

5. DATA REPORT: LATE QUATERNARY CALCAREOUS NANNOFOSSILS FROM THE NORTHWESTERN PACIFIC OCEAN, HOLES 1150A AND 1151C, LEG 186¹

Kyoko Hagino² and Hisatake Okada³

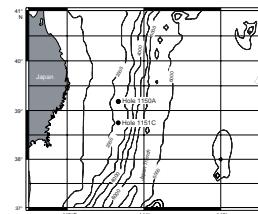
ABSTRACT

To assess the paleoceanographic potential of Leg 186 sediments, we investigated Quaternary calcareous nannofossil flora at Sites 1150 and 1151 in the Japan Trench. Because of the frequent occurrence of barren intervals and the lack of oxygen isotope data, a detailed paleoceanography is not feasible for these cores. We limited our study to the upper 26.07 m of the section from Hole 1150A and the upper 21.01 m of the section from Hole 1151C. The studied samples from Cores 186-1150A-1H through 3H are younger than 0.085 Ma. Core 186-1151C-1H (upper 1.92 meters below seafloor [mbsf]) is younger than 0.085 Ma, and samples between 2H-7, 5–7 cm, and 3H-CC, 5–7 cm, (9.99–21.01 mbsf) are older than 0.245 Ma and younger than 0.408 Ma.

INTRODUCTION

During Ocean Drilling Program (ODP) Leg 186, we drilled six holes at two sites, Sites 1150 ($39^{\circ}11'N$, $143^{\circ}20'E$) and 1151 ($38^{\circ}45'N$, $143^{\circ}20'E$) on the deep-sea terrace of the landward side of the Japan Trench (Fig. F1) (Sacks, Suyehiro, Acton, et al., 2000). These sites are located under the influence of the Oyashio Extension, the northwest boundary current of the Subpolar Gyre in the Pacific Ocean. Off the coast of northeast Japan, the Subpolar Gyre (Oyashio Current) meets head-on with the Subtropical Gyre (Kuroshio Current) (Tomczak and Godfrey, 1994).

F1. ODP sites, p. 9.



¹Hagino, K., and Okada, H., 2003. Data report: Late Quaternary calcareous nannofossils from the northwestern Pacific Ocean, Holes 1150A and 1151C, Leg 186. In Suyehiro, K., Sacks, I.S., Acton, G.D., and Oda, M. (Eds.), *Proc. ODP, Sci. Results*, 186, 1–15 [Online]. Available from World Wide Web: <http://www-odp.tamu.edu/publications/186_SR/VOLUME/CHAPTERS/108.PDF>. [Cited YYYY-MM-DD]

²Research Fellow of the Japan Society for the Promotion of Science, Department of Earth and Planetary Sciences, Graduate School of Environmental Studies, Nagoya University, Furo-cho, Chikusa-Ku, Nagoya 464-8602, Japan.

hagino@geobio.eps.nagoya-u.ac.jp

³Division of Earth and Planetary Sciences, Graduate School of Science, Hokkaido University, Sapporo N10 W8, 060-0810, Japan.

Distribution of the living calcareous nannoplankton is controlled by the water masses (e.g., Okada and Honjo, 1973a), and the stratigraphic variation in the Quaternary nannoflora reflects the environmental change in surface water (e.g., Chinzei et al., 1987). Therefore, detailed investigation of the nannofossil assemblages at the surface water frontal zones is important in monitoring Quaternary paleoceanography. Here, we present a data set of Quaternary calcareous nannofossil assemblages in the subpolar frontal zone of the western Pacific Ocean that supersedes the preliminary results described by Shipboard Scientific Party (2000a, 2000b).

MATERIALS AND METHODS

A total of 183 sediment samples were collected at 30-cm intervals from Cores 186-1150A-1H through 3H and 186-1151C-1H through 3H of the two Japan Trench sites Holes 1150A and 1151C (Fig. F1). Sediment samples were prepared as smear slides using standard techniques (Bown and Young, 1998), and calcareous nannofossils were examined at 1000 \times magnification under a Nikon E600 polarizing light microscope.

To understand a general trend in the biostratigraphy, samples were examined at intervals of approximately two samples per section in the first stage of investigation. Following this observation, the remaining samples were examined to refine the biostratigraphy.

For the Quaternary, relative abundances of placolith-type species are useful as datum events (e.g., Hine and Weaver, 1998). *Florisphaera profunda* is usually the only lower photic zone dweller that is preserved well in marine sediment, and it often dominates the fossil assemblage. Lower photic zone dwellers respond to environmental changes in a different manner than upper photic zone dwellers (Molfino and McIntyre, 1990a, 1990b; Hagino et al., 2000). For the purpose of biostratigraphy and paleoceanography, more than 300 upper photic zone specimens were identified and counted and the number of *F. profunda* coccoliths present in the same view fields were counted separately.

TAXONOMIC REMARKS

Late Quaternary calcareous nannoflora usually contain abundant placolith-type nannofossils that are produced by *Emiliania huxleyi* and *Gephyrocapsa* species. Placoliths of *Gephyrocapsa* species display a great degree of size variation (Bollmann, 1997; Matsuoka and Okada, 1989, 1990). Downcore size variation of *Gephyrocapsa* species is useful for stratigraphy (Matsuoka and Okada, 1989; Erba 1995; Okada and Wells, 1997); however, classification and taxonomy of the genus *Gephyrocapsa* is in a state of confusion. Because this study does not intend to refine the classification system, *Gephyrocapsa* species are classified into the following four size categories: very small (<2.0 μm), small (2.0–3.0 μm), medium (3.0–5.0 μm), and large (>5.0 μm).

In the small *Gephyrocapsa* specimens, only *Gephyrocapsa caribbeanica* is identified at the species level based on the closed central area. The other small and very small placoliths having a bridge and a wide central area are classified into *Gephyrocapsa* (small) or *Gephyrocapsa* (very small) categories based on size.

Under the light microscope, it is difficult to distinguish the small and very small *Gephyrocapsa* specimens that were lost or did not develop

bridge elements from the small and very small *Reticulofenestra* specimens, respectively. Moreover, it is difficult to identify the small and very small placoliths at species level under the light microscope. Therefore, the small and very small *Gephyrocapsa* specimens without bridge and the small and very small *Reticulofenestra* specimens are classified into the placolith (small) or placolith (very small) categories based on size.

