c	White
11H-4, 11-13	98.94	27.2	0.544	10.0	3.62	0.0408	2.68	25.1	1.22	2.37	320	72	885	90	44.8	0.0616	o	Green
11H-4, 70-72	99.52	2.92	0.0870	1.04	0.506	0.0791	0.390	51.9	0.88	0.23	74	18	1670	23	92.6	0.0952	c	
11H-4, 90-92	99.71	3.09	0.0938	1.08	0.511	0.0849	0.383	51.0	0.85	0.21	75	16	1620	21	91.0	0.0985	c	White
11H-5, 8-10	100.38	19.1	0.449	6.98	2.58	0.0515	1.95	33.4	1.10	1.50	370	57	1310	74	59.5	0.0728	o	Green
11H-5, 61-63	100.90	29.6	0.546	10.9	3.93	0.0435	3.00	22.2	1.19	2.62	331	77	824	93	39.6	0.0567	o	
11H-5, 75-77	101.05	29.3	0.548	11.0	3.98	0.0454	3.03	22.3	1.17	2.52	323	73	843	93	39.8	0.0567	o	Green
11H-6, 15-17	101.90	16.8	0.853	5.71	3.14	0.102	1.67	35.1	1.82	1.08	185	101	1340	90	62.6	0.169	v	
11H-6, 50-52	102.24	16.7	0.825	5.54	3.18	0.103	1.68	35.3	1.85	1.02	169	80	1330	86	62.9	0.169	v	Gray
11H-6, 121-123	102.94	25.1	0.476	9.21	3.24	0.0484	2.66	26.6	1.07	2.12	324	69	934	77	47.5	0.0585	o	Green
11H-7, 29-31	103.51	30.2	0.548	11.1	3.85	0.0403	3.00	22.0	1.27	2.50	307	83	847	99	39.2	0.056	o	Green
11H-CC, 4-6	103.85	26.0	0.480	9.85	3.45	0.0452	2.56	25.9	1.22	2.19	350	65	973	83	46.2	0.0552	o	Green
12H-1, 120-122	105.07	27.4	0.502	9.58	3.45	0.0466	2.67	25.5	1.11	2.10	336	79	962	90	45.4	0.0594	o	
12H-2, 19-21	105.54	26.3	0.481	9.05	3.40	0.0450	2.56	26.4	1.13	1.99	318	75	976	90	47.0	0.0602	o	Green
12H-2, 58-60	105.92	24.7	0.427	9.05	3.31	0.0899	2.86	28.0	1.01	2.12	293	64	947	73	50.0	0.0535	o	Gray
12H-2, 104-106	106.37	20.8	0.457	7.32	3.02	0.0761	2.03	32.1	1.20	1.66	413	50	1150	101	57.3	0.0708	o	
12H-3, 60-62	107.40	24.7	0.457	9.08	3.38	0.0522	2.57	27.8	1.15	2.00	333	71	1030	80	49.7	0.0571	o	Gray
12H-3, 119-121	107.97	24.2	0.447	9.05	3.05	0.0528	2.53	27.6	1.14	1.89	335	70	1020	78	49.2	0.0560	o	
12H-4, 30-32	108.56	24.9	0.457	9.21	3.15	0.0573	2.59	27.9	1.13	2.04	334	78	1040	80	49.7	0.0562	o	
12H-4, 96-98	109.20	26.1	0.471	8.09	3.10	0.0657	2.49	27.8	1.18	1.91	312	58	956	96	49.7	0.0660	o	
12H-5, 26-28	109.98	27.5	0.473	7.40	2.68	0.0642	2.32	27.2	1.20	1.77	302	47	904	105	48.5	0.0724	o	Green
12H-5, 90-92	110.60	27.5	0.493	9.82	3.58	0.0506	2.87	24.6	1.16	2.38	347	75	877	80	43.9	0.0569	o	
12H-5, 110-112	110.79	28.3	0.501	9.85	3.37	0.0501	2.90	24.9	1.14	2.37	357	80	881	88	44.5	0.0576	o	Green
12H-6, 27-29	111.44	31.3	0.577	11.4	3.89	0.0397	3.03	21.1	1.17	2.50	390	95	783	102	37.6	0.0576	o	
12H-6, 100-102	112.15	4.49	0.0939	1.57	0.632	0.0704	0.530	50.6	0.86	0.29	62	20	1610	24	90.3	0.0677	c	Green
12H-7, 30-32	112.93	5.59	0.115	1.96	0.737	0.0637	0.593	49.2	0.89	0.36	67	22	1540	28	87.9	0.0663	c	White
13H-1, 88-90	114.19	30.5	0.497	10.3	3.84	0.0314	2.01	22.4	1.34	1.59	464	110	878	113	40.0	0.0547	o	
13H-2, 49-51	115.30	30.1	0.487	10.1	3.67	0.0315	1.98	22.3	1.34	1.60	461	98	862	111	39.7	0.0547	o	
13H-2, 138-140	116.19	30.0	0.480	9.90	3.56	0.0329	1.93	23.7	1.28	1.54	480	99	909	111	42.2	0.0550	o	
13H-3, 81-83	117.12	29.6	0.465	9.46	3.60	0.0389	1.85	25.3	1.27	1.52	471	99	945	115	45.1	0.0556	o	Green
13H-4, 4-6	117.85	30.9	0.549	11.3	3.78	0.0584	2.98	22.1	1.17	2.51	322	89	821	112	39.4	0.0551	o	Green
13H-4, 52-54	118.33	21.9	1.24	7.34	4.10	0.0964	2.08	30.8	2.36	1.25	157	47	1210	138	55.0	0.191	v	
13H-4, 136-138	119.16	22.0	1.26	7.39	4.42	0.0809	2.17	28.7	2.39	1.17	163	54	1140	129	51.1	0.194	v	Gray
13H-CC, 15-17	119.87	20.6	0.384	7.41	2.51	0.0508	1.87	33.1	1.04	1.60	338	59	1207	67	59.0	0.0588	o	
14H-1, 32-34	123.23	27.0	0.494	9.99	3.79	0.0494	2.63	24.4	1.11	2.20	308	67	897	82	43.5	0.0560	o	Green
14H-1, 75-77	123.66	28.6	0.517	10.3	3.84	0.0497	2.72	24.7	1.12	2.22	293	74	911	85	44.1	0.0570	o	
14H-1, 140-142	124.31	26.4	0.482	9.38	3.26	0.0557	2.52	27.5	1.12	1.98	273	65	997	84	49.1	0.0582	o	Green
14H-2, 59-61	125.00	25.7	0.459	9.26	3.30	0.0503	2.75	28.4	1.07	2.15	307	74	1030	72	50.7	0.0562	o	Green
14H-2, 137-139	125.78	28.0	0.516	9.85	3.78	0.0494	2.69	26.2	1.13	2.13	321	75	972	93	46.7	0.0594	o	
14H-3, 77-79	126.68	27.6	0.505	9.42	3.50	0.0507	2.60	25.5	1.06	2.22	308	61	914	83	45.4	0.0608	o	
14H-4, 10-12	127.51	26.6	0.476	8.80	3.19	0.0562	2.44	28.3	1.11	2.01	280	72	1020	96	50.5	0.0613	o	Green
14H-4, 58-60	127.99	15.8	0.939	5.46	3.52	0.0942	1.64	35.7	2.07	0.93	117	38	1350	89	63.6	0.195	v	
14H-4, 90-92	128.31	16.6	0.965	5.62	3.71	0.0978	1.71	36.7	2.06	0.91	114	50	1400	103	65.5	0.195	v	Gray
14H-5, 15-17	129.06	14.6	0.721	5.04	2.59	0.0920	1.28	38.8	1.73	1.02	195	32	1410	119	69.2	0.162	v	
14H-6, 7-9	130.48	15.0	0.782	5.26	2.61	0.104	1.26	37.6	1.71	1.06	199	28	1380	126	67.1	0.169	v	
14H-6, 134-136	131.73	15.4	0.716	5.22	2.49	0.105	1.21	37.9	1.67	1.01	194	51	1420	119	67.7	0.156	v	Gray
15H-1, 7-9	132.48	22.1	0.356	7.56	3.56	0.0457	1.46	31.2	1.15	1.28	423	87	1170	79	55.7	0.0534	o	
15H-1, 50-52	132.91	21.9	0.352	7.51	3.11	0.0478	1.42	31.4	1.13	1.33	413	83	1170	77	56.1	0.0531	o	
15H-1, 88-90	133.29	21.3	0.347	7.35	2.71	0.0498	1.39	31.3	1.12	1.18	431	70	1150	75	55.8	0.0535	o	Green
15H-1, 130-132	133.71	26.0	0.516	10.1	3.45	0.0659	2.84	25.6	1.12	2.34	279	100	922	85	45.7	0.0581	o	Green
15H-2, 54-55	134.44	45.3	0.943	20.4	6.53	0.0445	1.70	4.48	1.36	2.22	417	150	269	165	8.00	0.0523	o	
15H-2, 58-59	134.48	44.5	0.933	20.0	6.51	0.0450	1.66	4.34	1.34	2.16	409	142	264	159	7.74	0.0528	o	
15H-2, 69-70	134.59	45.7	0.955	20.5														

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr				
15H-5, 118-120	139.59	24.3	0.675	8.74	3.33	0.0926	1.97	28.9	1.35	1.96	292	62	1080	116	51.6	0.0876	v	Gray
15H-6, 34-36	140.25	24.5	0.454	9.11	3.27	0.0841	2.50	28.2	1.04	2.09	308	62	1030	75	50.3	0.0565	o	
15H-6, 67-69	140.58	24.2	0.441	8.88	3.22	0.0908	2.45	28.5	1.05	2.20	304	70	1010	74	50.8	0.0562	o	Gray/green
15H-6, 116-118	141.07	34.8	0.638	13.5	5.09	0.0390	2.12	16.1	1.30	1.92	482	132	670	119	28.8	0.0534	o	
15H-CC, 9-11	141.23	34.8	0.640	13.5	6.11	0.0396	2.13	16.2	1.32	1.95	491	129	675	119	28.9	0.0537	o	
16H-1, 10-12	142.01	35.4	0.632	13.0	4.90	0.0382	2.14	16.8	1.30	1.90	467	121	701	125	30.0	0.0551	o	
16H-1, 66-68	142.55	34.7	0.604	12.3	5.97	0.0373	2.02	17.0	1.28	1.96	458	114	684	129	30.4	0.0556	o	
16H-1, 117-119	143.04	34.4	0.600	12.2	4.51	0.0372	2.00	17.9	1.27	1.97	455	111	711	132	32.0	0.0557	o	Dark green
16H-2, 28-30	143.62	27.9	0.528	10.3	3.76	0.0691	3.06	23.0	1.12	2.47	274	68	815	85	41.1	0.0583	o	
16H-2, 45-47	143.79	28.1	0.532	10.4	3.86	0.0665	3.12	23.1	1.15	2.52	273	67	808	87	41.3	0.0582	o	Green
16H-2, 102-104	144.32	28.3	0.510	10.3	3.96	0.106	3.12	24.4	1.05	2.44	263	73	831	83	43.6	0.0563	o	Gray
16H-3, 8-10	144.88	28.6	0.530	10.9	3.74	0.0633	2.60	23.0	1.14	2.38	327	71	863	93	41.1	0.0551	o	
16H-3, 68-70	145.46	28.3	0.523	10.5	4.01	0.0639	2.53	22.7	1.10	2.35	308	74	830	94	40.5	0.0564	o	
16H-3, 117-119	145.93	28.8	0.533	10.3	3.71	0.0676	2.50	24.2	1.14	2.28	315	64	901	97	43.2	0.0584	o	Gray
16H-4, 6-8	146.31	46.3	0.988	21.4	6.68	0.0387	1.71	3.93	1.30	2.22	358	173	247	171	7.02	0.0522	o	
16H-4, 29-31	146.53	44.6	0.956	20.8	7.07	0.0344	1.64	3.77	1.29	2.18	356	164	239	164	6.73	0.0520	o	
16H-4, 46-48	146.69	46.1	0.992	21.5	6.64	0.0385	1.68	4.09	1.28	2.21	359	170	251	173	7.29	0.0524	o	Dark green
16H-4, 117-119	147.38	5.24	0.126	1.87	0.824	0.0990	0.542	49.9	0.95	0.37	38	18	1460	34	89.1	0.0761	c	
16H-5, 2-5	147.72	2.04	0.0725	0.714	0.388	0.0635	0.464	51.0	1.11	0.14	16	3	1060	20	91.0	0.115	c	White
16H-5, 21-23	147.90	30.3	0.577	11.8	4.46	0.0700	3.32	20.7	1.09	2.97	244	76	727	131	36.9	0.0555	o	Gray
16H-5, 87-89	148.54	19.3	0.907	6.54	3.42	0.102	1.52	31.0	1.98	1.46	262	19	1170	194	55.3	0.157	v	
16H-5, 117-119	148.83	19.7	0.914	6.79	3.33	0.113	1.58	32.8	2.05	1.50	261	20	1240	178	58.6	0.152	v	Gray
16H-6, 2-4	149.17	30.8	0.564	11.8	3.88	0.0684	2.86	21.4	1.09	2.66	311	85	771	99	38.3	0.0544	o	Gray
16H-6, 60-62	149.73	20.4	0.384	7.73	2.63	0.0690	1.74	32.6	1.06	1.72	311	54	1170	74	58.2	0.0562	o	
16H-6, 70-72	149.83	20.3	0.384	7.70	2.62	0.0744	1.77	32.2	1.03	1.60	301	46	1170	75	57.5	0.0565	o	Gray
16H-6, 136-138	150.46	18.8	0.415	7.09	2.69	0.0858	1.90	33.3	1.09	1.58	260	51	1270	77	59.4	0.0663	o	
16H-7, 71-73	151.20	19.3	0.444	6.92	2.65	0.103	1.84	34.2	1.09	1.62	263	48	1300	84	61.0	0.0727	o	
16H-CC, 11-13	151.33	18.6	0.433	6.71	2.60	0.104	1.81	33.3	1.09	1.45	249	52	1280	83	59.5	0.0731	o	Gray/green
17H-1, 48-50	151.59	30.4	0.550	11.3	4.06	0.0554	2.93	21.2	1.16	2.51	315	84	772	100	37.8	0.0551	o	
17H-1, 69-71	151.78	30.7	0.564	11.3	4.11	0.0568	2.92	22.1	1.14	2.51	299	82	810	103	39.4	0.0564	o	Green
17H-2, 50-52	152.06	23.0	0.426	8.63	3.09	0.0587	2.52	29.4	0.98	2.17	262	73	1010	70	52.4	0.0559	o	
17H-2, 63-65	152.17	24.6	0.459	8.81	3.32	0.0638	2.74	29.4	1.00	2.12	253	68	1020	80	52.4	0.0590	o	Green
17H-3, 26-28	152.65	2.97	0.0721	1.10	0.570	0.115	0.367	50.8	0.79	0.21	32	20	1570	20	90.7	0.0744	c	
17H-3, 35-37	152.73	3.54	0.0826	1.20	0.699	0.108	0.421	51.1	0.88	0.23	33	15	1610	24	91.2	0.0783	c	White
17H-3, 122-123	153.43	41.5	0.879	18.3	5.60	0.023	1.70	9.06	1.17	1.98	362	137	429	148	16.2	0.0544	o	
17H-3, 130-131	153.50	40.8	0.863	18.0	5.57	0.0233	1.67	8.97	1.17	1.93	355	135	425	145	16.0	0.0543	o	
17H-3, 138-139	153.57	41.4	0.873	18.2	5.69	0.0236	1.70	9.12	1.35	2.00	373	137	432	154	16.3	0.0542	o	
17H-4, 14-16	153.78	44.1	0.931	19.2	6.35	0.0256	1.79	8.27	1.33	2.11	378	168	415	165	14.8	0.0549	o	Dark green
17H-4, 60-62	154.15	27.4	0.497	10.1	3.55	0.0542	2.72	25.0	1.08	2.22	290	78	916	93	44.7	0.0557	o	
17H-CC, 22-24	154.35	27.6	0.492	9.51	3.40	0.0559	2.72	24.3	1.06	2.13	287	83	857	97	43.4	0.0586	o	Green
18X-1, 111-113	154.65	28.1	0.503	9.59	3.97	0.0574	2.78	24.8	1.07	2.31	317	70	857	99	44.3	0.0594	o	
18X-1, 128-130	154.82	28.1	0.501	9.42	3.52	0.0585	2.79	25.1	1.16	2.19	297	65	861	101	44.8	0.0602	o	Pale green
18X-1, 143-145	154.97	15.4	0.283	5.83	1.97	0.0832	1.35	38.4	0.96	1.11	335	52	1360	58	68.6	0.0549	o	Gray
18X-2, 47-49	155.50	28.1	0.506	10.3	3.68	0.0589	2.86	24.0	1.07	2.50	306	73	838	95	42.8	0.0555	o	
18X-2, 94-96	155.97	27.6	0.496	9.88	4.64	0.0622	2.79	23.6	1.07	2.36	290	68	818	95	42.1	0.0569	o	
18X-2, 118-120	156.21	28.0	0.503	9.89	3.49	0.0632	2.77	24.9	1.06	2.33	300	68	860	98	44.4	0.0576	o	Green
18X-CC, 10-12	157.01	28.0	0.520	10.3	3.52	0.0473	2.60	24.4	1.09	2.27	300	75	858	98	43.5	0.0571	o	
18X-CC, 37-39	157.28	27.4	0.508	9.98	3.52	0.0486	2.57	24.4	1.06	2.30	288	68	849	96	43.5	0.0577	o	Green
19X-1, 70-72	160.84	13.8	0.258	5.09	1.74	0.0585	1.20	40.0	0.93	0.99	245	40	1410	54	71.4	0.0573	o	
19X-1, 87-89	161.01	15.3	0.279	4.99	1.76	0.0620	1.29	39.3	1.04	0.99	236	44	1320	68	70.1	0.0635	o	Pale green
19X-1, 135-137	161.49	8.69	0.321	3.15	1.39	0.0822	0.90	45.1	1.15	0.68	136	27	1560	59	80.5	0.115	c	
19X-2, 15-17	161.72	8.79	0.330	3.23	1.31	0.0849	0.89	44.7	1.26	0.71	140	22	1560	55	79.8	0.116	c	
19X-2, 40-42	161.97	9.43	0.362	3.45	1.36	0.0838	0.94	44.1	1.40	0.74	142	24	1550	55	78.8	0.119	c	Gray
19X-2, 66-68	162.23	27.4	0.503	10.3	3.73	0.0799	3.58	24.2	1.02	2.58	246	75	805	88	43.2	0.0554	o	Gray
19X-2, 103-105	162.60	3.13	0.0875	1.07	0.491	0.0754	0.394	51.2	0.80	0.26	46	9	1410	22	91.4	0.0923	c	White
19X-2, 141-143	162.98	3.02	0.0647	1.06	0.460	0.0767	0.384	51.7	0.80	0.25	43	8	1560	21	92.2	0.0689	c	White
19X-3, 38-40	163.45	3.44	0.0855	1.18	0.511	0.0685	0.412	50.9	0.80	0.27	46	7	1510	20	90.8	0.0823	c	White
19X-3, 118-120	164.25	23.8	0.437	8.63	3.08	0.0551	2.66	28.7	1.08	1.91	268	57	1010	84	51.3	0.0574	o	
19X-4, 39-41	164.96	24.3	0.440</td															

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements ($\mu\text{g/g}$)					CaCO_3 (wt%)	Ti/Al	Group	Color
		SiO_2	TiO_2	Al_2O_3	Fe_2O_3	MnO	MgO	CaO	Na_2O	K_2O	Ba	Cr	Sr	Zr				
19X-6, 62-64	168.19	22.6	0.431	7.13	2.90	0.0870	2.87	30.8	1.06	1.71	245	48	1010	75	55.0	0.0685	o	Gray/green
19X-CC, 9-11	168.47	23.1	0.427	8.60	2.93	0.0682	2.58	29.0	1.01	1.91	275	60	1030	81	51.7	0.0562	o	
19X-CC, 43-45	168.81	22.3	0.412	8.31	5.05	0.0704	2.42	27.9	1.02	1.85	258	57	995	80	49.9	0.0562	o	
20X-1, 20-22	169.97	23.8	0.444	8.38	3.60	0.0743	2.49	28.6	1.13	1.92	261	56	1010	86	51.0	0.0600	o	
20X-1, 61-63	170.34	23.8	0.448	8.47	2.92	0.0714	2.37	29.1	1.08	1.91	252	57	1030	91	52.0	0.0598	o	
20X-1, 93-95	170.62	15.8	0.496	5.55	2.28	0.130	1.21	38.5	1.27	1.16	184	29	1350	108	68.8	0.101	v	Green
20X-1, 126-128	170.91	13.5	0.781	4.78	2.24	0.108	1.24	39.2	1.67	0.99	206	25	1360	99	70.0	0.185	v	Gray
20X-2, 97-99	171.96	26.3	0.481	9.61	3.20	0.0533	2.22	27.3	1.15	2.03	265	62	988	92	48.8	0.0568	o	Gray
20X-2, 115-117	172.12	26.3	0.490	10.0	3.34	0.0622	2.28	29.2	1.10	2.12	259	68	1060	108	52.2	0.0553	o	Green
20X-3, 5-7	172.48	23.2	1.31	7.98	4.91	0.115	2.25	28.2	2.38	1.26	204	44	1070	167	50.4	0.185	v	
20X-3, 23-25	172.64	23.3	1.32	8.02	4.85	0.115	2.25	28.2	2.40	1.30	206	42	1070	170	50.4	0.186	v	
20X-3, 46-48	172.84	24.2	1.37	8.43	4.83	0.0978	2.27	26.6	2.48	1.39	225	39	1030	168	47.4	0.184	v	Dark gray
20X-3, 106-108	173.38	3.16	0.088	1.03	0.470	0.0860	0.600	50.8	1.03	0.24	41	10	1100	23	90.7	0.0966	c	Pale gray
20X-4, 41-43	174.13	8.94	0.288	3.11	1.40	0.101	0.910	44.6	1.21	0.67	162	24	1570	54	79.6	0.105	c	Gray
20X-4, 64-66	174.33	8.81	0.290	3.05	1.40	0.109	0.870	45.1	1.19	0.67	161	28	1600	52	80.5	0.108	c	Gray
20X-4, 109-111	174.73	17.9	0.309	6.43	2.16	0.0660	1.25	35.6	1.14	1.06	356	50	1290	61	63.6	0.0544	o	Gray
20X-5, 8-10	175.17	38.4	0.763	16.1	5.40	0.0277	1.82	11.2	1.35	1.86	530	129	502	137	20.0	0.0538	o	
20X-5, 49-51	175.52	39.3	0.774	16.2	5.38	0.0281	1.85	11.5	1.29	1.90	543	130	510	140	20.5	0.0542	o	
20X-5, 110-112	176.07	40.0	0.793	16.6	5.30	0.0287	1.84	11.0	1.33	1.90	506	134	496	144	19.7	0.0543	o	Dark green
20X-6, 39-41	176.77	21.9	0.409	8.10	2.92	0.0703	2.13	31.4	1.06	1.77	285	60	1110	80	56.0	0.0571	o	
20X-6, 112-114	177.42	21.9	0.406	8.09	2.86	0.0741	2.17	31.0	1.07	1.73	276	61	1100	78	55.4	0.0568	o	
20X-CC, 21-23	178.24	23.9	0.436	7.79	2.83	0.0710	2.20	29.8	1.09	1.71	265	55	1030	91	53.1	0.0634	o	Green
21X-1, 81-83	179.27	15.5	0.283	5.60	1.87	0.0613	1.28	39.2	1.04	0.97	300	52	1380	61	70.0	0.0572	o	
21X-1, 109-110	179.55	15.4	0.280	5.53	1.84	0.0628	1.29	38.1	1.10	1.02	284	53	1330	62	68.1	0.0574	o	Pale green
21X-2, 39-41	180.33	23.9	0.439	8.45	3.05	0.0778	2.37	29.4	1.03	1.76	319	70	1020	85	52.4	0.0589	o	
21X-2, 107-109	181.00	24.1	0.444	8.38	3.29	0.0853	2.37	29.0	1.06	1.84	309	65	1010	86	51.8	0.0601	o	
21X-3, 40-42	181.84	24.8	0.446	8.16	2.95	0.0875	2.28	29.2	1.11	1.74	315	66	1000	91	52.1	0.0619	o	Green
21X-3, 124-126	182.67	32.5	0.564	11.6	3.99	0.0374	1.98	18.9	1.16	1.73	470	107	709	124	33.7	0.0552	o	
21X-4, 6-8	182.99	33.1	0.575	11.6	4.18	0.0410	2.00	19.4	1.20	1.66	460	108	726	124	34.5	0.0562	o	
21X-4, 62-64	183.55	34.6	0.595	12.1	4.05	0.0446	2.06	20.3	1.26	1.76	467	109	762	129	36.3	0.0556	o	Dark green
21X-4, 119-121	184.11	28.4	0.504	10.5	3.53	0.0622	2.38	25.2	1.12	2.27	342	77	900	90	44.9	0.0543	o	
21X-4, 145-147	184.37	30.5	0.536	10.9	3.89	0.0656	2.49	26.2	1.18	2.43	355	79	925	98	46.8	0.0556	o	Green
21X-5, 19-21	184.61	26.6	0.516	9.79	3.30	0.0696	2.92	26.6	1.16	2.38	275	76	888	87	47.5	0.0597	o	Green
21X-5, 88-90	185.29	17.1	0.284	5.95	2.26	0.0791	1.24	38.1	0.90	1.10	323	55	1310	62	68.1	0.0541	o	
21X-5, 103-105	185.44	19.3	0.328	6.75	2.55	0.0647	1.37	36.8	1.01	1.21	357	59	1310	71	65.6	0.0551	o	Pale green
21X-5, 125-127	185.66	5.31	0.125	1.84	0.766	0.134	0.533	50.2	1.09	0.39	60	26	1500	34	89.5	0.0770	c	White
21X-6, 4-6	185.95	3.74	0.098	1.29	0.573	0.139	0.437	52.1	0.98	0.27	53	25	1520	29	93.0	0.0859	c	White
21X-6, 89-91	186.79	13.8	0.277	5.02	1.84	0.0726	1.12	40.6	1.12	0.98	243	48	1430	64	72.4	0.0625	o	
21X-6, 142-144	187.32	14.1	0.281	5.07	1.84	0.0835	1.12	41.3	1.05	1.01	243	49	1440	66	73.7	0.0629	o	
21X-CC, 7-9	187.97	15.0	0.300	5.11	1.79	0.0856	1.11	40.2	1.03	1.00	245	46	1410	70	71.7	0.0664	o	Gray
22X-1, 62-64	188.81	8.91	0.203	3.05	1.26	0.141	0.738	46.0	0.92	0.60	53	31	1360	46	82.0	0.0755	c	
22X-1, 84-86	189.02	8.58	0.217	2.88	1.18	0.143	0.788	46.4	1.07	0.58	56	31	1270	47	82.8	0.0850	c	White
22X-2, 3-5	189.69	26.9	0.481	9.62	3.30	0.0466	1.68	27.6	1.09	1.37	372	89	1010	104	49.3	0.0567	o	
22X-2, 48-50	190.12	25.8	0.457	9.16	3.28	0.0497	1.64	27.2	1.08	1.34	357	84	992	99	48.6	0.0565	o	
22X-2, 87-89	190.50	26.4	0.469	9.40	3.07	0.0542	1.66	28.7	1.14	1.37	363	91	1050	102	51.2	0.0566	o	Dark green
22X-3, 26-28	191.37	20.5	0.401	7.72	2.72	0.0728	1.89	34.4	1.10	1.58	244	57	1230	79	61.5	0.0588	o	
22X-3, 98-100	192.07	19.9	0.385	7.47	2.65	0.0692	1.86	33.4	0.99	1.60	243	57	1180	77	59.6	0.0584	o	
22X-4, 37-39	192.93	23.7	0.433	8.60	3.06	0.0785	2.10	30.7	1.00	1.83	253	65	1090	82	54.7	0.0570	o	Gray
22X-5, 58-60	194.59	20.7	0.401	7.57	2.73	0.0684	1.94	33.6	1.06	1.60	234	58	1180	81	60.0	0.0601	o	
22X-6, 92-94	196.39	20.7	0.402	7.53	2.73	0.0703	1.92	33.4	1.02	1.60	245	58	1180	82	59.5	0.0605	o	
22X-CC, 24-26	197.66	23.8	0.465	7.63	2.92	0.0820	1.99	32.1	1.06	1.66	259	59	1110	98	57.3	0.0691	o	
23X-1, 14-16	197.95	25.6	0.501	7.67	2.96	0.0847	2.01	30.6	1.08	1.65	267	58	1060	110	54.6	0.0740	o	
23X-1, 36-38	198.17	27.8	0.530	7.52	2.93	0.0819	2.00	28.7	1.15	1.59	273	55	978	127	51.2	0.0799	o	Gray
23X-2, 8-10	199.35	4.23	0.119	1.46	0.671	0.0919	0.451	50.7	0.99	0.34	37	27	1430	33	90.5	0.0925	c	
23X-2, 55-57	199.82	4.53	0.158	1.51	0.742	0.0785	0.511	50.4	0.98	0.34	49	29	1260	37	90.0	0.119	c	White
23X-2, 142-144	200.69	13.9	0.375	4.92	2.19	0.0833	1.21	41.0	0.93	1.03	173	42	1400	76	73.2	0.0864	v	Gray
23X-3, 26-28	201.03	12.9	0.381	4.68	1.90	0.0893	1.14	40.8	1.12	0.97	178	43	1380	78	72.7	0.0924	v	
23X-3, 119-121	201.96	26.3	0.485	9.95	3.12	0.0481	1.64	27.7	1.06	1.44	311							