RESULTS AND COMMENTS

Biostratigraphy and Paleoceanographic Comments

Hole 1150A

As a preliminary step, a total of 38 selected samples out of 108 samples were examined to check biostratigraphy, and 25 of them contained sufficient numbers of calcareous nannofossils to study the floral composition (Table T1).

The latest “standard” nannofossil event is the first occurrence (FO) of *E. huxleyi* at 0.268 Ma (Thierstein et al., 1977). *E. huxleyi* is present in Sample 186-1150A-3H-7, 0–4 cm (25.52 meters below seafloor [mbsf]), which is the lowest sample studied here (Table T1). The Shipboard Scientific Party (2000a) estimated the FO of *E. huxleyi* at 46.33–55.73 mbsf. Therefore, the entire part of the studied sections are younger than 0.268 Ma. The base of the *E. huxleyi* acme Zone is a diachronous event and was reported between 85 ka in low latitudes and 73 ka in transitional waters (Thierstein et al., 1977). This event was identified by the reversal in abundance of *G. caribbeanica*/*E. huxleyi* (e.g., Thierstein et al., 1977) or by reversal in abundance of *Gephyrocapsa muellerae*/*E. huxleyi* (Flores et al., 2000). *E. huxleyi* is more abundant at the lowest sample studied here than *G. caribbeanica* and *G. muellerae* from Core 186-1150A-1H up through Core 3H. Therefore, the entire range of studied samples belong to the *E. huxleyi* acme Zone (Fig. F2).

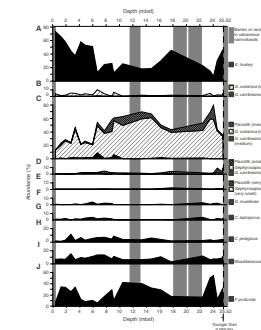
The barren intervals (Fig. F2) likely indicate the weakening of the Kuroshio Current and may correspond to glacial–stadial intervals, but without oxygen isotope data the hypothesis is difficult to test.

Hole 1151C

Seventy-five samples collected from Cores 186-1151C-1H through 3H were examined, and 52 of them contained sufficient numbers of calcareous nannofossils to study the floral composition (Table T2). Reworked specimens of *Pseudoemiliania lacunosa* are commonly present in the two upper cores (186-1151C-1H and 2H) but are scarce in the lower core (3H) (Table T2; Fig. F3). The preliminary study observed the last occurrence (LO) of *P. lacunosa* (0.408 Ma) in Sample 186-1151C-5H-3, 98 cm (Shipboard Scientific Party, 2000b). Therefore, the entire range of studied sections is younger than 0.408 Ma. Weaver and Thomson (1993) reported an abrupt decrease of *G. caribbeanica* at the boundary between marine oxygen isotope Stages (MISs) 7 and 8 (0.245 Ma). The abrupt decrease of *G. caribbeanica* observed between 9.99 and 8.85 mbsf in Core 186-1151C-3H is likely to correspond to this event. (Table T2; Fig. F3). Based on the LO of *P. lacunosa* (0.408 Ma) (Shipboard Scientific Party, 2000b) and the abrupt decrease of *G. caribbeanica* (0.245 Ma), the lowest sample studied here can be estimated as 0.36 Ma in age (horizon-

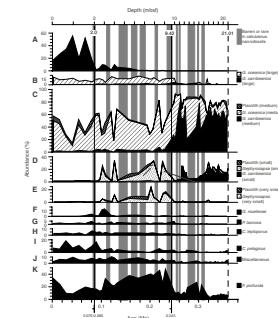
T1. Calcareous nannofossils in Cores 186-1150A-1H through 3H, p. 13.

F2. Stratigraphic variation, Hole 1150A, p. 10.



T2. Calcareous nannofossils in Cores 186-1151C-1H through 3H, p. 14.

F3. Stratigraphic variation, Hole 1151C, p. 11.



tal axis of Fig. F3). In this study, the FO of *E. huxleyi* was observed in Sample 186-1151C-2H-5, 65–67 cm (8.56 mbsf); however, the event has been reported between Samples 2H-CC (12.10 mbsf) and 3H-CC (21.44 mbsf) (Shipboard Scientific Party, 2000b). Because of the vulnerability to dissolution, the FO of *E. huxleyi* is often difficult to identify in a poorly preserved assemblage. Therefore, the true FO of this species may be lower than the identified FO of *E. huxleyi* in this study. *E. huxleyi* is very abundant in the entire Core 186-1151C-1H, except the for Sample 186-1151C-1H-CC. Sections 186-1151C-1H-1 and 1H-2, therefore, are assigned to the *E. huxleyi* acme Zone.

Calcidiscus leptoporus prefers tropical to transitional waters; on the other hand, *Coccolithus pelagicus* prefers Arctic to subarctic waters (e.g., Winter et al., 1994). In the central Pacific Ocean, *Florisphaera profunda* is abundant in the lower photic zone of the tropical to transitional waters and is barren in the subarctic Oyashio Extension water (Okada and Honjo, 1973a). In the studied samples at this site, *C. leptoporus*, *C. pelagicus*, and *F. profunda* are the major species (Table T2; Fig. F3). Therefore, it is clear that the surface water at Site 1151 has been affected by both warm Kuroshio and cold Oyashio Extensions during the last 0.38 m.y.

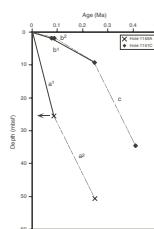
Stratigraphic variation of the lower photic zone species including *F. profunda* is useful as a paleoceanographic indicator, and a lower abundance of lower photic zone species can be interpreted as an indicator of surface water mixing that triggers a shallower nutricline and higher primary productivity (Molfino and McIntyre, 1990a, 1990b; Okada and Matsuoka, 1996; Beaufort et al., 1997). Studies of living calcareous nanoplankton revealed that the relative abundance of lower photic zone species within the water column is controlled by the absolute abundance of the upper photic zone species (Hagino et al., 2000), and small placolith-bearing species flourish in the upper photic zone of the eutrophic equatorial Pacific (Okada and Honjo, 1973b; Hagino, 1999; Hagino and Okada, 2001). Although *Gephyrocapsa* (small) and *Gephyrocapsa* (very small) are only minor components of the flora in the uppermost 2 mbsf, their abundances increased significantly between 2 and 10 mbsf, showing a concordant stratigraphic trend to that of *F. profunda*. (Fig. F3). In the lower core (10–21 mbsf), *Gephyrocapsa* (small) and *Gephyrocapsa* (very small) are minor components and *Gephyrocapsa caribeanica* (small) becomes a major component, showing an opposite trend to that of *F. profunda* (Fig. F3). According to Molfino and McIntyre (1990b) and Okada (2000), *F. profunda* and small to very small *Gephyrocapsa* are good indicators of stratified and mixed photic layer conditions, respectively. The stratigraphic trend of *G. caribeanica* (small) is concordant with this theory, but the trends of *Gephyrocapsa* (small) and *Gephyrocapsa* (very small) are contrary to this theory.

The unavailability of oxygen isotope data and the presence of many barren intervals limits what can be learned about (Fig. F3) the paleoceanography from the nannoflora in this area.

Sedimentation Rate

Sedimentation rate was calculated based on the depth of the nannofossil datums (Fig. F4). The base of the *E. huxleyi* acme Zone has been reported from 0.073 to 0.085 Ma (e.g., Thierstein et al., 1977). The only datum information in Cores 186-1150A-1H through 3H is that the deepest studied sample (Sample 186-1150A-3H-7, 0–4 cm; 25.52 mbsf), is younger than 0.085 Ma. This result indicates that the sedimentation

F4. Age-depth plot of biostratigraphic events, Holes 1150A and 1151C, p. 12.



rate between Cores 186-1150A-1H and 3H (upper 25.52 mbsf) is >300 m/m.y. The Shipboard Scientific Party (2000a) estimated that the sedimentation rate for the upper 50.72 mbsf is 204 m/m.y., based on the FO of *E. huxleyi*. Therefore, the sedimentation rate between 25.52 and 50.72 mbsf would be <94 m/m.y. (Fig. F4).

In Hole 1151C, the base of the *E. huxleyi* acme Zone (0.076–0.085 Ma) lies within Samples 186-1151C-1H-2, 35–37 cm (1.92 mbsf), through 1H-CC (2.08 mbsf); therefore, the sedimentation rate is >23.5 m/m.y. and <26.3 m/m.y in the upper 2.0 mbsf. An abrupt decrease of *G. caribbeanica* (0.245 Ma) observed between Samples 186-1151C-2H-5, 95–97 cm (8.85 mbsf), and 2H-6, 65–67 cm (9.99 mbsf), indicates that the sedimentation rate is >43.9 m/m.y. and <46.4 m/m.y. (between 2.00 and 9.42 mbsf). The preliminary study reported the LO of *P. lacunosa* (0.408 Ma) in Samples 186-1151C-5H-3, 98 cm (23.69 mbsf), and 4H-2, 98 cm (34.68 mbsf) (Shipboard Scientific Party, 2000b). On the basis of an abrupt decrease of *G. caribbeanica* and the LO of *P. lacunosa* (0.408 Ma), the sedimentation rate between 9.42 and 29.18 mbsf can be estimated as 121.1 m/m.y.

ACKNOWLEDGMENTS

This research used samples and/or data provided by the Ocean Drilling Program (ODP). ODP is sponsored by the U.S. National Science Foundation (NSF) and participating countries under management of Joint Oceanographic Institutions (JOI), Inc. Funding for this research was provided in part by Research Fellowships from the Japan Society for the Promotion Science for Young Scientists (grant no. 0104014) and by a Grant-in-Aid for Scientific Research from the Ministry of Education, Culture, Sports, Science, and Technology (grant no. 11691113).

We are truly grateful to Gary Acton, Lorri Peters, and Bryan C. Ladner for their critical reviews and constructive opinions. We wish to express thanks to T. Sakamoto (Hokkaido University) and Y. Tanaka (National Institute of Advanced Industrial Science and Technology) for their helpful suggestions.

REFERENCES

- Beaufort, L., Lancelot, Y., Camberlin, P., Cayre, O., Vincent, E., Bassinot, F., and Labeyrie, L., 1997. Insolation cycles as a major control of equatorial Indian Ocean primary production. *Science*, 278:1451–1454.
- Böllmann, J., 1997. Morphology and biogeography of *Gephyrocapsa* coccoliths in Holocene sediments. *Mar. Micropaleontol.*, 29:319–350.
- Bown, P.R., and Young, J.R., 1998. Techniques. In Bown, P.R. (Ed.), *Calcareous Nannofossil Biostratigraphy*: London (Champman and Hall), 16–28.
- Chinzei, K., Fujioka, K., Kitazato, H., Koizumi, I., Oba, T., Oda, M., Okada, H., Sakai, T., and Tanimura, Y., 1987. Postglacial environmental change of the Pacific Ocean off the coast of central Japan. *Mar. Micropaleontol.*, 11:273–291.
- Erba, E., 1995. Quantitative nannofossil biostratigraphy of Quaternary sequences from guyots in the central and western Pacific Ocean. In Haggerty, J.A., Premoli Silva, I., Rack, F., and McNutt, M.K. (Eds.), *Proc. ODP, Sci. Results*, 144: College Station, TX (Ocean Drilling Program), 3–20.
- Flores, J.A., Gersonde, R., Sierro, F.J., and Niebler, H.S., 2000. Southern Ocean Pleistocene calcareous nannofossil events: calibration with isotope and geomagnetic stratigraphies. *Mar. Micropaleontol.*, 40:377–402.
- Hagino, K., 1999. Population dynamics of living coccolithophores in relation to hydrography of the northwest and equatorial Pacific Ocean [Ph.D. thesis]. Hokkaido Univ., Sapporo.
- Hagino, K., and Okada, H., 2001. Morphological observations of living *Gephyrocapsa crassipons*. *J. Nannoplankton Res.*, 23:3–7.
- Hagino, K., Okada, H., and Matsuoka, H., 2000. Spatial dynamics of coccolithophore assemblages in the equatorial western-central Pacific Ocean. *Mar. Micropaleontol.*, 39:53–72.
- Hine, H., and Weaver, P.P.E., 1998. Quaternary. In Bown, P.R. (Ed.), *Calcareous Nannofossil Biostratigraphy*: London (Champman and Hall), 266–283.
- Matsuoka, H., and Okada, H., 1989. Quantitative analysis of Quaternary nannoplankton in the subtropical northwestern Pacific Ocean. *Mar. Micropaleontol.*, 14:97–118.
- _____, 1990. Time-progressive morphometric changes of the genus *Gephyrocapsa* in the Quaternary sequence of the tropical Indian Ocean, Site 709. In Duncan, R.A., Backman, J., Peterson, L.C., et al., *Proc. ODP, Sci. Results*, 115: College Station, TX (Ocean Drilling Program), 255–270.
- Molfino, B., and McIntyre, A., 1990a. Nutricline variations in the equatorial Atlantic coincident with the Younger Dryas. *Paleoceanography*, 5:997–1008.
- _____, 1990b. Precessional forcing of nutricline dynamics in the equatorial Atlantic. *Science*, 249:766–769.
- Okada, H., 2000. Neogene and Quaternary calcareous nannofossils from the Blake Ridge, Sites 994, 995, and 997. In Paull, C.K., Matsumoto, R., Wallace, P.J., and Dillon, W.P. (Eds.), *Proc. ODP, Sci. Results*, 164: College Station, TX (Ocean Drilling Program), 331–341.
- Okada, H., and Honjo, S., 1973a. The distribution of oceanic coccolithophorids in the Pacific. *Deep-Sea Res. Part A*, 20:355–374.
- _____, 1973b. Distribution of coccolithophorids in the North and equatorial Pacific Ocean: quantitative data on samples collected during Leg 30, Oshoro-Maru, 1968 and Leg HK69-4, Hakuho Maru, 1969. *WHOI Tech. Rep.*, WHOI-73-81.
- Okada, H., and Matsuoka, M., 1996. Lower-photic nannoflora as an indicator of the late Quaternary monsoonal palaeo-record in the tropical Indian Ocean. In Moguillevsky, A., and Whatley, R., (Eds.), *Proc. ODP Mar. Biosphere. Int. Conf.*: Aberystwyth, UK (Univ. of Wales, Aberystwyth Press), 231–245.
- Okada, H., and Wells, P., 1997. Late Quaternary nannofossil indicators of climate change in two deep-sea cores associated with the Leeuwin Current off Western Australia. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 131:413–432.