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr				
24X-4, 86-88	212.83	29.3	0.983	9.91	4.63	0.178	2.23	24.1	1.59	2.24	228	55	860	218	43.1	0.112	v	
24X-5, 15-17	213.61	31.4	0.995	10.6	4.84	0.149	2.34	22.1	1.61	2.37	206	57	812	227	39.5	0.106	v	Gray
24X-5, 112-114	214.58	16.8	0.454	6.09	2.49	0.231	1.20	37.7	1.12	1.47	171	38	1250	133	67.3	0.0846	v	Pale gray
24X-6, 17-19	215.11	25.1	0.513	7.93	2.61	0.0810	1.67	30.6	0.97	1.84	270	55	1050	108	54.7	0.0733	o	Gray
24X-6, 39-41	215.33	4.13	0.119	1.51	0.672	0.189	0.439	49.5	0.85	0.32	33	17	1440	30	88.3	0.0893	c	
24X-6, 71-73	215.65	5.76	0.161	2.10	0.942	0.184	0.542	49.2	0.85	0.40	39	19	1460	39	87.8	0.0867	c	White
24X-6, 111-113	216.05	19.1	0.694	6.96	3.21	0.111	1.46	33.8	1.36	1.50	125	38	1210	140	60.3	0.113	v	Gray
24X-CC, 20-22	217.00	26.3	1.45	8.70	6.70	0.185	3.12	24.2	1.94	1.78	218	61	807	185	43.2	0.189	v	Gray
25X-1, 49-51	217.62	34.9	0.625	13.0	3.81	0.0454	2.09	18.9	1.17	1.90	372	115	729	114	33.8	0.0546	o	Green
25X-1, 89-91	218.02	35.4	0.645	13.6	4.56	0.0944	3.02	17.3	1.09	2.87	267	91	644	114	31.0	0.0539	o	Gray
25X-1, 123-125	218.36	17.9	0.483	6.57	2.82	0.170	1.51	35.6	1.00	1.48	175	49	1210	101	63.6	0.0832	v	
25X-1, 146-148	218.59	16.1	0.460	5.87	2.56	0.191	1.39	38.4	1.02	1.36	160	46	1290	99	68.5	0.0888	v	Gray
25X-2, 54-56	219.17	16.5	0.576	5.61	3.02	0.139	1.60	37.9	1.02	1.27	139	44	1130	95	67.7	0.117	v	
25X-2, 111-113	219.74	15.7	0.567	5.26	2.88	0.141	1.56	38.2	0.96	1.13	149	41	1140	92	68.1	0.122	v	
25X-3, 5-7	220.18	15.9	0.577	5.32	2.74	0.143	1.56	37.4	1.06	1.19	165	44	1110	93	66.8	0.123	v	Gray
25X-3, 90-92	221.03	19.5	0.358	7.40	2.48	0.0932	1.37	34.4	0.88	1.37	245	65	1260	75	61.4	0.0547	o	Pale green
25X-4, 9-11	221.72	34.3	0.614	12.4	4.38	0.0729	2.62	19.3	1.12	2.70	275	80	711	115	34.4	0.0562	o	
25X-4, 37-39	222.00	33.1	0.590	12.0	4.26	0.0702	2.54	19.4	1.23	2.57	280	80	727	111	34.7	0.0555	o	Gray
25X-4, 80-82	222.43	33.6	0.568	11.8	4.21	0.0376	2.01	18.9	1.16	1.89	410	112	730	111	33.7	0.0544	o	
25X-4, 107-109	222.70	31.7	0.535	11.1	7.96	0.0354	1.90	17.4	1.09	1.70	415	109	672	109	31.1	0.0547	o	Dark green
25X-5, 13-15	223.26	4.44	0.126	1.48	0.746	0.163	0.463	50.3	0.86	0.33	49	23	1530	29	89.7	0.0961	c	
25X-5, 30-32	223.43	4.66	0.132	1.54	0.790	0.161	0.471	49.8	0.83	0.33	49	21	1520	29	88.8	0.0970	c	White
25X-5, 97-100	224.10	41.2	0.738	15.7	5.31	0.0295	2.21	10.7	1.25	2.16	551	132	466	129	19.1	0.0532	o	
25X-5, 134-137	224.47	41.0	0.728	15.3	6.13	0.0315	2.19	11.0	1.28	2.08	527	130	473	130	19.6	0.0540	o	
25X-6, 25-27	224.88	40.5	0.707	14.6	6.80	0.0318	2.09	10.8	1.22	2.09	489	126	456	134	19.3	0.0548	o	
25X-6, 74-76	225.37	41.3	0.726	15.2	5.33	0.0335	2.06	11.2	1.17	2.17	482	128	471	135	20.0	0.0543	o	Dark green
25X-6, 135-137	225.98	34.8	0.626	13.2	4.50	0.0944	2.65	18.2	1.11	2.94	314	88	681	108	32.4	0.0537	o	Gray/brown
25X-CC, 29-31	226.42	26.1	0.470	9.40	3.59	0.141	1.81	27.4	1.04	2.06	238	66	966	101	48.9	0.0567	o	
26X-1, 36-38	227.05	25.7	0.463	9.11	3.59	0.127	1.74	28.7	1.05	2.12	247	62	1020	101	51.3	0.0576	o	
26X-1, 121-123	227.90	24.3	0.448	8.75	3.07	0.0962	1.65	29.8	1.03	1.92	258	61	1080	99	53.3	0.0580	o	
26X-2, 22-24	228.41	24.8	0.457	8.99	3.13	0.101	1.67	30.2	1.04	1.94	256	60	1090	100	53.9	0.0577	o	Pale gray
26X-2, 141-143	229.60	26.6	0.500	9.79	3.57	0.0974	2.65	26.2	0.97	2.45	256	68	901	86	46.8	0.0578	o	
26X-3, 17-19	229.86	26.7	0.500	9.72	3.59	0.104	2.69	26.4	1.01	2.43	255	70	897	88	47.0	0.0583	o	
26X-3, 61-63	230.30	26.9	0.502	9.68	3.68	0.114	2.71	26.3	1.00	2.44	252	69	892	89	46.9	0.0587	o	Pale gray
26X-3, 134-135	231.02	37.9	0.688	14.2	4.87	0.0790	2.16	15.0	1.17	2.03	432	127	595	119	26.7	0.0549	o	
26X-3, 140-141	231.08	38.5	0.696	14.5	4.99	0.0755	2.20	14.3	1.18	2.04	414	130	572	124	25.5	0.0545	o	
26X-3, 146-147	231.14	37.6	0.682	14.1	4.98	0.0728	2.14	15.0	1.19	2.00	416	129	597	124	26.7	0.0547	o	Dark green
26X-4, 79-81	231.98	10.8	0.257	3.83	1.69	0.239	0.908	43.0	1.18	0.73	52	32	1240	47	76.8	0.0762	c	White
26X-5, 19-21	232.88	34.0	0.611	13.0	5.43	0.206	2.60	18.7	1.10	3.05	270	84	678	105	33.4	0.0533	o	
26X-5, 64-66	233.33	29.6	0.543	11.4	3.69	0.119	2.30	23.6	1.01	2.59	295	74	829	95	42.2	0.0540	o	Pale brown
26X-5, 145-147	234.14	38.3	0.712	15.0	4.73	0.0850	2.47	14.3	1.22	3.29	348	88	553	138	25.6	0.0538	o	
26X-6, 28-30	234.47	38.0	0.710	14.9	5.14	0.0959	2.45	14.8	1.16	3.25	343	89	569	137	26.4	0.0541	o	Green
26X-6, 88-90	235.07	51.4	0.956	19.1	7.53	0.103	3.36	1.09	1.47	429	346	109	151	191	1.94	0.0569	o	Green
26X-CC, 3-5	235.72	37.6	0.693	14.0	4.77	0.101	2.47	15.0	1.32	3.03	349	83	558	142	26.7	0.0561	o	
27X-1, 12-14	236.63	38.5	0.717	14.3	4.91	0.101	2.49	15.3	1.22	2.77	346	87	577	151	27.3	0.0567	o	
27X-1, 42-44	236.92	37.9	0.712	14.5	4.99	0.123	2.49	15.2	1.20	2.91	328	88	576	150	27.1	0.0557	o	Green
27X-1, 103-105	237.51	50.4	1.07	22.9	6.87	0.0355	1.74	0.67	1.30	2.13	470	160	136	179	1.20	0.0529	o	
27X-1, 135-137	237.82	49.0	1.04	22.3	6.70	0.0344	1.68	1.67	1.25	2.11	462	156	169	176	2.98	0.0528	o	Dark green
27X-2, 21-23	238.15	7.34	0.19	2.60	1.29	0.324	0.639	46.8	0.97	0.55	32	35	1430	42	83.5	0.0828	c	White
27X-2, 80-82	238.73	39.6	0.731	15.5	6.35	0.404	2.87	12.2	1.23	3.05	414	96	479	132	21.7	0.0534	o	Brown
27X-3, 14-16	239.55	33.5	0.599	12.9	5.32	0.377	2.31	17.8	1.10	2.81	340	83	649	118	31.8	0.0527	o	Brown
27X-3, 60-62	239.99	37.2	0.644	13.5	4.64	0.0899	2.08	15.6	1.17	1.89	417	110	590	118	27.9	0.0540	o	
27X-3, 98-100	240.36	37.3	0.640	13.4	4.54	0.0822	2.07	15.8	1.19	1.88	397	110	595	122	28.2	0.0543	o	Green
27X-3, 134-136	240.71	8.85	0.196	3.11	1.38	0.400	0.772	45.2	1.00	0.58	40	27	1330	38	80.7	0.0715	c	White
27X-4, 50-52	241.36	32.0	0.600	11.8	5.31	0.277	3.05	21.8	1.03	2.80	321	84	709	102	38.9	0.0579	o	Brown
27X-4, 110-112	241.94	50.3	1.07	23.0	7.19	0.0404	1.55	0.26	1.24	2.34	256	140	102	195	0.46	0.0526	o	Green
27X-5, 42-44	242.75	50.8	1.10	22.8	6.07	0.0188	1.54	0.27	1.17	1.99	330	144	111	196	0.47	0.0545	o	
27X-5, 72-74	243.04	51.8	1.10	22.6	6.33	0.0300	1.57	1.30	1.18	2.03	316	143	147	202	2.33	0.0553	o	
27X-5, 100-102	243.31	50.1	1.06	22.0</														

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)									Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr					
28X-1, 42-44	246.43	39.4	0.649	13.9	6.16	0.0766	2.22	11.1	1.28	2.25	383	121	432	120	19.7	0.0531	o	Dark green	
28X-1, 123-125	247.24	37.0	0.707	14.2	4.64	0.131	2.43	15.9	1.11	2.93	364	84	596	143	28.3	0.0566	o		
28X-2, 11-13	247.62	37.1	0.709	14.1	4.56	0.135	2.42	16.1	1.11	2.99	375	84	599	149	28.8	0.0570	o	Dark green	
28X-2, 76-78	248.27	40.1	0.637	13.6	5.24	0.111	2.20	13.7	1.23	2.21	397	120	552	120	24.5	0.0531	o	Dark green	
28X-3, 10-12	249.11	40.2	0.672	15.1	5.49	0.149	2.07	12.7	1.23	2.14	427	120	524	111	22.6	0.0506	o	Dark green	
28X-3, 45-47	249.46	40.3	0.668	14.2	5.61	0.104	2.21	12.2	1.20	2.10	387	132	488	123	21.8	0.0532	o	Dark green	
28X-CC, 7-9	249.64	41.4	0.662	13.9	5.49	0.112	2.18	12.4	1.24	2.18	349	133	489	132	22.0	0.0541	o	Green	
28X-CC, 36-38	249.93	40.3	0.690	14.7	5.69	0.157	2.20	12.1	1.22	2.29	356	132	492	119	21.5	0.0533	o	Dark green	
29X-1, 88-90	256.59	42.9	0.756	15.9	5.84	0.109	2.19	9.32	1.27	2.23	377	139	407	134	16.6	0.0540	o	Dark green	
29X-1, 134-136	257.05	48.3	0.984	22.2	6.77	0.0332	2.02	0.65	1.31	2.14	431	174	136	152	1.16	0.0503	o	Dark green	
29X-2, 22-24	257.43	41.2	0.795	16.2	6.36	0.309	2.46	11.5	1.15	3.32	388	93	505	142	20.5	0.0557	o	Brown	
29X-2, 54-56	257.75	5.03	0.124	1.76	0.963	0.448	0.488	49.0	0.84	0.33	88	19	1470	30	87.4	0.0801	c		
29X-2, 109-111	258.30	5.19	0.143	1.77	0.858	0.340	0.58	48.4	1.17	0.38	85	17	1190	30	86.4	0.0919	c	White	
29X-3, 34-36	259.05	42.7	0.773	16.7	5.45	0.137	2.18	9.93	1.18	2.33	409	146	444	134	17.7	0.0524	o	Dark green	
29X-3, 73-75	259.44	7.01	0.185	2.46	1.05	0.564	0.600	47.2	0.99	0.46	49	23	1400	35	84.2	0.0850	c	White	
29X-3, 146-148	260.17	51.7	1.17	22.0	5.89	0.0561	1.91	0.64	1.13	2.14	326	165	141	201	1.14	0.0600	o	Dark green	
29X-4, 36-38	260.56	43.7	0.774	16.7	5.57	0.163	2.40	9.12	1.23	2.33	415	150	437	132	16.3	0.0524	o	Dark green	
29X-4, 78-80	260.98	49.8	0.965	21.5	6.74	0.0713	2.10	1.17	1.28	2.37	348	172	160	155	2.09	0.0510	o	Dark green	
29X-5, 23-25	261.93	51.3	1.10	23.5	5.91	0.0192	1.90	0.38	1.27	2.41	285	198	154	186	0.68	0.0531	o	Dark green	
29X-5, 57-59	262.27	49.5	0.920	19.8	6.64	0.225	2.65	1.35	1.26	2.56	543	187	173	152	2.41	0.0526	o	Dark green	
29X-5, 116-118	262.86	50.3	1.06	21.9	6.09	0.0690	1.96	1.59	1.17	2.25	319	168	180	184	2.84	0.0547	o	Dark green	
29X-5, 140-142	263.10	51.0	1.12	17.3	6.81	0.206	2.99	1.12	1.60	3.14	246	93	169	212	2.00	0.0730	o	Dark brown	
29X-6, 27-29	263.48	5.53	0.164	1.83	0.866	0.266	0.521	49.2	0.96	0.39	125	24	1400	36	87.8	0.102	c	White	
29X-6, 78-80	263.99	5.36	0.148	1.75	0.942	0.285	0.500	49.0	0.93	0.32	163	21	1510	34	87.5	0.0956	c	White	
29X-6, 131-133	264.52	5.75	0.151	1.87	0.937	0.289	0.488	49.0	0.88	0.36	107	18	1570	33	87.5	0.0911	c	White	
29X-CC, 28-30	265.13	51.8	1.01	21.1	6.09	0.185	2.21	0.60	1.22	2.38	241	175	133	179	1.07	0.0541	o	Dark gray	
30X-1, 25-27	265.54	7.23	0.190	2.26	1.14	0.203	0.610	47.5	1.04	0.42	170	30	1490	40	84.8	0.0949	c	White	
30X-1, 99-101	266.26	44.6	0.749	15.3	5.56	0.0593	2.66	8.74	1.29	2.47	463	130	404	147	15.6	0.0553	o	Green	
30X-1, 147-149	266.74	50.0	0.913	20.4	6.08	0.0213	2.28	0.57	1.37	2.22	345	169	139	166	1.02	0.0508	o	Dark green	
30X-2, 44-46	267.21	39.7	0.661	13.5	4.94	0.0981	2.67	12.1	1.20	2.09	344	125	508	132	21.6	0.0553	o	Dark green	
30X-2, 115-117	267.92	47.8	0.830	17.0	6.04	0.0509	2.79	5.51	1.46	2.23	341	143	300	162	9.84	0.0553	o	Dark green	
30X-2, 144-146	268.21	7.30	0.163	2.30	1.00	0.252	0.587	48.4	0.94	0.49	120	30	1520	44	86.4	0.0806	c	White	
30X-3, 49-51	268.76	32.9	0.660	9.72	3.32	0.0736	2.06	22.8	1.17	1.77	340	62	844	145	40.7	0.0770	o	Dark green	
30X-3, 118-120	269.45	38.4	0.616	12.5	4.86	0.0622	2.73	14.9	1.27	1.95	443	109	620	125	26.6	0.0560	o	Dark green	
30X-4, 66-68	270.43	33.4	0.725	12.5	4.20	0.0861	2.25	19.3	1.11	2.49	466	79	792	130	34.4	0.0656	o	Dark green	
30X-4, 115-117	270.92	34.4	0.757	12.4	4.15	0.0902	2.26	19.3	1.17	2.35	428	83	782	137	34.4	0.0690	o	Pale green	
30X-5, 77-79	272.04	43.5	0.752	14.4	6.73	0.0499	2.82	7.90	1.35	1.88	437	128	345	153	14.1	0.0590	o	Dark green	
30X-5, 123-125	272.50	46.1	0.714	14.1	6.05	0.0506	2.96	9.13	1.47	2.10	324	127	405	142	16.3	0.0575	o		
30X-6, 20-22	272.97	45.0	0.704	14.2	5.48	0.0486	2.96	9.14	1.48	2.14	309	127	403	136	16.3	0.0562	o	Dark green	
30X-6, 73-75	273.50	40.0	0.599	12.0	5.00	0.0649	2.67	14.0	1.32	1.94	405	113	569	129	25.0	0.0564	o		
30X-CC, 22-24	274.17	41.0	0.586	11.6	5.05	0.0658	2.50	14.0	1.26	1.93	403	107	552	135	25.0	0.0571	o	Dark green	
31X-1, 58-60	275.49	43.2	0.700	14.4	5.10	0.0556	2.77	10.2	1.36	1.87	344	128	489	129	18.3	0.0552	o	Dark green	
31X-2, 41-43	276.82	52.8	0.953	18.1	5.79	0.0263	2.69	2.32	1.59	2.07	280	149	205	168	4.14	0.0595	o	Dark green	
31X-2, 127-129	277.68	36.8	0.513	10.6	4.38	0.0660	2.37	18.4	1.25	1.61	366	100	743	120	32.8	0.0550	o		
31X-3, 27-30	278.19	37.8	0.497	10.1	3.92	0.0635	2.26	17.4	1.26	1.64	363	96	698	129	31.0	0.0556	o	Gray/brown	
31X-3, 66-68	278.58	5.68	0.143	1.69	0.693	0.271	0.485	49.1	0.95	0.34	157	26	1590	36	87.7	0.0963	c		
31X-3, 102-104	278.94	7.77	0.206	2.32	0.860	0.256	0.597	46.8	1.00	0.46	141	30	1460	44	83.6	0.101	c	White	
31X-3, 138-140	279.30	47.1	0.721	14.1	6.50	0.126	3.14	8.04	1.83	1.68	236	128	425	141	14.4	0.0582	o	Green	
31X-4, 36-38	279.78	7.41	0.171	2.16	0.816	0.299	0.561	48.4	0.96	0.47	137	30	1570	42	86.4	0.0896	c	White	
31X-4, 108-110	280.50	35.9	0.529	11.0	4.07	0.0828	2.60	18.3	1.29	1.57	355	109	736	112	32.6	0.0547	o	Green/brown	
31X-4, 146-148	280.88	39.3	0.584	11.7	4.47	0.0806	2.67	16.0	1.36	1.87	331	110	680	125	28.5	0.0566	o		
31X-5, 30-32	281.22	39.1	0.578	11.7	4.40	0.0798	2.66	15.8	1.30	1.82	305	103	679	122	28.3	0.0561	o	Green	
31X-5, 63-65	281.55	6.32	0.136	1.90	0.715	0.379	0.807	48.8	0.94	0.34	116	32	1310	36	87.2	0.0812	c	White	
31X-CC, 19-21	282.40	35.8	0.480	10.2	4.13	0.0781	2.57	17.7	1.28	1.52	353	100	733	104	31.6	0.0531	o	Green	
32X-1, 16-18	284.97	38.4	0.502	10.7	4.03	0.0803	2.70	18.4	1.42	1.62	339	106	763	112	32.9	0.0531	o		
32X-1, 108-110	285.88	42.9	0.483	9.80	3.86	0.0658	2.57	15.1	1.42	1.51	435	101	621	116	27.0	0.0559	o		
32X-2, 58-60	286.86	43.4	0.468	9.43	4.19	0.0589	2.40	14.9	1.39	1.52	414	96	599	116	26.5	0.0562	o	Green/brown	
32X-2, 123-125	287.50	46.5	0.506	10.3	4.13	0.0667	2.54	12.5											

Table 2 (continued).

Core, section, interval (cm)	Depth (mbsf)	Major elements (wt%)								Trace elements (µg/g)					CaCO ₃ (wt%)	Ti/Al	Group	Color
		SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	Ba	Cr	Sr	Zr				
32X-CC, 16-18	294.26	50.3	0.675	13.6	5.72	0.0342	2.81	5.72	1.62	1.73	629	122	314	146	10.2	0.0565	o	Green/brown
33X-1, 76-78	295.08	53.4	0.786	16.5	5.73	0.0245	3.50	0.76	1.67	2.22	465	153	169	153	1.36	0.0541	o	Dark gray
33X-1, 133-135	295.63	53.4	0.796	16.2	5.91	0.0282	3.60	1.18	1.66	2.33	323	144	169	150	2.11	0.0556	o	Dark green
33X-2, 44-46	296.31	53.6	0.703	15.2	6.07	0.0469	4.05	1.56	1.81	2.36	412	162	165	132	2.79	0.0524	o	Dark green
33X-2, 98-100	296.84	34.7	0.702	10.9	4.09	0.0887	2.50	19.6	1.21	2.02	478	94	816	108	35.0	0.0731	v	
33X-2, 138-140	297.23	35.0	0.773	10.9	4.14	0.0923	2.44	20.2	1.29	2.03	502	93	819	117	36.0	0.0808	v	Gray
33X-3, 62-64	297.96	48.4	0.649	13.6	5.58	0.0528	3.56	7.40	1.54	2.19	409	138	341	125	13.2	0.0540	o	Green/brown
33X-3, 143-146	298.75	54.5	0.736	15.4	6.00	0.0522	3.73	2.06	1.66	2.29	340	151	178	139	3.68	0.0542	o	Dark gray
33X-4, 34-36	299.15	52.2	0.786	16.5	5.60	0.0196	3.38	0.84	1.73	2.06	592	166	181	138	1.50	0.0541	o	Dark gray
33X-4, 56-58	299.37	54.7	0.885	18.1	6.04	0.0277	3.43	1.15	1.63	2.18	359	163	196	150	2.05	0.0555	o	Dark gray
33X-4, 145-147	300.24	53.8	1.28	15.0	8.50	0.104	3.55	1.42	1.80	2.71	365	124	182	158	2.53	0.0970	v	Gray
33X-5, 55-57	300.83	55.3	1.18	16.2	7.21	0.0408	3.41	1.27	1.77	2.86	259	99	188	196	2.27	0.0827	v	Gray
33X-5, 127-129	301.54	54.1	0.845	17.9	5.86	0.0213	3.25	1.01	1.52	2.06	721	154	169	154	1.80	0.0536	o	Dark gray
33X-6, 22-24	301.98	52.7	0.979	19.7	5.62	0.0172	2.93	0.75	1.53	2.11	337	157	183	183	1.33	0.0563	o	Dark gray
33X-7, 27-29	303.51	50.8	1.21	16.4	6.64	0.0306	3.23	4.46	1.59	2.87	488	134	271	152	7.96	0.0837	v	Dark gray
33X-CC, 28-30	303.99	53.1	1.10	15.6	7.37	0.0478	3.55	1.21	1.73	2.35	213	146	169	159	2.16	0.0797	v	Dark gray
34X-1, 79-81	304.82	43.6	0.997	13.9	5.51	0.0418	2.85	9.39	1.54	2.02	453	139	423	140	16.8	0.0814	v	
34X-1, 104-106	305.07	43.3	0.783	13.3	5.15	0.0420	2.87	10.0	1.50	1.90	456	131	438	123	17.9	0.0666	v	Dark gray
34X-2, 16-18	305.67	39.6	0.571	11.9	4.93	0.0369	2.92	14.0	1.50	1.42	534	138	491	100	25.0	0.0543	o	Dark gray
34X-2, 33-35	305.84	33.3	0.564	9.62	3.94	0.0664	2.26	23.1	1.17	1.31	450	99	573	126	41.1	0.0664	c	
34X-2, 44-46	305.95	15.8	0.359	4.59	2.07	0.173	1.45	38.4	0.78	0.73	176	55	658	77	68.5	0.0886	c	
34X-2, 60-62	306.11	6.30	0.213	1.87	0.806	0.244	1.08	48.2	0.68	0.37	70	28	621	43	86.0	0.130	c	
34X-2, 95-97	306.46	4.21	0.147	1.21	0.531	0.219	0.987	49.9	0.65	0.25	43	33	572	43	89.1	0.137	c	
34X-CC, 27-29	306.78	2.49	0.0983	0.766	0.320	0.128	0.999	51.4	0.41	0.12	31	30	565	32	91.8	0.146	c	
35X-1, 71-73	314.32	6.83	0.174	1.95	0.712	0.123	1.15	47.6	0.79	0.34	44	34	405	88	85.0	0.101	c	White to brown
35X-CC, 36-37	314.80	58.8	1.08	5.47	2.33	0.538	1.20	13.5	1.12	1.08	292	123	260	421	24.1	0.223	sst	White
36X-1, 91-93	324.05	0.85	0.0319	0.287	0.124	0.179	0.863	53.0	0.80	0.06	4	29	439	26	94.5	0.126	c	White

Notes: Total iron expressed as Fe₂O₃; CaCO₃ calculated from CaO data, assuming that all Ca is present as carbonate. Lithologic groups are: c = calcareous; v = volcanic; o = organic-rich and other turbidites; sst = volcanoclastic sandstone. Italicized letters in the Color column are Quaternary turbidite identifiers of Weaver and Kuijpers (1983), De Lange et al. (1987), Weaver, Buckley et al., 1989, and Weaver, Thomson, et al., 1989; bed divisions are indicated by horizontal lines.

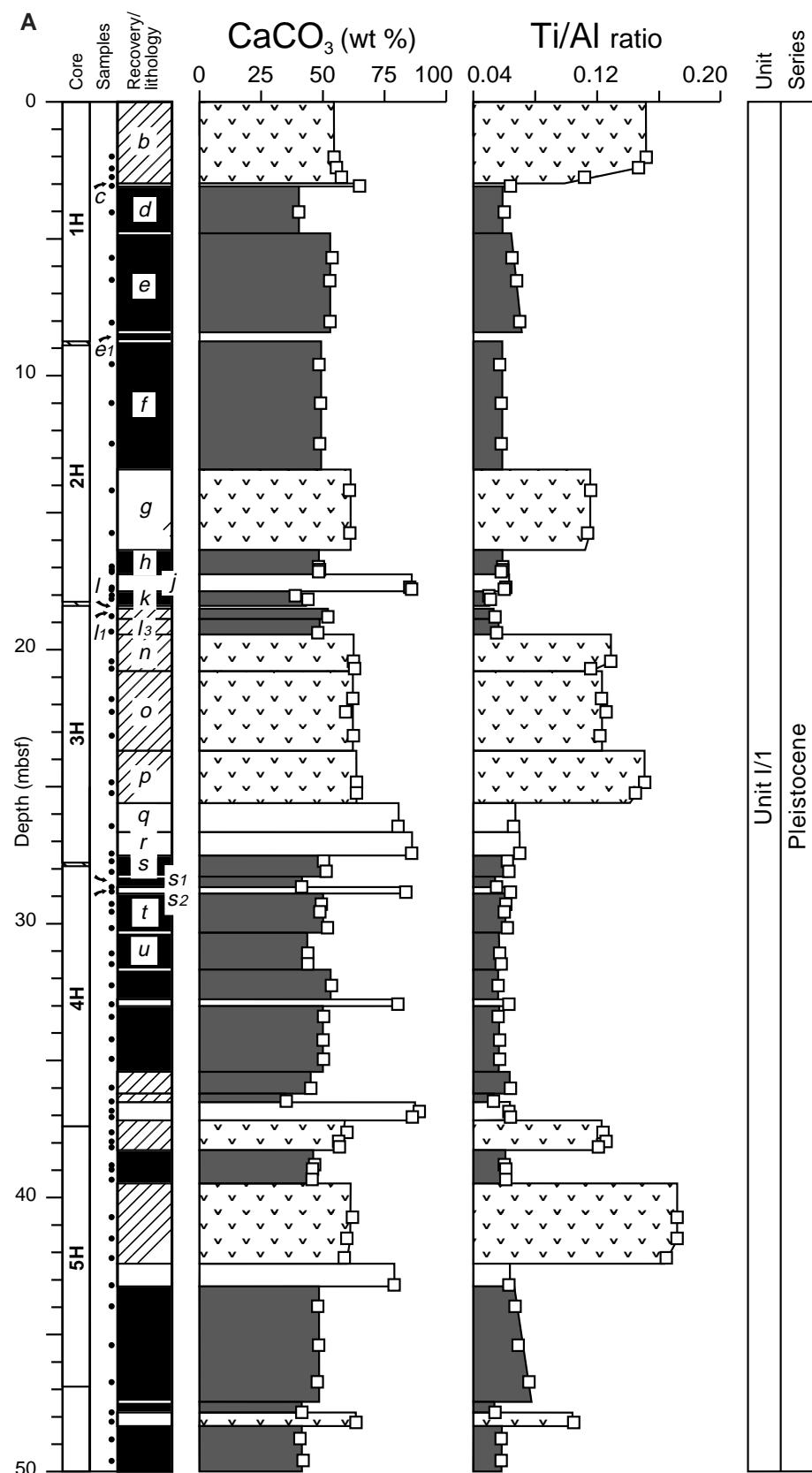


Figure 2. A–G. Carbonate and Ti/Al ratio profiles plotted against lithology for Site 950. Only turbidite data are plotted; pelagic intervals are excluded. Turbidite classification: (a) calcareous turbidites, >75% CaCO₃ (calculated from CaO data); (b) volcanic turbidites, Ti/Al > 0.8 and CaCO₃ < 75%; (c) organic-rich and other turbidites. Turbidite lettering in Sections 157–950A–1H through 4H after Weaver and Kuijpers (1983) and Weaver et al. (1989a, 1989b); series boundaries derived from Howe and Sblendorio-Levy (Chap. 29, this volume).

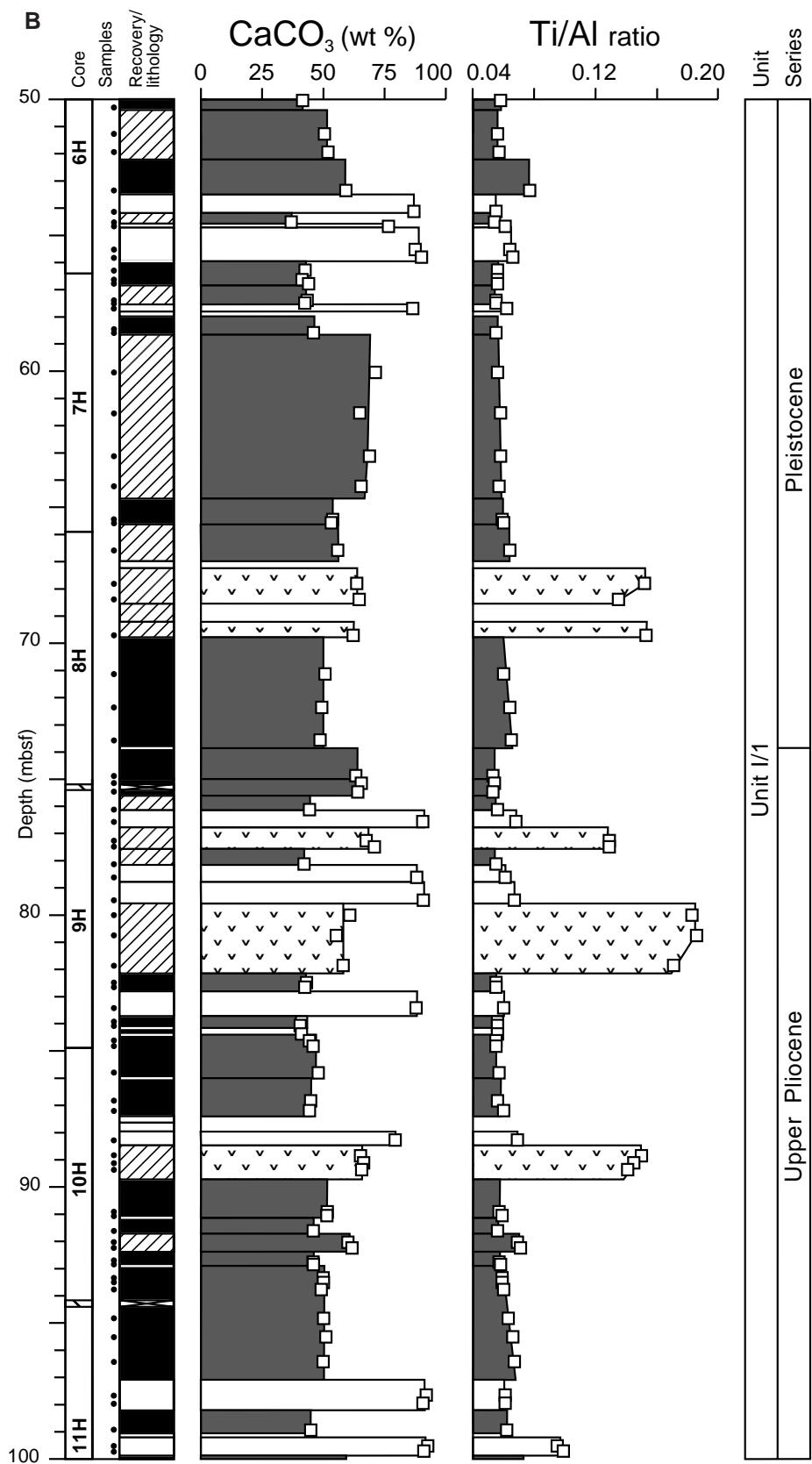


Figure 2 (continued).