- Sacks, I.S., Suyehiro, K., Acton, G.D., et al., 2000. *Proc. ODP, Init. Repts.*, 186 [CD-ROM]. Available from: Ocean Drilling Program, Texas A&M University, College Station TX 77845-9547, USA.
- Shipboard Scientific Party, 2000a. Site 1150. In Sacks, I.S., Suyehiro, K., Acton, G.D., et al., *Proc. ODP, Init. Repts.*, 186, 1-209 [CD-ROM]. Available from: Ocean Drilling Program, Texas A&M University, College Station TX 77845-9547, USA.
- _____, 2000b. Site 1151. In Sacks, I.S., Suyehiro, K., Acton, G.D., et al., *Proc. ODP, Init. Repts.*, 186, 1-125 [CD-ROM]. Available from: Ocean Drilling Program, Texas A&M University, College Station TX 77845-9547, USA.
- Tomczak, M., and Godfrey, J.S., 1994. *Regional Oceanography: An Introduction*: Oxford (Pergamon Press).
- Thierstein, H.R., Geitzenauer, K., Molfino, B., and Shackleton, N.J., 1977. Global synchronicity of late Quaternary coccolith datum levels: validation by oxygen isotopes. *Geology*, 5:400-404.
- Weaver, P.P.E., and Thomson, J., 1993. Calculation erosion by deep-sea turbidity currents during initiation and flow. *Nature*, 364:136-138.
- Winter, A., Jordan, R.W., and Roth, P.H., 1994. Biogeography of living coccolithophores. In Winter, A., and Siesser, W.G. (Eds.), *Coccolithophores*: Cambridge (Cambridge Univ. Press), 161-177.

TAXONOMIC LIST

Calcidiscus leptoporus (Murray and Blackman, 1989) Loeblich and Tappan, 1978.
Calcidiscus macintyreai (Bukry and Bramlette, 1969) Loeblich and Tappan, 1978.
Ceratolithus cristatus Kamptner, 1950.
Coccolithus pelagicus (Wallich, 1877) Schiller, 1930.
Emiliania huxleyi (Lohmann, 1902) Hay and Mohler in Hay et al., 1967.
Florisphaera profunda Okada and McIntyre, 1977.
Gephyrocapsa caribbeana Boudreaux and Hay, 1967.
Gephyrocapsa oceanica Kamptner 1943.
Gephyrocapsa muellerae Bréhéret et al., 1987.
Helicosphaera carteri (Wallich, 1877) Kamptner, 1954.
Helicosphaera inversa Gartner, 1980.
Neosphaera coccolithomorpha Lecal-Schlauder, 1950.
Oolithotus fragilis (Lohmann, 1912) Martini and Müller, 1972.
Pseudoemiliania lacunosa (Kamptner, 1963) Gartner, 1969.
Reticulofenestra asanoi Sato and Takayama, 1992.
Rhabdosphaera clavigera Murray and Blackman, 1898.
Syracosphaera historica Kamptner, 1941.
Syracosphaera pulchra Lohmann, 1902.
Umbellosphaera irregularis Paasche in Markali and Paasche, 1955.
Umbellosphaera tenuis (Kamptner, 1937) Paasche in Markali and Paasche, 1955.
Umbilicosphaera hulbertiana Gaarder, 1970.
Umbilicosphaera sibogae var. *foliosa* (Kamptner, 1963) Okada and McIntyre, 1977.
Umbilicosphaera sibogae var. *sibogae* (Weber-van Bosses, 1901) Gaarder, 1970.

Figure F1. Location of the ODP sites studied in this report. Contours = meters below sea level.

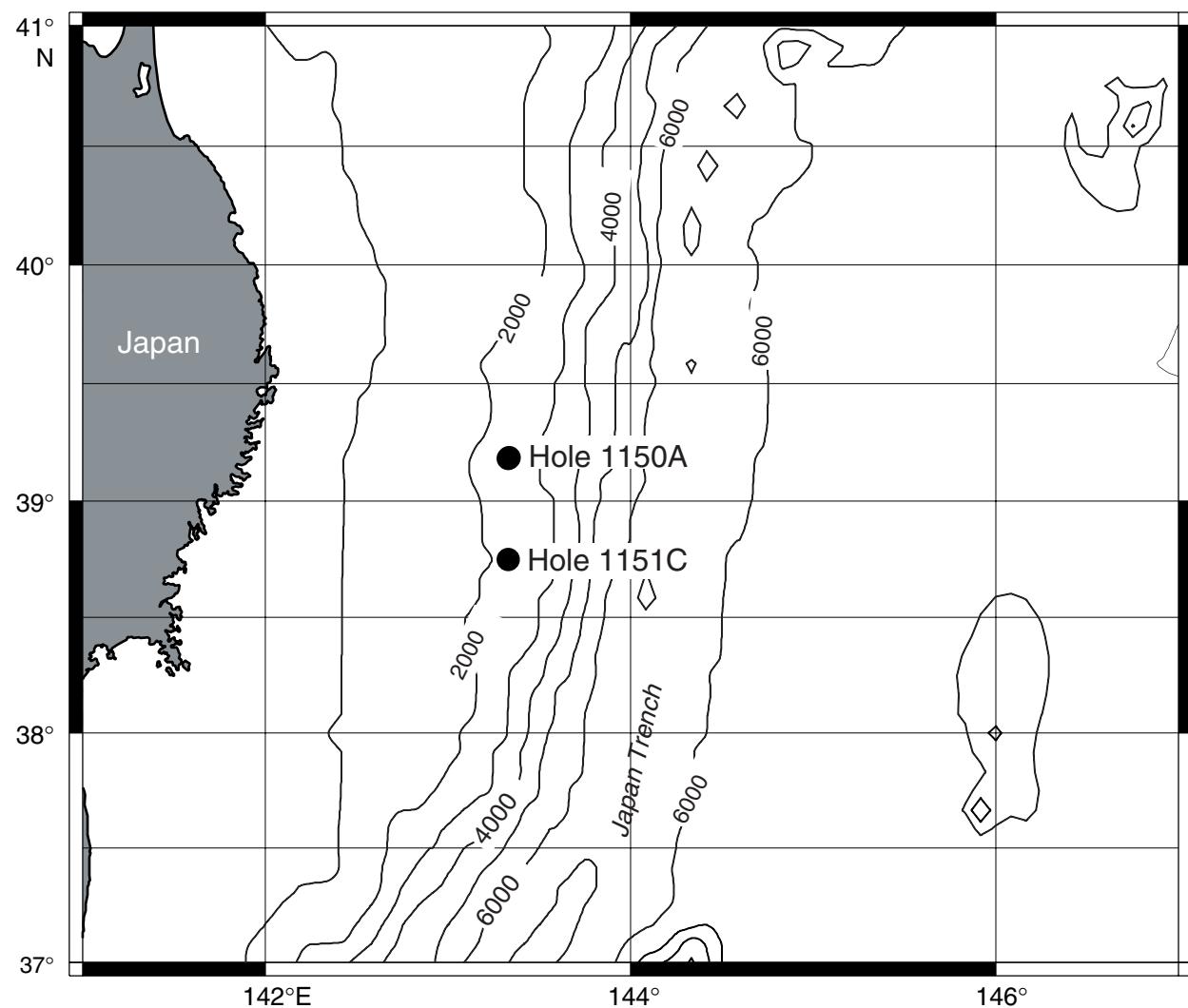


Figure F2. Stratigraphic variation in percentage abundance of the major components of the flora observed in Hole 1150A. Abundance of figs: A–I. Percentage abundance of each species within the upper photic zone flora. J. Percentage abundance of *F. profunda* (lower photic zone dweller) within the entire calcareous nannofossil assemblage.