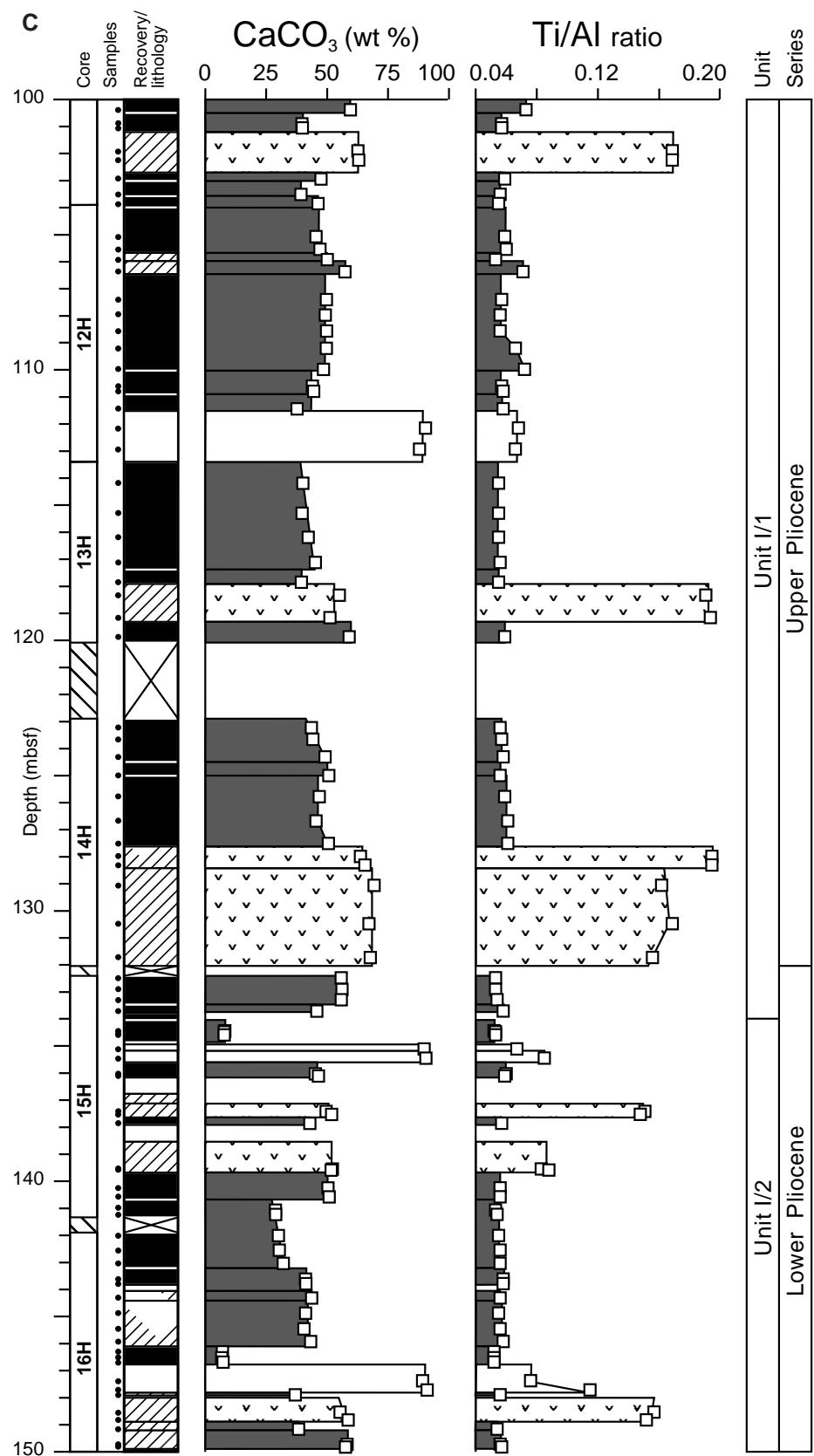


Figure 2 (continued).

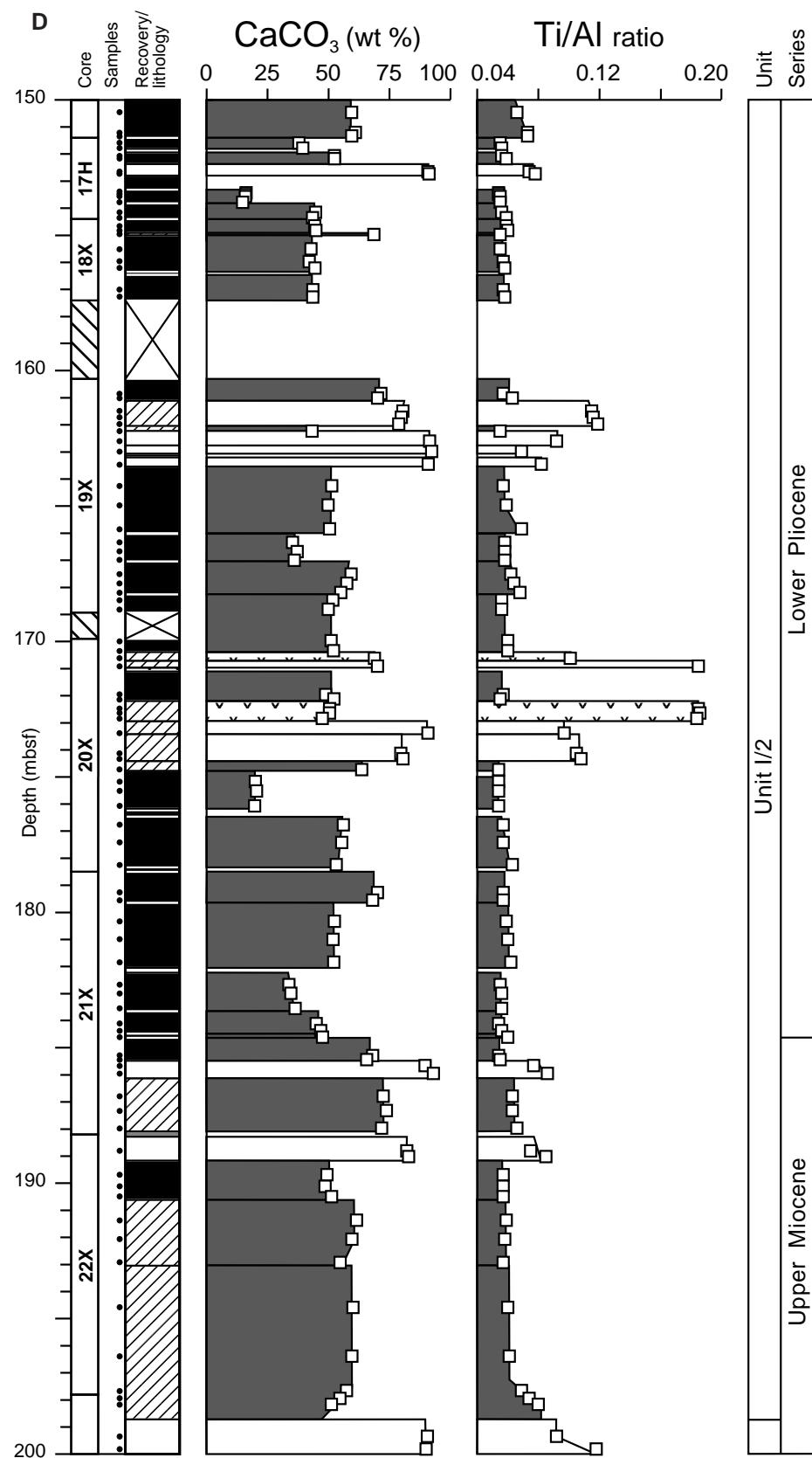


Figure 2 (continued).

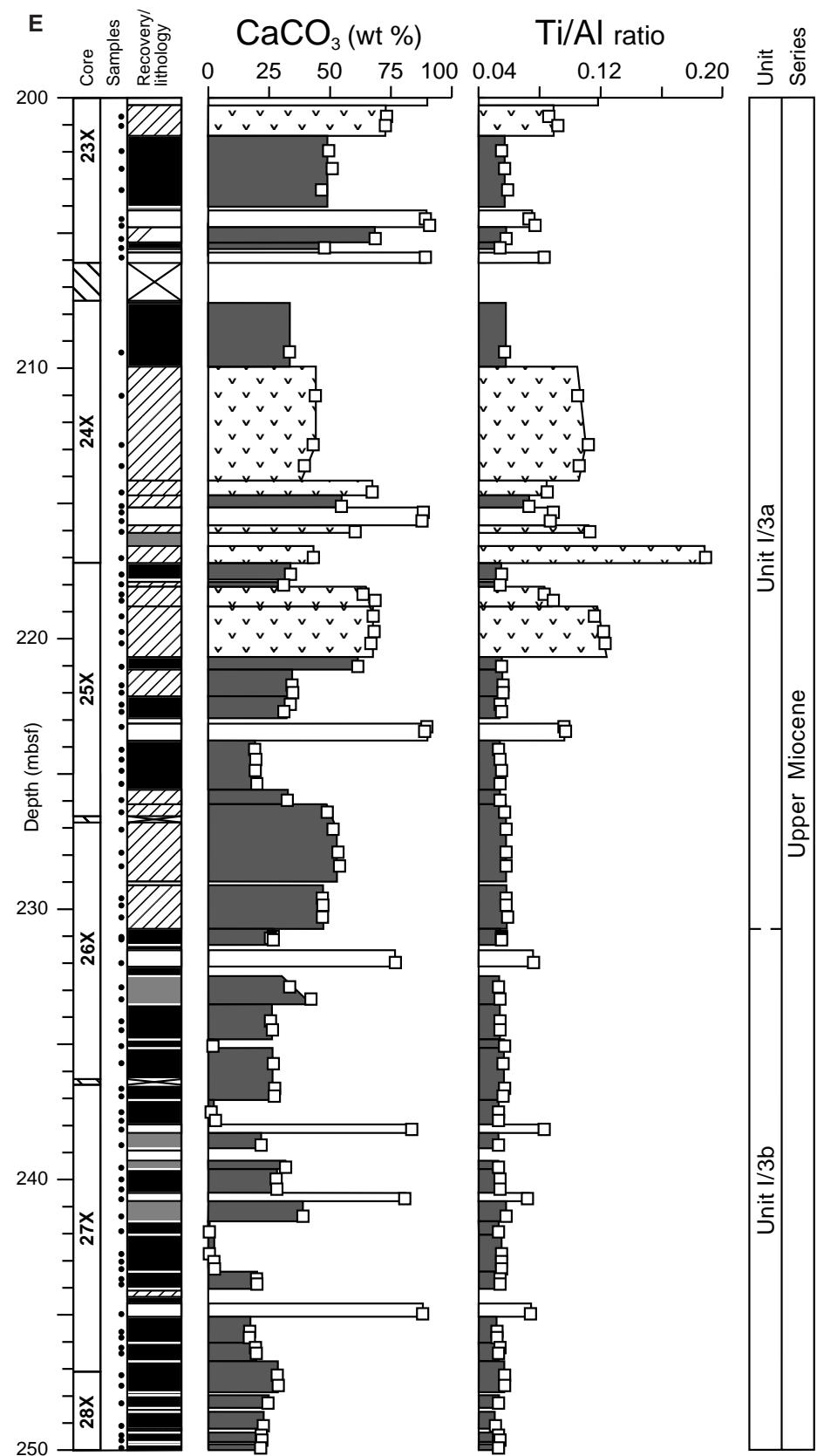


Figure 2 (continued).

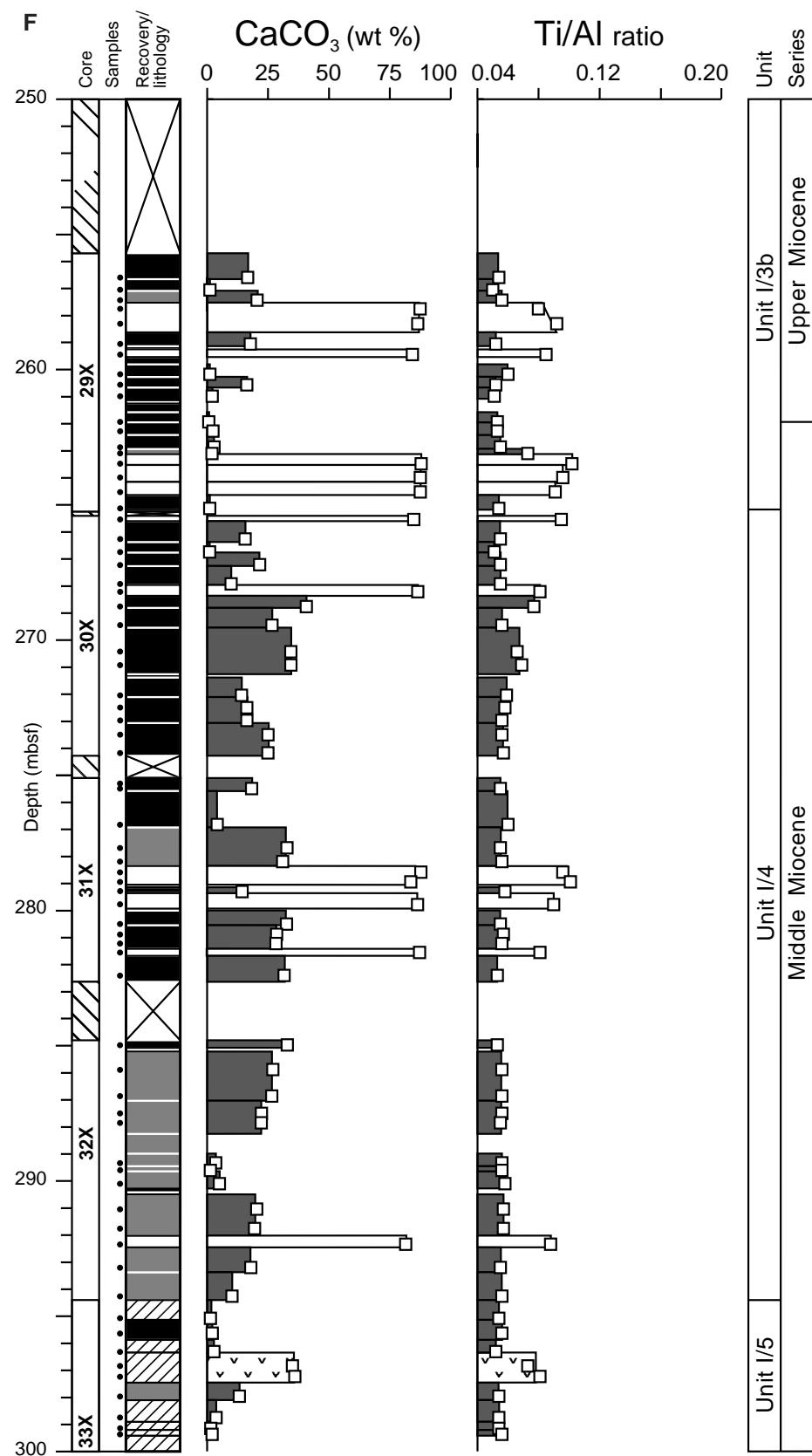


Figure 2 (continued).