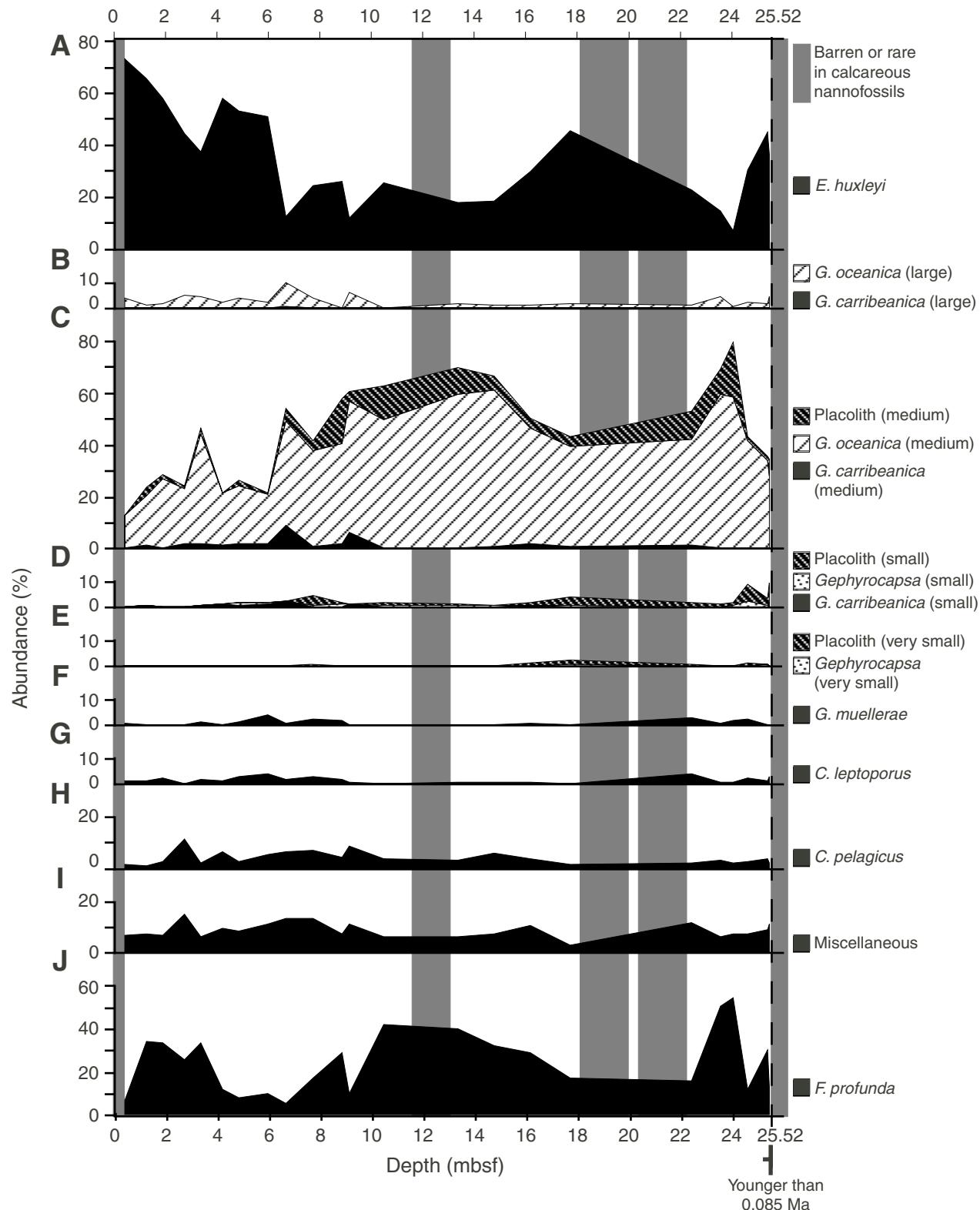


Figure F3. Stratigraphic variation in percentage abundance of the major components of the flora observed in Hole 1151C. A–J. Percentage abundance of each species within the upper photic zone flora. K. Percentage abundance of *F. profunda* (lower photic zone dweller) within the entire calcareous nannofossil assemblage.

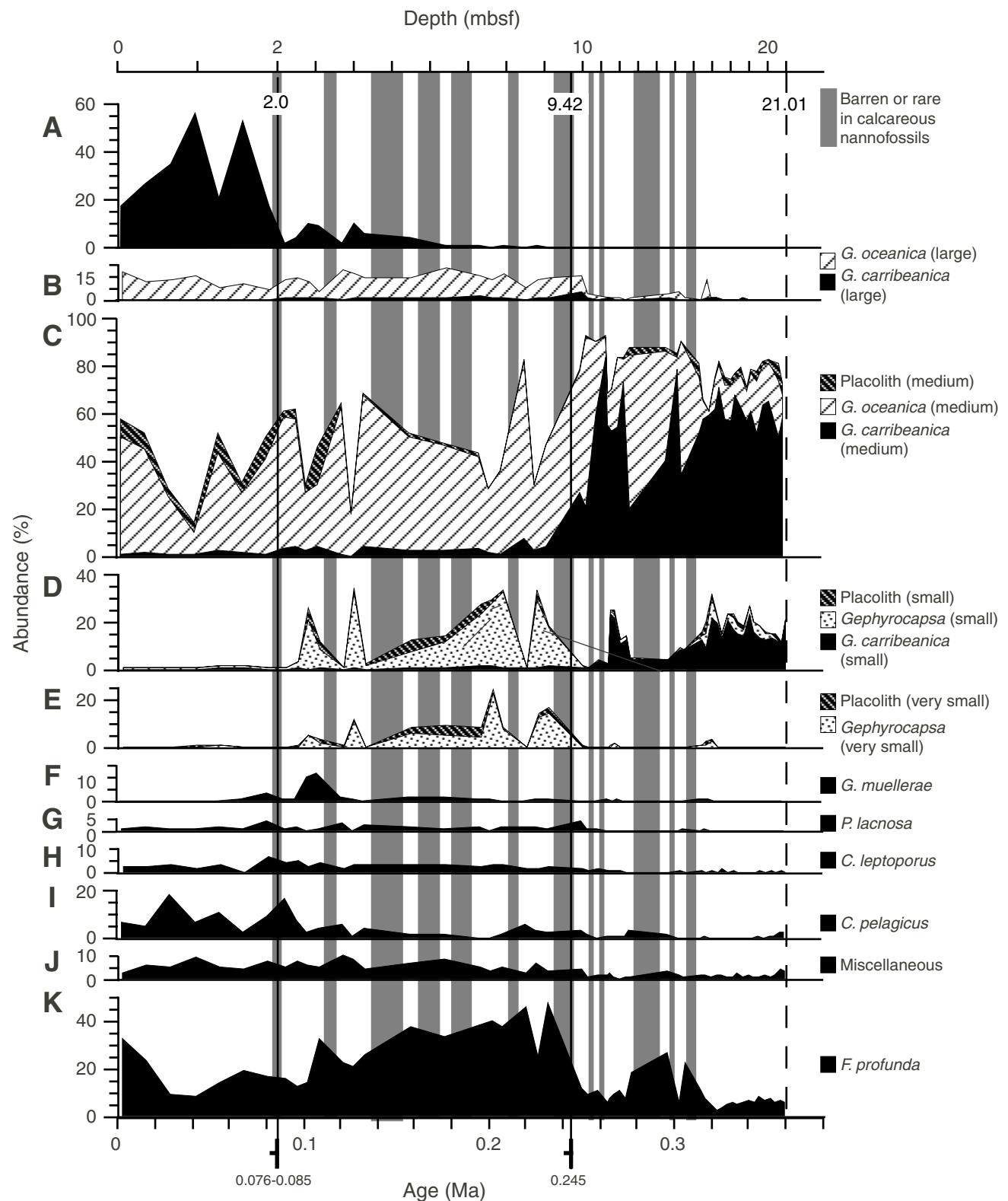


Figure F4. Age-depth plot of biostratigraphic events in Holes 1150A and 1151C. Solid line a^1 = age-depth relationship in Cores 186-1150A-1H through 3H in Hole 1150A, where the age of Sample 186-1150A-3H, 0–4 cm (25.52 mbsf), is assumed to be 0.085 Ma. The age control point at 25.52 mbsf in Hole 1150A can be moved leftward along the horizontal axis because the true acme of *E. huxleyi* is expected to be found in a lower sample. Dotted line a^2 = age-depth relationship between 25.52 and 50.72 mbsf based on the lowest studied sample (25.52 mbsf) and the FO of *E. huxleyi* observed by Shipboard Scientific Party (2000a). Kinked solid line b^1 and kinked dotted line b^2 = age-depth relationship between 9.42 and 0 mbsf in Hole 1151C, when the age of the 2.00-mbsf sample is posited as 0.073 and 0.085 Ma, respectively. The dotted line c = age-depth relationship between 34.68 and 9.42 mbsf in Hole 1151C based on the abrupt decrease of *G. caribeanica* (9.42 mbsf) and the LO of *P. lacunosa* (Shipboard Scientific Party, 2000b).