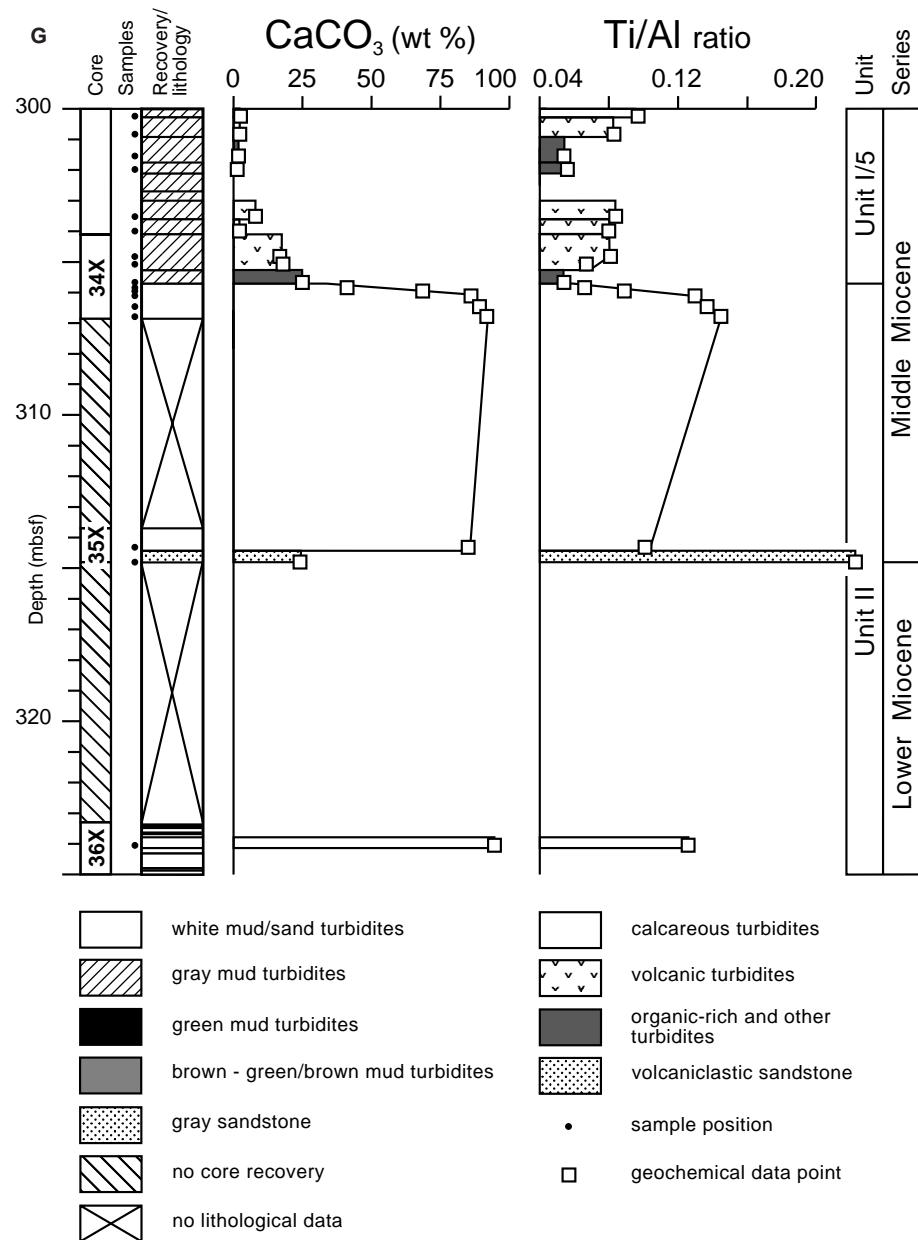


Figure 2 (continued).

Figure 2 demonstrates that shipboard identification of all gray turbidites as being members of the ‘volcanic’ group was erroneous, and that the frequency of volcanic turbidite deposition is much lower than originally thought. The detailed log (Fig. 2), combined with summary geochemical profiles (Fig. 3) for CaCO₃ and Ti/Al ratios through shipboard lithostratigraphic Unit I (Schmincke, Weaver, Firth, et al., 1995), enable five new subdivisions to be designated here (defined here as chemostratigraphic Units I/1 through I/5):

Unit I/1, 0–134 mbsf (upper Pliocene–Pleistocene)

The pattern of turbidite deposition seen in the upper Pleistocene–Holocene, continues down to ~134 mbsf (low–upper Pliocene). Turbidites in this interval fall clearly into one of the three previously defined turbidite groups. Ti/Al ratios for the volcanic turbidites form a broad array with increasing values downward, attaining maximum

values of ~0.2 toward the base of the unit. There is a suggestion that the high-Ti (turbidites *b*, *p*; Fig. 2) and low-Ti (*g*, *o*) subgroups defined by Pearce and Jarvis (1995) in the Quaternary sediments, continue downcore with a trend toward increasing Ti/Al ratios in both groups.

Carbonate contents of both the volcanic and calcareous turbidites also generally increase downward through this interval. Ti/Al ratios of the organic-rich turbidites are remarkably consistent throughout the section, with small excursions toward higher values being caused by samples from the lower portions of beds, which commonly contain a proportion of entrained volcaniclastic grains (Pearce and Jarvis, 1992a, 1992b, 1995). A thick gray turbidite at 59–65 mbsf, originally assigned to the volcanic group (Schmincke, Weaver, Firth, et al., 1995), contains a high carbonate content of 68%, but a low Ti/Al ratio of 0.057 (Figs. 2B, 3). A sample (157-950A-7H-3, 78–79 cm; 61.68 mbsf) from this turbidite analyzed for C_{org} aboard ship,

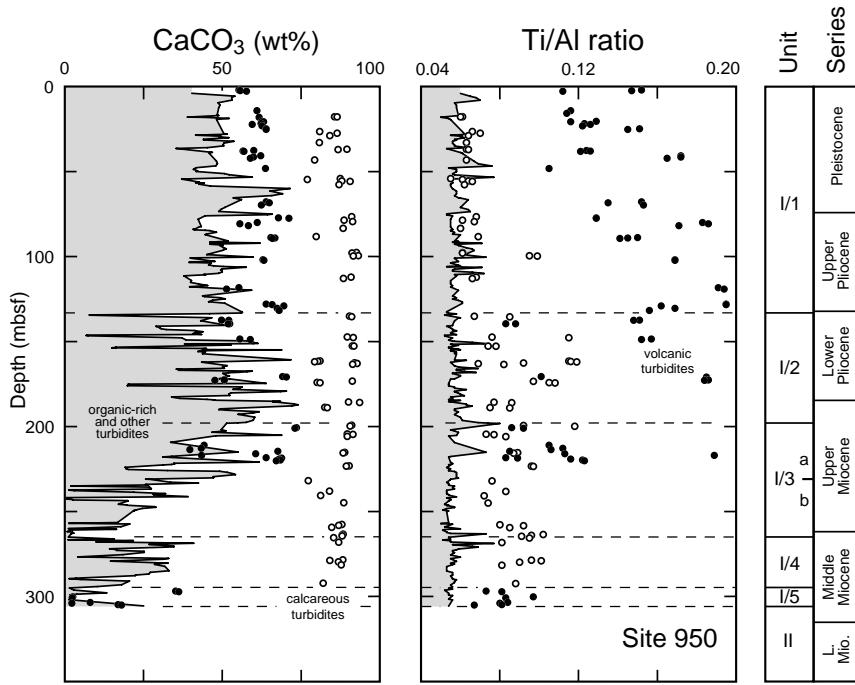


Figure 3. Geochemical variation in turbidite carbonate contents and Ti/Al ratios, Unit I, Site 950. Note the clear distinction between volcanic (solid circles; 71 data points), calcareous (open circles; 70 points), organic-rich and other (shaded area and continuous line; 338 analyses) turbidites in Unit I/1.

yielded 0.51% (Schmincke, Weaver, Firth, et al., 1995), indicating that despite its color, the turbidite belongs to the organic-rich group.

The clear distinction between volcanic and other turbidites is shown by data for other elements known to be enriched in the volcanoclastic fraction (De Lange et al., 1987, 1989), particularly Fe/Al, Cr/Al, and Zr/Al (Fig. 4). Fe/Al ratios, are also lowest in the organic-rich turbidites, occur at intermediate levels in the calcareous turbidites, and are highest in the volcanic category. Chromium contents overlap between the organic-rich and volcanic groups, but trend to the highest values in the latter category. Zirconium contents are invariably highest in volcanic turbidites. Unlike Ti, only Cr/Al displays a clear downhole increase within the volcanic group of Unit I/1, and in this case a maximum is reached in a turbidite at 102 mbsf, some distance above the base of the unit.

Pearce and Jarvis (1995) recognized two subgroups within Quaternary organic-rich turbidites, based largely on potassium contents. Such a distinction is apparent in the K/Al ratios of the green turbidites in Unit I/1 at Site 950 (Fig. 5), with high-K examples being characterized by ratios of >0.3, and a second subgroup having ratios of 0.2–0.3; however, separation between the fields of the two subgroups becomes less clear in the mid-Pleistocene. A similar pattern is shown by Mg/Al (Fig. 6), with organic-rich turbidites at 18 (turbidite *k*), 56, 84, 85, 115, and 133 mbsf, in particular, displaying markedly K- and Mg-depleted compositions. Si/Al ratios vary considerably in the organic-rich turbidites of Unit I/1 (Fig. 7), but show no clear stratigraphic trends. Volcanic and calcareous turbidites display much less variation, with overlapping arrays on K, Mg, and Si plots, and little distinction between beds.

Unit I/2, 134–199 mbsf (uppermost Miocene–lower Pliocene)

A transition toward increased proportions of thin-bedded turbidites ~134 mbsf (Fig. 2C), is accompanied by the appearance of low-carbonate (<10% CaCO₃) dark-green organic-rich turbidites, gray volcanic turbidites with low Ti/Al ratios (~0.09), and calcareous turbidites with high Ti/Al ratios of >0.07 (Fig. 3). Samples from calcareous turbidites with the highest Ti/Al ratios (up to 0.12) are generally from the lower portions of beds. Together, these changes produce overlap between the calcareous and volcanic turbidite fields on the

Ti/Al ratio profile (Fig. 3), a characteristic that continues down through to the remainder of Unit I. Indeed, two gray calcareous turbidites within this interval (at 161 and 174 mbsf), have carbonate contents of 80%, but high Ti/Al ratios of ~0.11.

Despite the occurrence of a few dark green turbidites with very low carbonate contents, there is a general downward increase in the carbonate contents of the nonvolcanic turbidites through Unit I/2; CaCO₃ contents attain a maximum in a series of three thick, gray turbidites with low Ti/Al ratios (Fig. 2D) at 199 mbsf. The bottom of Unit I/2 is defined by this shift in the carbonate profile (Fig. 3); calcareous turbidites also attain their highest CaCO₃ contents of >90% toward the base of the unit. Shipboard C_{org} determination (Schmincke, Weaver, Firth, et al., 1995) of a sample (157-950A-23X-1, 11–12 cm; 197.91 mbsf) taken from one of the thick, gray nonvolcanic beds at the base of Unit I/2, gave a low value of 0.19%. Given the considerable thickness (>5 m) of this turbidite, it is unlikely that it represents an oxidized member of the organic-rich group. This points to a different provenance for this, and possibly the other thick, gray nonvolcanic beds; additional organic carbon and other geochemical data are required to test this hypothesis.

A number of high-Ti volcanic turbidites continue to occur in Unit I/2, and also have high Fe and Zr, but low Cr contents (Fig. 4). Occasional high Fe values in samples from organic-rich turbidites are attributed to the occurrence of scattered pyrite in the deeper turbidites. The two subgroups of organic-rich turbidites identified in Unit I/1 continue through Unit I/2, and display even clearer separation on K and Mg profiles (Figs. 5, 6). The low-carbonate organic-rich turbidites that first appear in Unit I/2 are also distinguished by having the lowest K/Al (<0.2), Mg/Al, and Si/Al ratios in the sequence (Figs. 5 through 7), forming a distinctive third compositional subgroup on the geochemical profiles.

Unit I/3, 199–265 mbsf (uppermost middle Miocene–upper Miocene)

Turbidites below 199 mbsf generally display declining CaCO₃ contents, reaching a minimum in a sequence of thin-bedded dark-green turbidites at the base (Fig. 2E–G), which contains no significant carbonate. Even the calcareous turbidites become increasingly

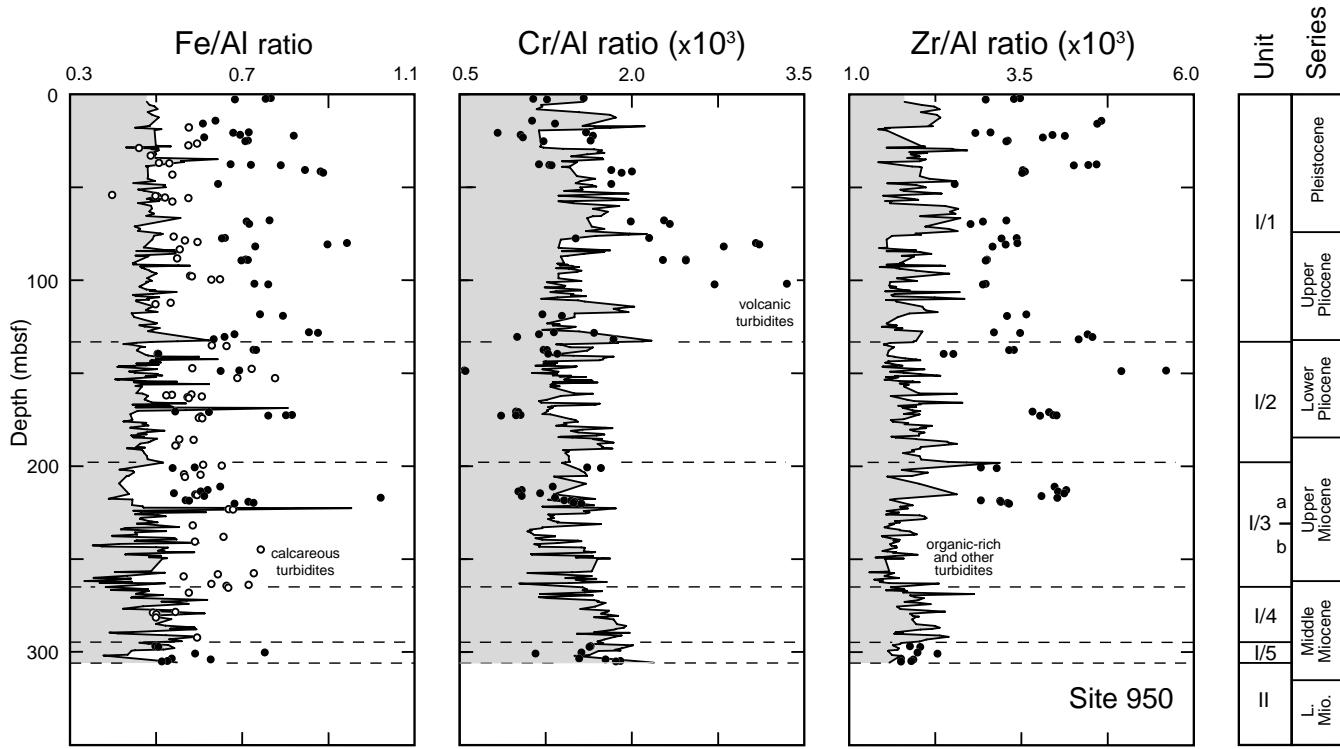


Figure 4. Geochemical variation in turbidite Fe/Al, Cr/Al, and Zr/Al ratios, Unit I, Site 950. Volcanic (solid circles), calcareous (open circles), organic-rich and other (shaded area and continuous line) turbidites follow discrete trends similar to those displayed by Ti/Al ratio plots (Fig. 3). Calcareous turbidites have been omitted from the Cr/Al and Zr/Al plots because scattering produced by poor analytical reproducibility masks any geochemical trends.

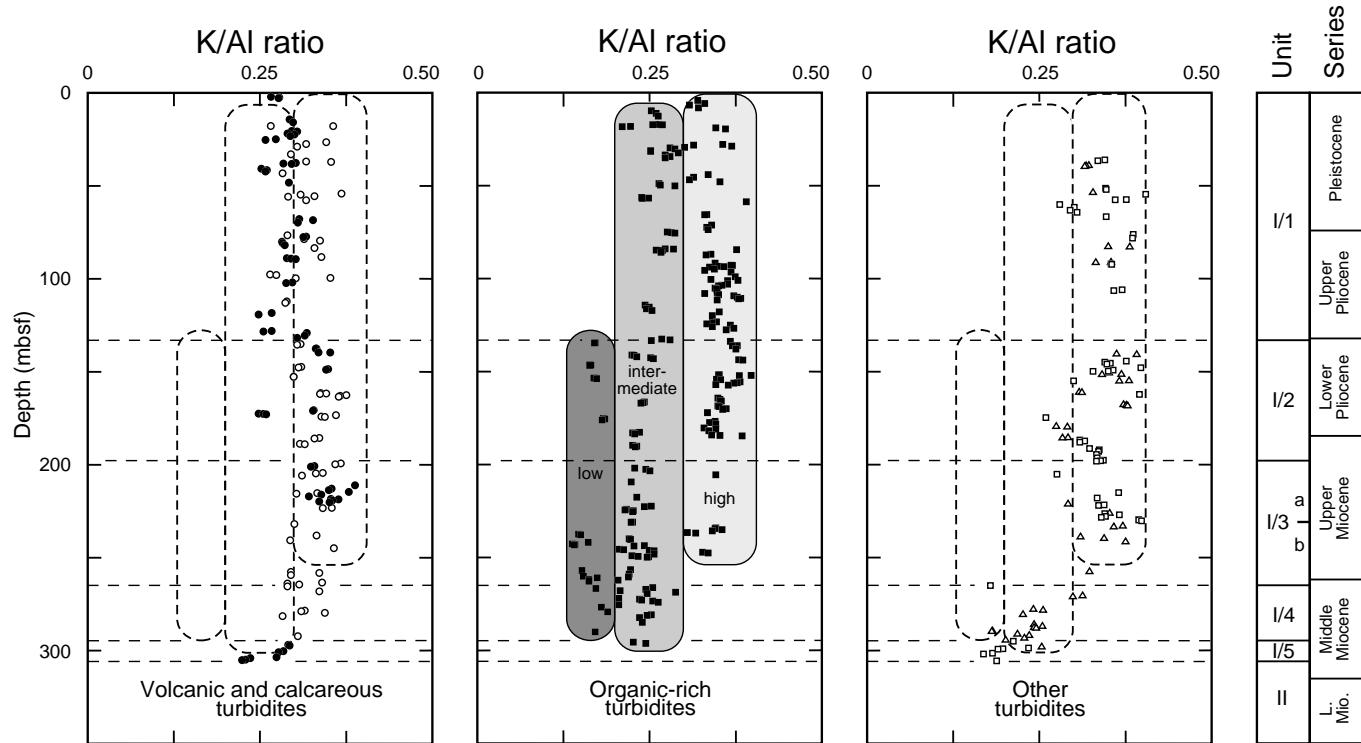


Figure 5. Stratigraphic variation in turbidite K/Al ratios, Unit I, Site 950. Volcanic (solid circles) and calcareous (open circles) turbidites have similar ratios, typically >0.25 . Green organic-rich turbidites (filled squares; 229 analyses) display three distinct compositional arrays (shaded): with high (>0.3), intermediate (0.2–0.3) and low (<0.2) K/Al ratios. Gray nonvolcanic (open squares; 60 data points) and other (open triangles; 48 analyses) turbidites fall mostly within the field (outlined) of the high K/Al organic-rich group, except near the base on Unit I.

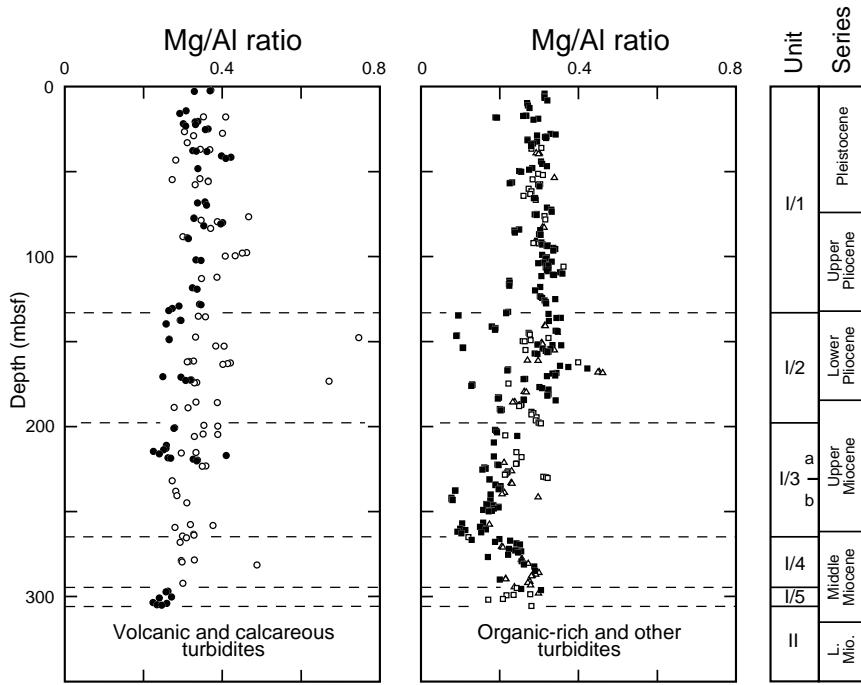


Figure 6. Stratigraphic variation in turbidite Mg/Al ratios, Unit I, Site 950. Symbols as in Figure 5.

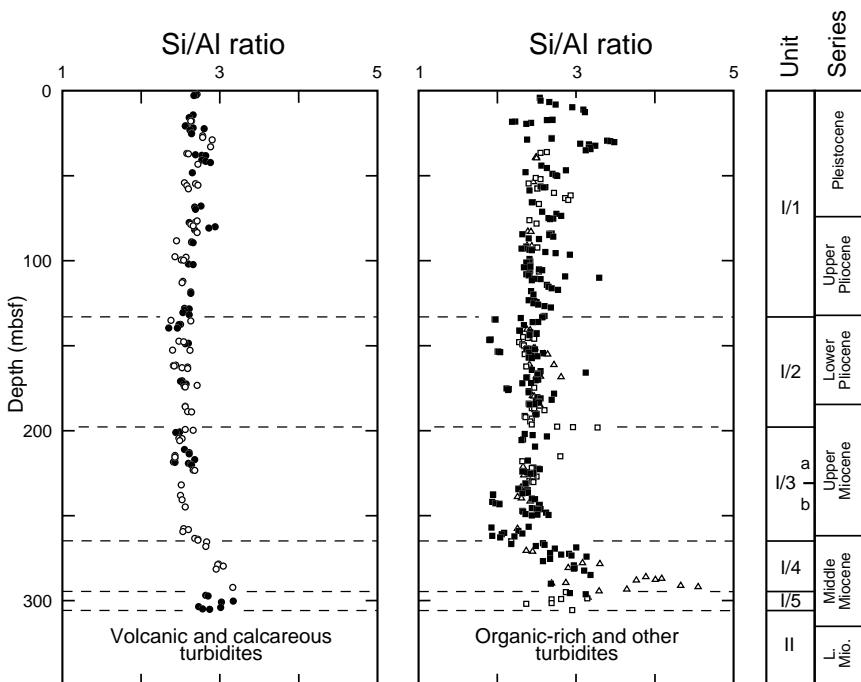


Figure 7. Stratigraphic variation in turbidite Si/Al ratios, Unit I, Site 950. Symbols as in Figure 5.

impure through this interval (Fig. 3). The last thick, gray volcanic turbidites occur ~221 mbsf, and these, with one exception (at 217 mbsf), are characterized by low Ti/Al ratios of ≤ 0.12 . There is a noticeable overall decrease in bed thickness below 231 mbsf (Fig. 2E), with older turbidites typically < 2 m thick. This change in bedding style and lithology is used to subdivide the unit into I/3a (above) and I/3b (below).

Organic-rich turbidites in Unit I/3 display an increasing dominance of the intermediate K subgroup (K/Al ratios, 0.2–0.3) toward the base, with common low K (K/Al ratios, < 0.2) examples in Sub-unit I/3b (Fig. 5). Separation between the subgroups is also seen on the Mg/Al (Fig. 6), and to a lesser extent the Si/Al (Fig. 7) profiles.

However, there is a general shift toward lower Mg contents through Unit I/3, a consequence of declining carbonate, which contains significant Mg. The Mg/Al minimum, therefore, corresponds to the carbonate minimum at the base of Unit I/3; K and Si profiles remain unaffected.

Unit I/4, 265–294 mbsf (mid- to high middle Miocene)

Unit I/4 is dominated by thinly bedded green and green/brown turbidites (Fig. 2; Table 2), with low and variable CaCO_3 contents of 0% to 34%. Calcareous turbidites have moderate carbonate contents (87%) and high Ti/Al ratios of ~ 0.09 . The oldest dark-green high K

organic-rich turbidite occurs toward the top of Unit I/4 at 247 mbsf (Fig. 5), although brown turbidites with similar compositions occur below, at 257 and 263 mbsf.

All turbidite groups display increasing Si/Al ratios downward through Unit I/4, with the highest Si/Al ratios occurring in a unique suite of green/brown turbidites between 286 and 294 mbsf (Figs. 2, 7; Table 2), at the very base of the unit.

Unit I/5, 294–306 mbsf (low middle Miocene)

The oldest beds in Unit I consist of a thin-bedded sequence of dark-gray turbidites (Unit I/5; Fig. 2), with variable Ti/Al ratios and low carbonate contents. Many have Ti/Al ratios of >0.08, and so are assigned to the volcanic group, but they are lithologically and geochemically distinct from younger members of the group, which only appear above 221 mbsf, in the upper Miocene. Despite their Ti-enriched signatures (Fig. 3), these middle Miocene volcanic turbidites have lower Zr contents (Fig. 4), than other volcanic turbidites in the core; they are further distinguished by having low Mg and K, and high Si to Al ratios. Organic-rich and other turbidites in this interval are all medium to low K/Al varieties (Fig. 5), with relatively high Si contents (Fig. 7) and no clear separation between subgroups. No calcareous turbidites occur in Unit I/5.

Unit II, 306–333 mbsf (lower Miocene–lowest middle Miocene)

Unit II is characterized by very poor core recovery; only a few sections of carbonate sand and gravel were obtained. Downhole logs indicate that three coarse-grained carbonate units occur in Unit II (Schmincke, Weaver, Firth, et al., 1995), and that its base coincides with the base of the nonrecovered interval (top of Core 157-950A-37X). A unique feature in this sequence is the occurrence at its top of a hybrid turbidite, which grades upward from carbonate sands and gravels containing 90% CaCO₃, at its base (Fig. 2), to brown muds with 40% CaCO₃ at its summit. High Ti/Al ratios of 0.14 in the coarser lower part of the turbidite reflect a high proportion of volcanioclastic material, dominantly zeolitic vitric tuff and basaltic rock fragments with occasional glass shards and mineral crystals; this fraction is diluted by increasing proportions of clay minerals upward, and at the top of the bed the Ti/Al ratio is only 0.07.

A sample of a thinner calcareous turbidite sand at 324 mbsf (Section 157-950A-36X-1, 91–93 cm; Fig. 2G; Table 2) also has a very high Ti/Al ratio of 0.13, suggesting a similar source to the thick hybrid turbidite. A gray volcanioclastic sandstone at 315 mbsf (Section 157-950A-35X-CC, 36–37 cm; Fig. 2G; Table 2), displays the highest Ti/Al ratios measured, at 0.22. Both of these sediments have moderate Fe but very high Cr and Zr contents. Very high Mg/Al ratios characterize the carbonate sediments within this interval, attaining a value of 3.4 in the turbidite sand at 324 mbsf. The volcanioclastic sandstone, on the other hand, displays low Mg but high Si to Al ratios.

Turbidite Provenance

Organic-Rich Turbidites

The deposition of organic-rich turbidites at Site 950 began in the early middle Miocene (~15 Ma), with the first dark-green turbidite, now at 296 mbsf (two thin, green turbidites occur below this, ~324 mbsf, but these contain negligible C_{org}). Organic-rich turbidites have dominated deposition on the MAP since that time.

Organic-rich turbidites display major changes in carbonate contents and other geochemical parameters through the sequence. Deposition on the plain during the middle Miocene was predominantly thin-bedded, low-carbonate sediments with low to medium K/Al and high Si/Al ratios. Major changes occurred during the late Miocene

(6–11 Ma), with a progressive increase in bed thicknesses and carbonate contents, and a shift toward the deposition of turbidites containing higher K/Al and Mg/Al ratios. Carbonate contents generally fell slightly through the latest Miocene to latest early Pliocene (3.5–6 Ma), a trend accompanied by decreasing bed thicknesses and an increasing dominance of high-K over medium-K sediments, but with the continued intermittent influx of flows with very low carbonate, K-depleted compositions. The low-K source switched off in the latest early Pliocene, and deposition during the last 3.5 m.y. has been remarkably uniform, with deposition being dominated by thick-bedded, high-carbonate, high-K organic-rich turbidites, accompanied by the occasional input of medium-K beds.

Preliminary interpretation of these trends is that, since the latest Miocene, organic-rich turbidites have originated predominantly from a chlorite- and illite-rich sediment source area on the northwest African continental slope off Morocco (Fig. 1), but with regular input of more kaolinitic sediment from a southerly source area, probably off Western Sahara. A third source, possibly a high-productivity area located even further to the south, was particularly active during the early late Miocene and mid-Pliocene. The less potassic sediments that dominate the middle and upper Miocene may be a consequence of the increased importance of the southerly source area, and/or climate change, promoting mineralogical changes in the northern source. Diagenetic processes are not considered to be a likely cause of the observed geochemical trends, because pore-water profiles (Schmincke, Weaver, Firth, et al., 1995) demonstrate minor uptake rather than the loss of potassium from deeper sediments at Site 950.

Volcanic Turbidites

Pearce and Jarvis (1995) concluded that during the late Quaternary, high-Ti volcanic turbidites on the MAP were derived from the younger western Canary Islands (Fig. 1), or possibly the basaltic complexes of Tenerife, whereas the low-Ti subgroup reflected a more fractionated volcanic source, and probably originated from the northern flanks of the central and eastern Canary Islands.

The oldest members of the volcanic turbidite group at Site 950 are confined to the early middle Miocene, ~14–16 Ma. These thin-bedded, dark-gray turbidites are lithologically quite distinct from younger members of the group and have trace-element-depleted signatures. Their origin is currently uncertain.

The first thick-bedded, gray volcanic turbidites were deposited in the mid-late Miocene, ~6.5 Ma. This corresponds to volcanic hiatuses on Gran Canaria and Fuerteventura, and the early submarine stage of Tenerife (Schmincke, 1976, 1982, 1994). It has been demonstrated that Gran Canaria was supplying very little sediment to the deep ocean at that time (Schmincke, Weaver, Firth, et al., 1995). Low Ti, Fe, and Cr, and moderate Zr concentrations predominate in the late Miocene turbidites, pointing toward an evolved volcanic source area, probably the slopes of Lanzarote or Gomera. However, an increasing proportion of Ti-rich turbidites were deposited during the early Pliocene,

A major change in volcanic turbidite geochemistry occurred at the beginning of the late Pliocene, ~3.5 Ma, with the disappearance of low-Ti sediments and an influx of common turbidites with very high Ti/Al and Fe/Al ratios. This points to a major shift toward a more basaltic source area. The island of La Palma was initiated around this time (Schmincke, 1994), while activity had ceased on Gomera and a volcanic hiatus was occurring on Tenerife. Gran Canaria was undergoing the development of a large stratocone, with the eruption of alkali basalt, trachyte, basanite, and phonolitic lavas and pyroclastics. It is speculated that the development of the La Palma Shield may have caused sediment instability on the slopes of the western Canary Islands, providing a significant sediment source at that time.