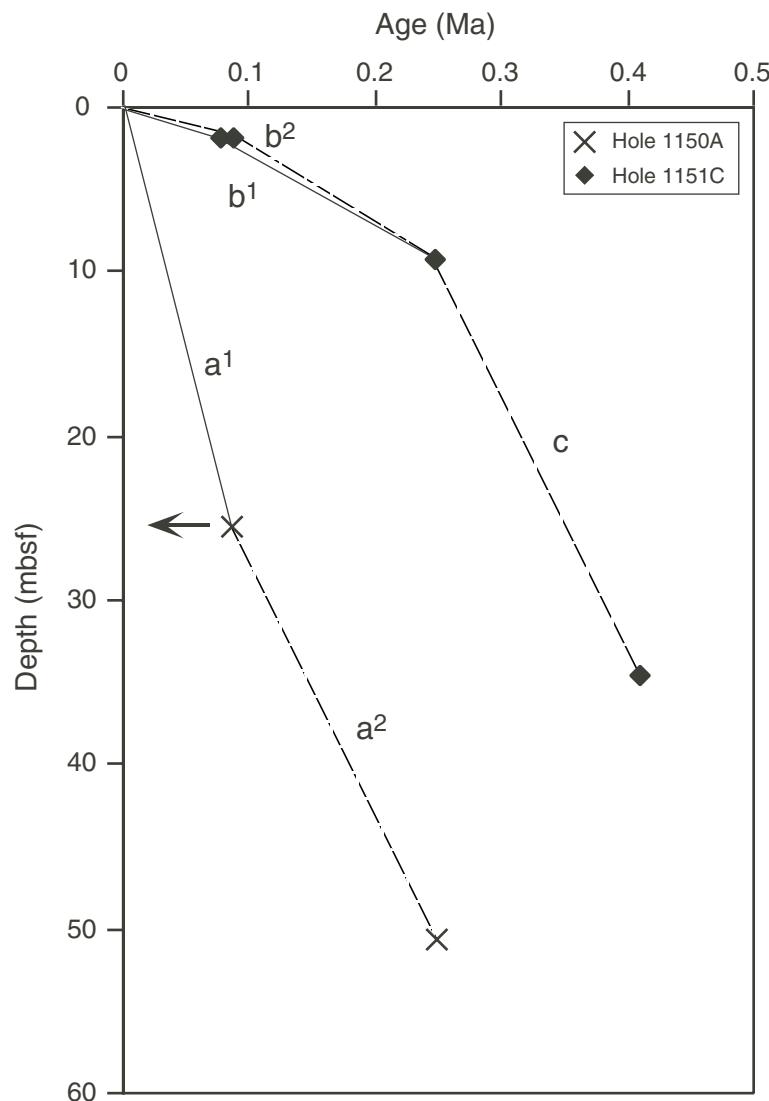


Table T1. Quantitative distribution of calcareous nannofossils, Cores 186-1150A-1H through 3H.

Table T2. Quantitative distribution of calcareous nannofossils, Cores 186-1151C-1H through 3H. (Continued on next page.)

Core, section, interval (cm)	Average depth (mbsf)	<i>Calcidiscus leptoporus</i>	<i>Ceratolithus cristatus</i>	<i>Coccolithus pelagicus</i>	<i>Emiliana huxleyi</i>	<i>Gephyrocapsa caribbeanica</i> (large)	<i>Gephyrocapsa caribbeanica</i> (medium)	<i>Gephyrocapsa caribbeanica</i> (small)	<i>Gephyrocapsa muellerae</i>	<i>Gephyrocapsa oceanica</i> (large)	<i>Gephyrocapsa oceanica</i> (medium)	<i>Gephyrocapsa (small)</i>	<i>Gephyrocapsa (very small)</i>	<i>Helicosphaera carteri</i>	<i>Helicosphaera inversa</i>	<i>Neosphaera coccolithomorpha</i>	<i>Pseudemiliania acunosa</i>	<i>Placolith</i> (medium)	<i>Placolith</i> (small)	<i>Placolith</i> (very small)	<i>Sphenolithus</i> spp.	<i>Syracosphaera histrica</i>	<i>Syracosphaera pulchra</i>	<i>Syracosphaera</i> spp.	<i>Umbilicosphaera irregularis</i>	<i>Umbilicosphaera sibogae</i> var. <i>sibogae</i>	<i>Umbilicosphaera sibogae</i> var. <i>foliosa</i>	Subtotal	<i>Florisphaera profunda</i>	Total	Remarks
186-1151C-1H-1, 5-7	0.06	10	22	60		5			38	165		3		3			2	24					1			333	162	495			
1H-1, 35-37	0.37	8	16	87		6			22	139		1		2			3	21					1			318	98	416			
1H-1, 65-67	0.68	13	64	124		5			28	85		1		3			1	12					1			348	35	383			
1H-1, 95-97	0.99	6	23	207		6			37	32		2		7			1	15	1	3			24			1	365	32	397		
1H-1, 125-127	1.30	14	41	77		10	1		19	154		4	4	2			4	29					7	1	1	369	58	427			
1H-2, 5-7	1.61	2	8	176		6	1	2	22	84		3		1	1		2	14					4	1	1	330	79	409			
1H-2, 35-37	1.92	24	3	33	63	5			11	13	143	3		1			12	32					8			1	1	353	69	422	
1H-CC, 5-7	2.08																								0	0	0				
2H-1, 5-7	2.26	17	1	67	9	1	14		3	31	220	1	1	1	2		2	11	1				12	1		395	73	468	Barren		
2H-1, 35-37	2.54	20	31	20	1	17	2	2	36	219	9	2	12	1		6	15	1				7			2	403	56	459			
2H-1, 65-67	2.83	8	9	34	2	9			33	21	81	68	16	2			11	15	1				2	11	1	2	326	53	379		
2H-1, 95-97	3.12	13	14	32	1	16	3	37	10	85	25	5	2	1		2	52	11	7			1	9		1	327	158	485			
2H-1, 125-127	3.40																								0	0	0	Barren			
2H-2, 5-7	3.69	7	1	20	8	4			4	42	207	2	1	2			10	4	1				18			331	94	425			
2H-2, 35-37	3.98	11	3	36	1	1	1	1	36	58	108	37	3	1		6	3	6	4			1	2	18	2	1	334	88	422		
2H-2, 65-67	4.26	12	15	21	2	15			28	209	6					6	3	1				5			326	114	440				
2H-2, 95-97	4.55																								0	0	0	Barren			
2H-2, 125-127	4.84																								0	0	0	Rare			
2H-3, 5-7	5.12																								0	0	0	Barren			
2H-3, 32-34	5.38	13	5	16	2	9	2	5	28	162	23	21	1			3	3	15	8			17		1	334	202	536				
2H-3, 65-67	5.69																								0	0	0	Barren			
2H-3, 95-97	5.98																								0	0	0	Barren			
2H-3, 125-127	6.27	10	5	4	1	10	3	4	40	136	32	15	5			1	2	8	14			1	11	1	2	3	308	156	464		
2H-4, 5-7	6.55																								0	0	0	Barren			
2H-4, 36-38	6.84																								0	0	0	Barren			
2H-4, 65-67	7.13	9	1	5	4	12	6	2	27	126	69	15	4			4	5	13	13			6	1	1	323	202	525				
2H-4, 95-97	7.41	13	1	12	8	4	3	26	89	87	75					1	7	7	1	6		1	2	334	220	554					
2H-4, 125-127	7.70	12	6	4	3	4	3	33	115	104	25	2				4	2	3	1	1	4	1	2	329	197	526					
2H-5, 5-7	7.99		1	1					1	6												1		10	3	13		Rare			
2H-5, 35-37	8.27	7	17	2	25	1	15	231	2							4	1		1			3			309	266	575				
2H-5, 65-67	8.56	6	10	3	2	10	1	3	26	93	97	42	7			3	15	8				1	8	2	1	1	339	112	451		
2H-5, 95-97	8.85	10	9	2	15	4	1	28	143	54	49					1	2	8				7			2	335	308	643			
2H-5, 125-127	9.13																								0	0	0	Barren			
2H-6, 5-7	9.42																								0	0	0	Barren			
2H-6, 35-37	9.70																								0	0	0	Rare			
2H-6, 65-67	9.99	7	10	10	87	2	20	163	2	1						11	2		1			1			317	40	357				
2H-6, 95-97	10.28	4	5	1	71	3	6	237								1	4					1			333	32	365				

Table T2 (continued).

Core, section, interval (cm)	Average depth (mbsf)	<i>Calcidiscus leptopus</i>	<i>Ceratolithus cristatus</i>	<i>Coccolithus pelagicus</i>	<i>Emiliania huxleyi</i>	<i>Gephyrocapsa caribeanica</i> (large)	<i>Gephyrocapsa caribeanica</i> (medium)	<i>Gephyrocapsa caribeanica</i> (small)	<i>Gephyrocapsa muellerae</i>	<i>Gephyrocapsa oceanica</i> (large)	<i>Gephyrocapsa oceanica</i> (medium)	<i>Gephyrocapsa</i> (small)	<i>Gephyrocapsa</i> (very small)	<i>Helicosphaera carteri</i>	<i>Helicosphaera inversa</i>	<i>Neosphaera cocolithomorpha</i>	<i>Pseudoeumiliaria acunosa</i>	Placolith (medium)	Placolith (small)	Placolith (very small)	Sphenolithus spp.	<i>Syracosphaera histrica</i>	<i>Syracosphaera pulchra</i>	<i>Syracosphaera</i> spp.	<i>Umbilicosphaera irregularis</i>	<i>Umbilicosphaera hulbertiana</i>	Subtotal	<i>Florisphaera profunda</i>	Total	Remarks	
2H-6, 125–127	10.56																														
2H-7, 5–7	10.85	5	1			204	11		4	103							1	2	3							0	0	0	Barren		
2H-7, 35–37	11.14																														Barren
2H-7, 64–66	11.41	2	3	1		252	7	1	26			1						3		1						300	19	319			
2H-CC, 5–7	11.49	2	2		1	178	74	2	2	42		3						2	3	2					319	24	343				
3H-1, 5–7	11.76	2	3			172	71		2	58		4	5						5							325	31	356			
3H-1, 35–37	12.05	3	3			177	35	2	1	96		4							1							322	38	360			
3H-1, 65–67	12.34	2	3			245	41		29			4							4							329	27	356			
3H-1, 95–97	12.63	2	12			68	13		3	211			1						12	4						327	73	400			
3H-1, 125–127	12.92																									0	0	0	Barren		
3H-2, 5–7	13.21																									0	0	0	Barren		
3H-2, 35–37	13.50																									0	0	0	Barren		
3H-2, 65–67	13.80																									0	0	0	Barren		
3H-2, 95–97	14.09																									0	0	0	Barren		
3H-2, 125–127	14.38																									0	0	0	Rare		
3H-3, 5–7	14.67	1	4	3		121	12		5	140			1						4	1						300	107	407			
3H-3, 35–37	14.96																									0	0	0	Rare		
3H-3, 65–67	15.25	2				245	25		9	16		3						4							308	17	325				
3H-3, 95–97	15.54	1				104	24		1	170			1					1							302	89	391				
3H-3, 125–127	15.83																									0	0	0	Barren		
3H-4, 5–7	16.13																									0	0	0	Barren		
3H-4, 35–37	16.42	3				162	39	1	74			4	2					16	5	4					309	37	346				
3H-4, 65–67	16.71	1	2	3		203	31	3	24	26		22	4				1	2	15	4				346	28	374					
3H-4, 95–97	17.00	3				194	71	3	2	5		26	9					6	1						322	16	338				
3H-4, 125–127	17.29		1	1		204	62		40			4						3	4						324	7	331				
3H-5, 5–7	17.58	5				227	43		31			4	1					3							317	11	328				
3H-5, 35–37	17.87					198	70		48			8						10	2						338	18	356				
3H-5, 65–67	18.16	3				201	58		52	18		1						9	4	1					348	22	370				
3H-5, 95–97	18.45	1				221	49		21	11								8	3						322	16	338				
3H-5, 125–127	18.75	2		1		208	45		44	10	1							12	4						329	18	347				
3H-6, 5–7	19.04	1				194	78		44	12								4	1						338	23	361				
3H-6, 35–37	19.33	1	1			208	53		53	13								6	1						337	19	356				
3H-6, 65–67	19.62	3	3			186	48		76	16		1						17	2	1					354	30	384				
3H-6, 95–97	19.91	2	1			211	42		55	6								5	3						329	24	353				
3H-6, 125–127	20.20	3	2			208	41		52	5								3	1						316	25	341				
3H-7, 5–7	20.49	1	4			198	41		66	4	1							9	3						331	20	351				
3H-7, 35–37	20.78	2	1	7		162	34		77	1		1						24	1						321	22	343				
3H-CC, 5–7	21.01		8			192	63		29	3								16							322	20	342				