There is a decline in Ti/Al and Fe/Al and a general increase in Zr/Al ratios through the subsequent 3.5 m.y., indicating increasing input of sediment from areas having more evolved volcanic compositions. La Palma, Tenerife, and, more recently, Hierro, seem likely sources for the more basic material, with Gran Canaria, Lanzarote, and Fuerteventura providing a more fractionated volcaniclastic component.

Calcareous Turbidites

The deposition of thin calcareous turbidites on the MAP has occurred since at least the late Eocene. They occur regularly through the middle Miocene to Pleistocene record. Prior to the late Pliocene (~3.5 Ma), geochemical evidence indicates that calcareous turbidites entrained significant amounts of basaltic volcaniclastic debris, implying the active erosion of the westerly seamount chains (Fig. 1), which are believed to have acted as their source area. Younger turbidites contain little volcaniclastic material, indicating that the seamounts have been covered by a pelagic drape since that time.

Other Turbidites

A number of turbidites that do not fall clearly into any of the above groups have been observed at Site 950. Many of these are thin, gray turbidites that probably represent the now completely oxidized members of the organic-rich group. Such turbidites fall entirely within the geochemical compositional arrays defined by green turbidites. A smaller number of beds fall between the main arrays and may have mixed compositions. The source area of several thick, gray turbidites with nonvolcanic signatures remain uncertain, and the distinctive package of middle Miocene green/brown Si-rich turbidites also merits further investigation.

Thick, carbonate-debris flows and calcareous turbidites deposited during the early Miocene have distinctive trace-element-enriched geochemical signatures, reflecting the inclusion of a high proportion of basaltic and other volcaniclastic debris. In the absence of dolomite, high Mg/Al ratios point to the inclusion of high-Mg calcite in the carbonate-sand fraction. The latter is consistent with a shallow-water origin for these sediments. They probably originated from the upper flanks of the Cruiser/Hyères/Great Meteor Seamount chain, which may have been emergent at that time.

CONCLUSIONS

Significant turbidite deposition on the MAP at Site 950 began in the early Miocene with the deposition of thin, carbonate-sand turbidites and three thick carbonate-debris flows. These beds have distinctive geochemical signatures, reflecting the incorporation of significant amounts of basaltic material and shallow-water carbonate grains derived from the erosion of seamounts lying to the west of the plain. The middle Miocene to Pleistocene was dominated by the deposition of distal mud turbidites, beginning at ~15 Ma. Three major turbidite compositional groups (organic-rich, volcanic, and calcareous), originally defined in the Quaternary of the MAP, have been recognized in the older sedimentary record on the plain. Geochemical data define five chemostratigraphic units within the turbidite succession, reflecting sediment evolution in turbidite source areas and changes in provenance during the history of deposition on the plain.

Organic-rich turbidites dominate the sedimentary record, and become progressively more K and Mg rich with time. Three subgroups are evident from the geochemical data, indicating significant changes in sediment sources, particularly during the early late Miocene (10–11 Ma), with a shift toward more potassic- (illitic and chloritic) and carbonate-rich compositions, and during the latest early Pliocene (~3.5 Ma) with the final disappearance of very low-K (kaolinite-

rich), carbonate-poor turbidites. The high-K subgroup probably originated principally from a northern source area on the upper continental slope off Morocco, whereas high-Al sediments were derived from the south, off Western Sahara. Climatic changes are also likely to have modified sediment mineralogy in the competing sources areas.

Volcanic turbidites are volumetrically the second most important sediment type through most of the sequence, although carbonate turbidites are more frequent. A thin package of thinly bedded dark gray volcanic turbidites with distinctive trace-element-depleted geochemical signatures was deposited during the early–middle Miocene (14–16 Ma). These are of uncertain affinity. The first typical thick-bedded, Ti-, Fe-, and Zr-rich volcaniclastic turbidites were deposited on the MAP in the mid-late Miocene, ~6.5 Ma, and probably originated from the Canary Island slopes of Lanzarote or Gomera. A major change to turbidites with very high Ti and Fe contents occurred around the beginning of the late Pliocene (~3.5 Ma), possibly associated with sediment failure during the early growth of La Palma. Younger volcanic turbidites display a clear trend toward progressively more fractionated volcanic sources since 3.5 Ma, although their wide range of trace-element compositions indicate continued supply from a variety of different source areas on the Canary Island slopes.

Thin-bedded calcareous turbidites occur throughout the sequence. They originated predominantly from the seamount chains to the west of the MAP and, up until the latest early Pliocene, incorporated a significant proportion of basaltic material derived from erosion of the exposed volcanic edifices. Since ~3.5 Ma, these seamounts have been largely covered by a pelagic sediment drape, which now provides the main sediment source for calcareous turbidites on the plain.

ACKNOWLEDGMENTS

We thank the crew, marine technicians, and fellow scientists of Leg 157 for their help and enthusiasm. In particular, Sten Lindblom, Anne Pimmel, and Robert Kemp managed to survive the shipboard geochemical program. Erinn McCarty remained jovial despite our requests for difficult samples. Meryl Batchelder provided valuable technical assistance in the initial stages of this project. Constructive reviews were provided by Gert De Lange, Tim Pearce, Guy Rothwell, and Phil Weaver. We gratefully acknowledge support of this work through NERC ODP grant GST/02/1097.

REFERENCES

- Bouma, A.H., 1962. *Sedimentology of Some Flysch Deposits: A Graphic Approach to Facies Interpretation*: Amsterdam (Elsevier).
- Colley, S., and Thomson, J., 1985. Recurrent uranium relocations in distal turbidites emplaced in pelagic conditions. *Geochim. Cosmochim. Acta*, 49:2339–2348.
- _____, 1992. Behavior and mobility of U-series radionuclides in Madeira Abyssal Plain turbidites over the past 750,000 years. *Mar. Geol.*, 109:141–158.
- Colley, S., Thomson, J., and Toole, J., 1989. Uranium relocations and derivation of quasi-isochrons for a turbidite/pelagic sequence in the Northeast Atlantic. *Geochim. Cosmochim. Acta*, 53:1223–1234.
- Colley, S., Thomson, J., Wilson, T.R.S., and Higgs, N.C., 1984. Post-depositional migration of elements during diagenesis in brown clay and turbidite sequences in the North East Atlantic. *Geochim. Cosmochim. Acta*, 48:1223–1235.
- De Lange, G.J., 1992a. Distribution of exchangeable, fixed, organic and total nitrogen in interbedded turbiditic/pelagic sediments of the Madeira abyssal plain, eastern North Atlantic. *Mar. Geol.*, 109:95–114.
- _____, 1992b. Distribution of various extracted phosphorus compounds in the interbedded turbiditic/pelagic sediments of the Madeira Abyssal Plain, eastern North Atlantic. *Mar. Geol.*, 109:115–139.
- De Lange, G.J., Jarvis, I., and Kuipers, A., 1987. Geochemical characteristics and provenance of late Quaternary sediments from the Madeira

- Abyssal Plain, North Atlantic. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:147–165.
- De Lange, G.J., Middelburg, J.J., Jarvis, I., and Kuijpers, A., 1989. Geochemical characteristics and provenance of late Quaternary sediments from the Madeira and southern Nares Abyssal Plains (North Atlantic). In Schuttenhelm, R.T.E., Auffret, G.A., Buckley, D.E., Cranston, R.E., Murray, C.N., Shephard, L.E., and Spijkstra, A.E. (Eds.), *Geoscience Investigations of Two North Atlantic Abyssal Plains—The ESOPE International Expedition*: Luxembourg (Commission of the European Community), 2:785–851.
- Govindaraju, K., 1994. 1994 compilation of working values and sample description for 383 geostandards. *Geostand. Newslet.*, 18 (spec. iss.).
- Jarvis, I., 1992. Sample preparation for ICP-MS. In Jarvis, K.E., Gray, A.L., and Houk, R.S. (Eds.), *Handbook of Inductively Coupled Plasma Mass Spectrometry*: Glasgow (Blackie), 172–224.
- Jarvis, I., and Higgs, N., 1987. Trace-element mobility during early diagenesis in distal turbidites: late Quaternary of the Madeira Abyssal Plain, N Atlantic. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:179–214.
- Jarvis, I., and Jarvis, K.E., 1992. Inductively coupled plasma-atomic emission spectrometry in exploration geochemistry. In Hall, G.E.M. (Ed.), *J. Geochem. Expl.*, 44:139–200.
- Jones, K.P.N., 1988. Studies of fine-grained, deep-sea sediments [Ph.D. thesis]. Univ. of Cambridge, UK.
- Jones, K.P.N., McCave, I.N., and Weaver, P.P.E., 1992. Textural and dispersal patterns of thick mud turbidites from the Madeira Abyssal Plain. *Mar. Geol.*, 107:149–173.
- Kidd, R.B., Hunter, P.M., and Simm, R.W., 1987. Turbidity-current and debris-flow pathways to the Cape Verde Basin: status of long-range sidescan sonar (GLORIA) surveys. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:33–48.
- Kuijpers, A., and Weaver, P.P.E., 1985. Deep-sea turbidites from the northwest African continental margin. *Dtsch. Hydrogr. Z.*, 38:147–164.
- McArthur, J.M., Tyson, R.V., Thomson, J., and Matthey, D., 1992. Early diagenesis of marine organic matter: alteration of the carbon isotopic composition. *Mar. Geol.*, 105:51–61.
- McCave, I.N., and Jones, K.P.N., 1988. Deposition of ungraded muds from high-density non-turbulent turbidity currents. *Nature*, 333:250–252.
- Middelburg, J.J., 1993. Turbidites provide a unique opportunity to study diagenetic processes. *Geol. Mijnbouw*, 72:15–21.
- Middelburg, J.J., and De Lange, G.J., 1988. Geochemical characteristics as indicators of the provenance of Madeira Abyssal Plain turbidites: a statistical approach. *Oceanol. Acta*, 11:159–165.
- Pearce, T.J., 1991. The geology, geochemistry, sedimentology and provenance of Late Quaternary turbidites, Madeira Abyssal Plain [Ph.D. thesis]. CNA, Kingston Polytechnic.
- Pearce, T.J., and Jarvis, I., 1992a. Applications of geochemical data to modelling sediment dispersal patterns in distal turbidites: Late Quaternary of the Madeira Abyssal Plain. *J. Sediment. Petrol.*, 62:1112–1129.
- _____, 1992b. Composition and provenance of turbidite sands: Late Quaternary, Madeira Abyssal Plain. *Mar. Geol.*, 109:21–53.
- _____, 1995. High-resolution chemostratigraphy of Quaternary distal turbidites: a case study of new methods for the analysis and correlation of barren sequences. In Dunay, R.E., and Hailwood, E.A. (Eds.), *Non-bioservatigraphical Methods of Dating and Correlation*. Geol. Soc. Spec. Publ. London, 89:107–143.
- Rothwell, R.G., Pearce, T.J., and Weaver, P.P.E., 1992. Late Quaternary evolution of the Madeira Abyssal Plain, Canary Basin, NE Atlantic. *Basin Res.*, 4:103–131.
- Schmincke, H.-U., 1976. The geology of the Canary Islands. In Kunkel, G. (Ed.), *Biogeography and Ecology in the Canary Islands*: The Hague (W. Junk), 67–184.
- _____, 1982. Volcanic and chemical evolution of the Canary Islands. In von Rad, U., Hinz, K., Sarnthein, M., and Seibold, E. (Eds.), *Geology of the Northwest African Continental Margin*: Berlin (Springer), 273–306.
- _____, 1994. *Geological Field Guide: Gran Canaria* (6th ed.): Kiel, Germany (Pluto Press).
- Schmincke, H.-U., Weaver, P.P.E., Firth, J.V., et al., 1995. *Proc. ODP, Init. Repts.*, 157: College Station, TX (Ocean Drilling Program).
- Searle, R.C., 1987. Regional setting and geophysical characterization of the Great Meteor East area in the Madeira Abyssal Plain. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:49–70.
- Stow, D.A.V., and Shanmugam, G., 1980. Sequences of structures in fine-grained turbidites: comparison of recent deep-sea and ancient flysch sediments. *Sediment. Geol.*, 25:23–42.
- Thomson, J., Colley, S., Higgs, N.C., Hydes, D.J., Wilson T.R.S., and Sørensen, J., 1987. Geochemical oxidation fronts in NE Atlantic distal turbidites and their effects in the sedimentary record. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:167–177.
- Thomson, J., Higgs, N.C., Croudace, I.W., Colley, S., and Hydes, D.J., 1993. Redox zonation of elements at an oxic/post-oxic boundary in deep-sea sediments. *Geochim. Cosmochim. Acta*, 57:579–595.
- Thomson, J., Higgs, N.C., Jarvis, I., Hydes, D.J., Colley, S., and Wilson, T.R.S., 1986. The behaviour of manganese in Atlantic carbonate sediments. *Geochim. Cosmochim. Acta*, 50:1807–1818.
- Totland, M.M., Jarvis, I., and Jarvis, K.E., 1992. An assessment of dissolution techniques for the analysis of geological samples by plasma spectrometry. *Chem. Geol.*, 95:35–62.
- Weaver, P.P.E., 1993. High resolution stratigraphy of marine Quaternary sequences. In Hailwood, E.A., and Kidd, R.B. (Eds.), *High Resolution Stratigraphy*. Geol. Soc. Spec. Publ. London, 70:137–153.
- Weaver, P.P.E., Buckley, D.E., and Kuijpers, A., 1989. Geological investigations of ESOPE cores from the Madeira Abyssal Plain. In Schuttenhelm, R.T.E., Auffret, G.A., Buckley, D.E., Cranston, R.E., Murray, C.N., Shephard, L.E., and Spijkstra, A.E. (Eds.), *Geoscience Investigations of Two North Atlantic Abyssal Plains: The ESOPE International Expedition*: Luxembourg (Commission of the European Community), 1:535–555.
- Weaver, P.P.E., and Kuijpers, A., 1983. Climatic control of turbidite deposition on the Madeira Abyssal Plain. *Nature*, 306:360–363.
- Weaver, P.P.E., Masson, D.G., and Kidd, R.B., 1994. Slumps, slides and turbidity currents: sea-level change and sedimentation in the Canary Basin. *Geoscientist*, 4:14–16.
- Weaver, P.P.E., and Rothwell, R.G., 1987. Sedimentation on the Madeira Abyssal Plain over the last 300,000 years. In Weaver, P.P.E., and Thomson, J. (Eds.), *Geology and Geochemistry of Abyssal Plains*. Geol. Soc. Spec. Publ. London, 31:71–86.
- Weaver, P.P.E., Rothwell, R.G., Ebbing, J., Gunn, D., and Hunter, P.M., 1992. Correlation, frequency of emplacement and source directions of megaturbidites on the Madeira Abyssal Plain. *Mar. Geol.*, 109:1–20.
- Weaver, P.P.E., Searle, R.C., and Kuijpers, A., 1986. Turbidite deposition and the origin of the Madeira Abyssal Plain. In Summerhayes, C.P., and Shackleton, N.J. (Eds.), *North Atlantic Palaeoceanography*. Spec. Publ. Geol. Soc. London, 21:131–143.
- Weaver, P.P.E., and Thomson, J., 1993. Calculation erosion by deep-sea turbidity currents during initiation and flow. *Nature*, 364:136–138.
- Weaver, P.P.E., Thomson, J., and Jarvis, I., 1989. The geology and geochemistry of Madeira Abyssal Plain sediments: a review. In Freeman, T.J. (Ed.), *Advances in Underwater Technology, Ocean Science and Offshore Engineering* (Vol. 18): London (Graham and Trotman), 51–78.
- Wilson, T.R.S., Thomson, J., Colley, S., Hydes, D.J., Higgs, N.C., and Sørensen, J., 1985. Early organic diagenesis: the significance of progressive subsurface oxidation fronts in pelagic sediments. *Geochim. Cosmochim. Acta*, 49:811–822.
- Wilson, T.R.S., Thomson, J., Hydes, D.J., Colley, S., Culkin, F., and Sørensen, J., 1986. Oxidation fronts in pelagic sediments: diagenetic formation of metal-rich layers. *Science*, 232:972–975.

Date of initial receipt: 8 July 1996

Date of acceptance: 8 January 1997

Ms 157SR-129